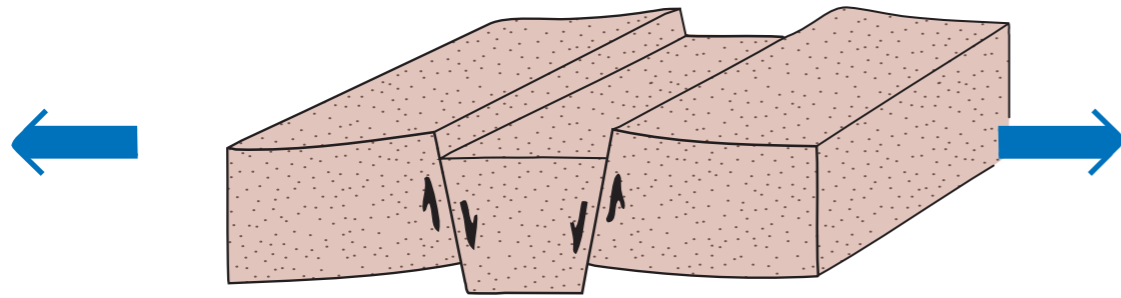


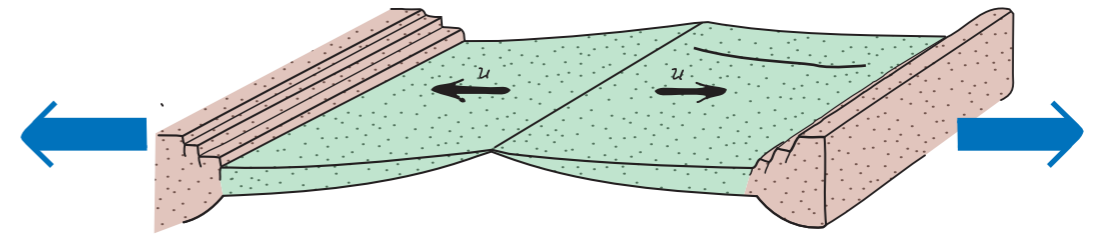
14- Obduction/Kontinental Kruste

Wilson Cycle

Rifting

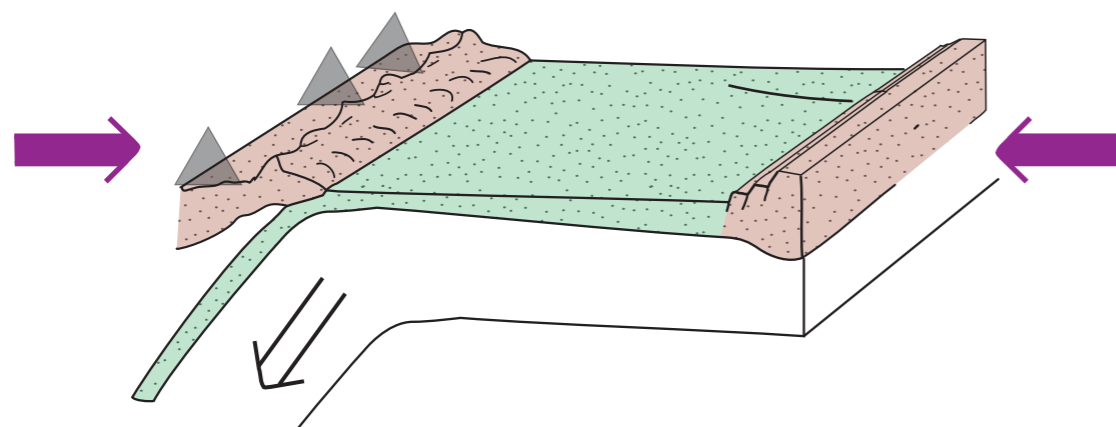


Ocean formation

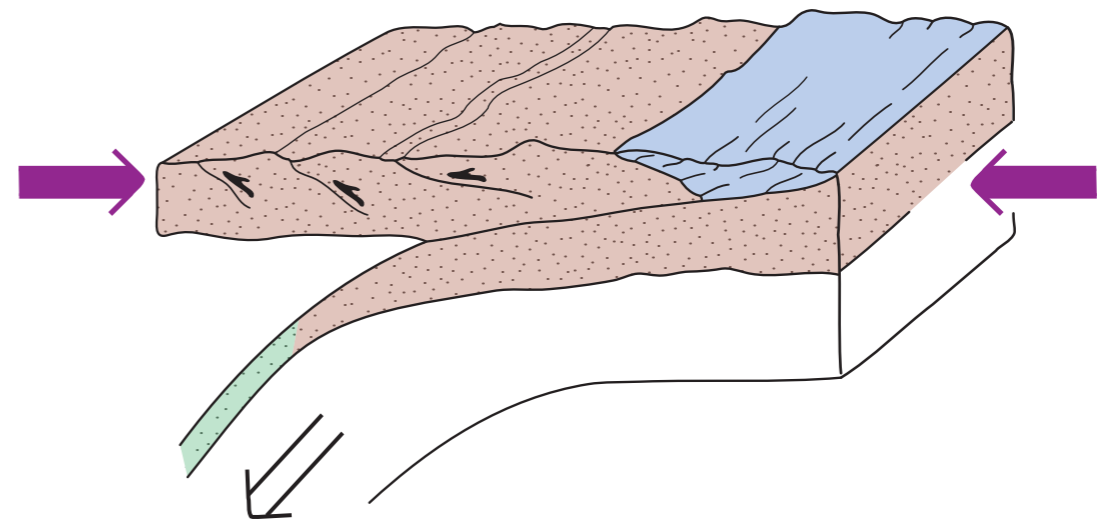


Obduction?

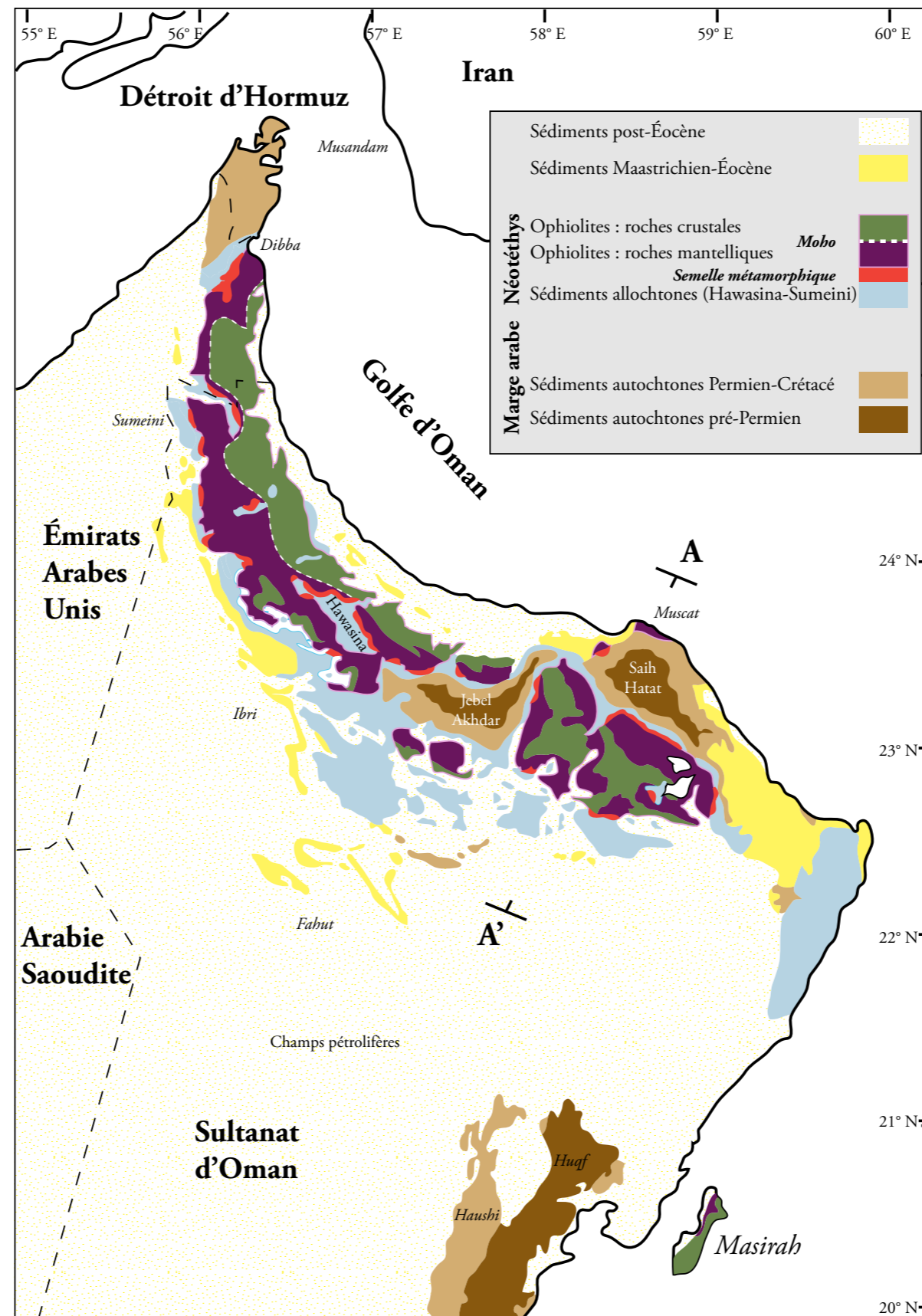
Subduction



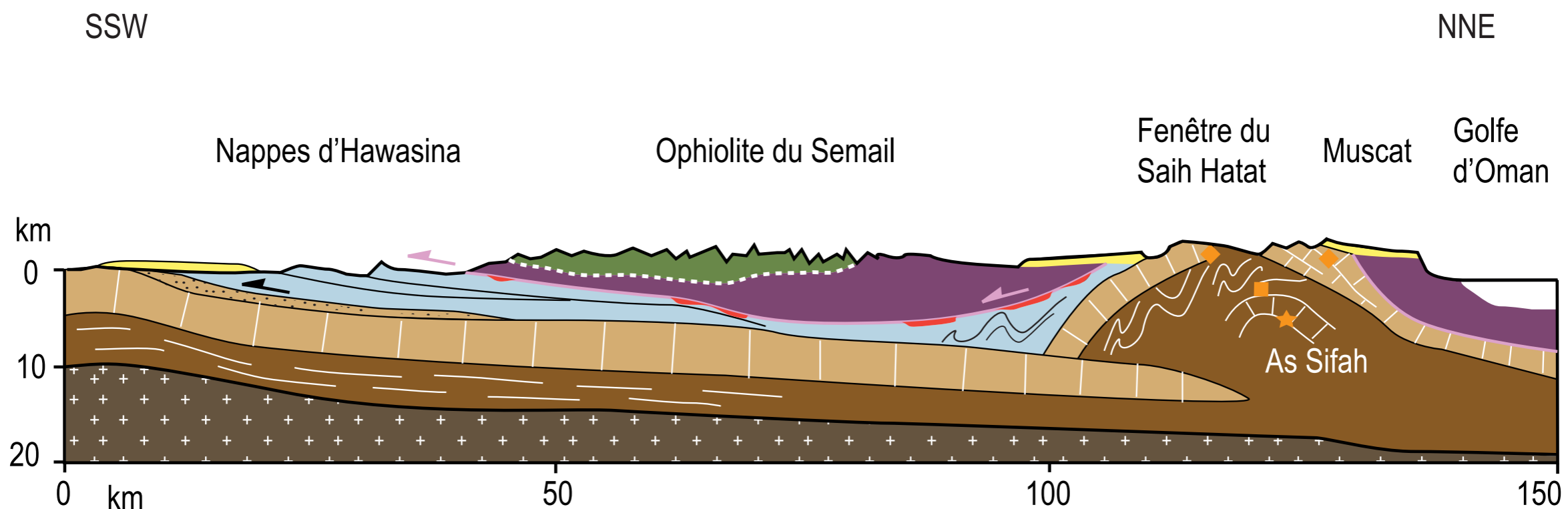
Collision



Oman, the best obducted ophiolite



Oman, the best obducted ophiolite



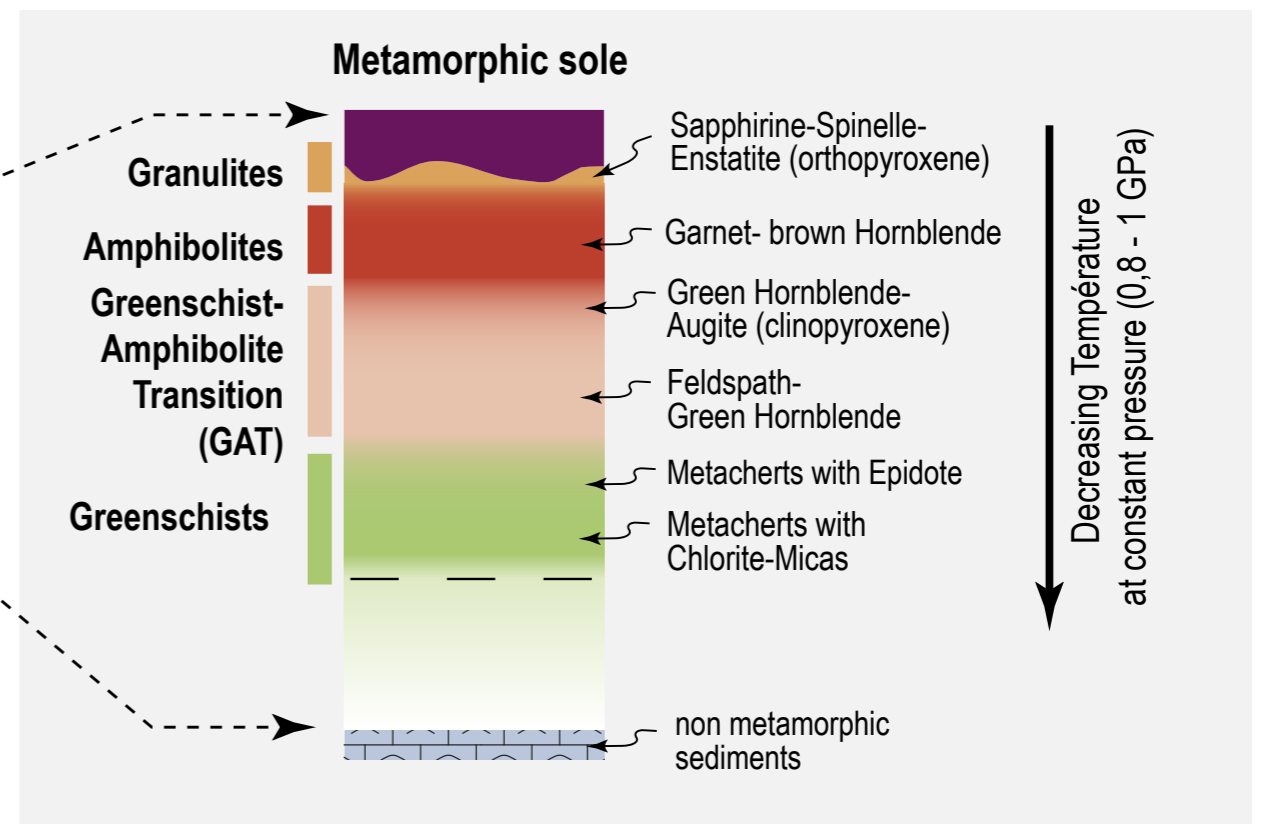
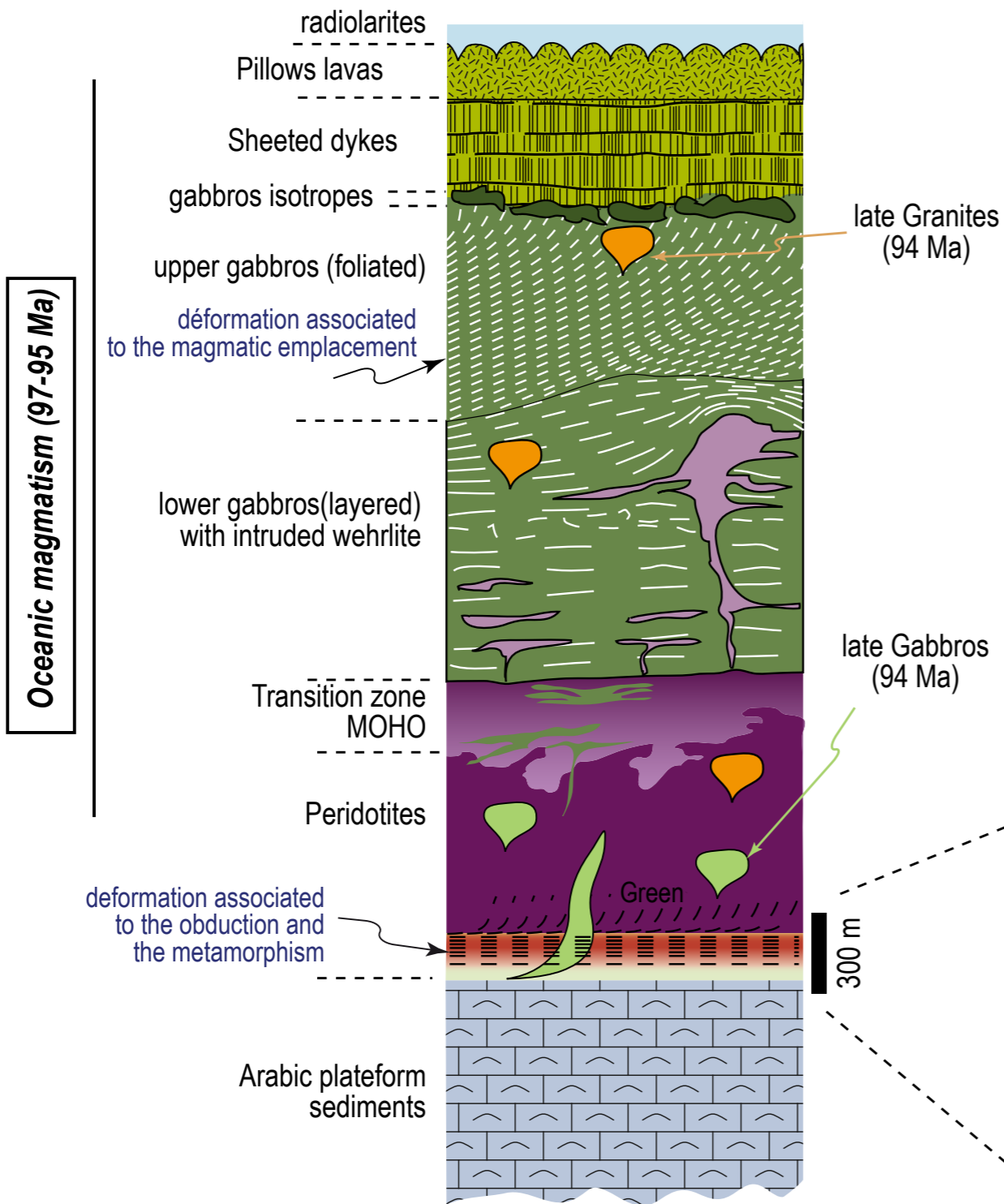
Plateforme arabe

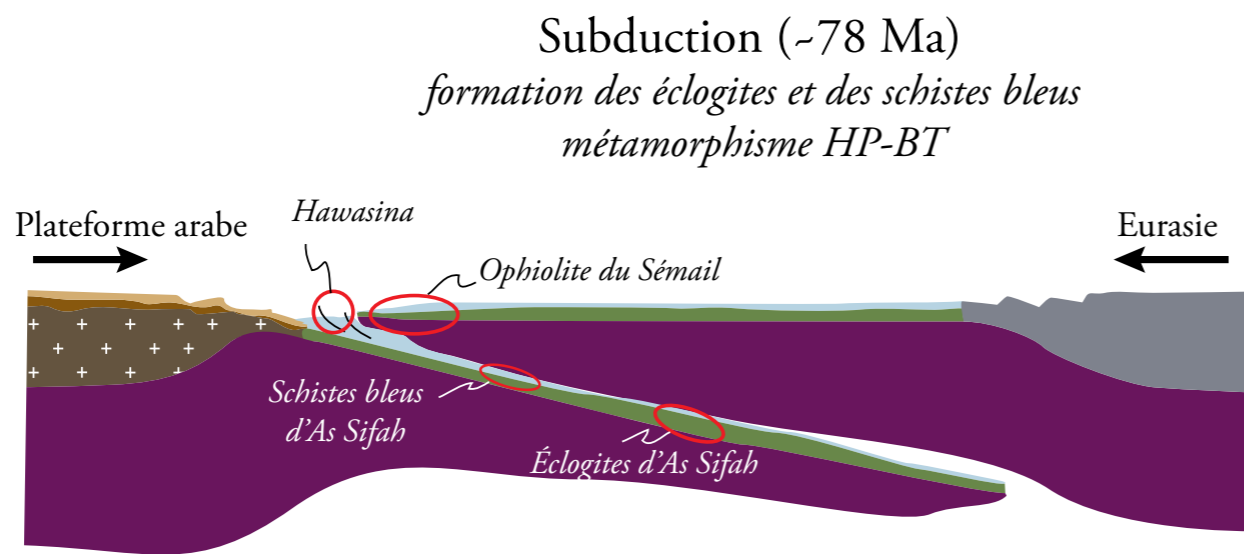
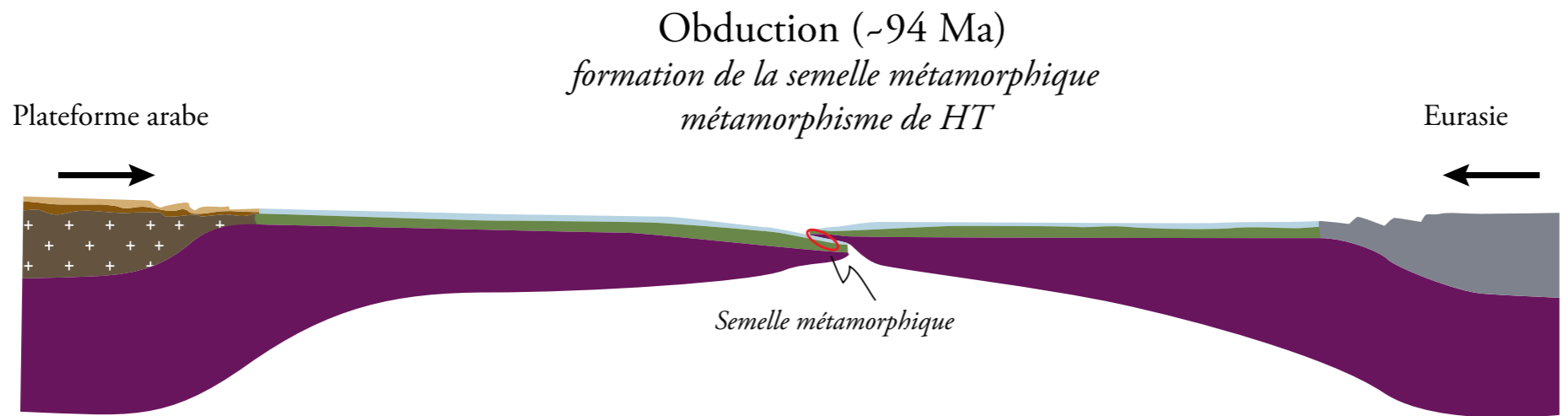
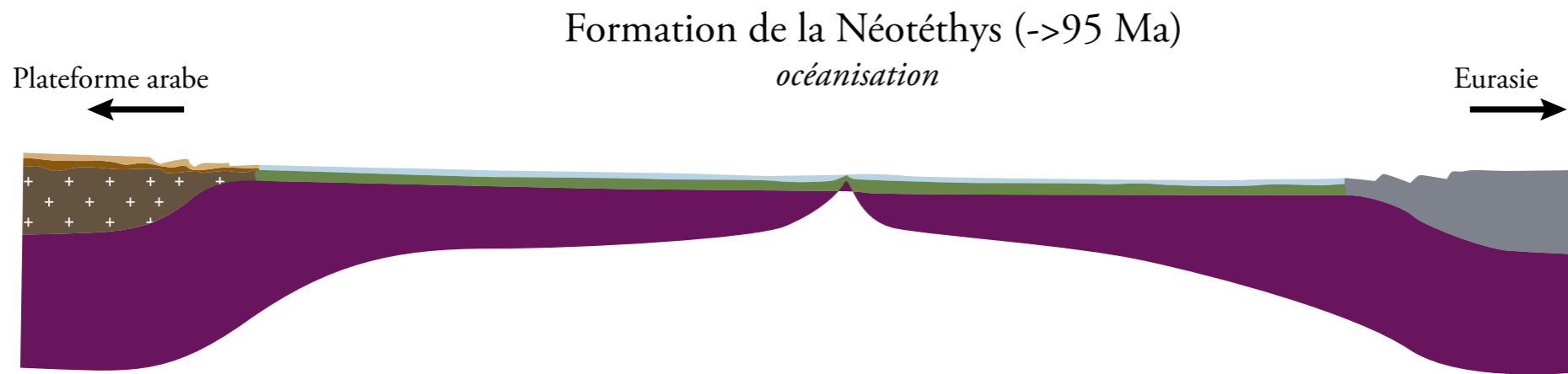
- Roches sédimentaires Permien-Crétacé supérieur
- Roches sédimentaires Protérozoïque supérieur-Paléozoïque
- Croûte continentale arabe

Ophiolite de Semail

- Croûte (Gabbros, Basaltes)
- Paléo-Moho
- Manteau (Hazburgites)

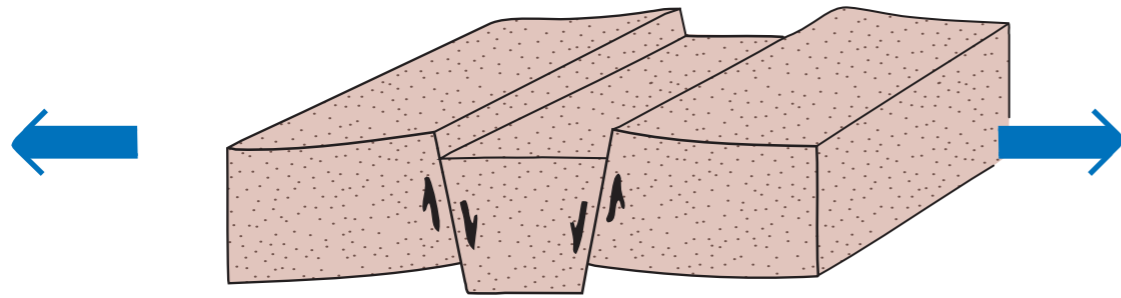
- Sédiments océaniques (nappes de l'Hawasina)
- Tertiaire discordant



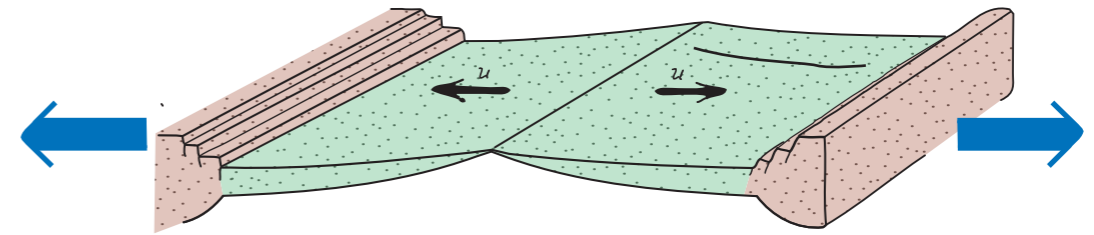


Wilson Cycle

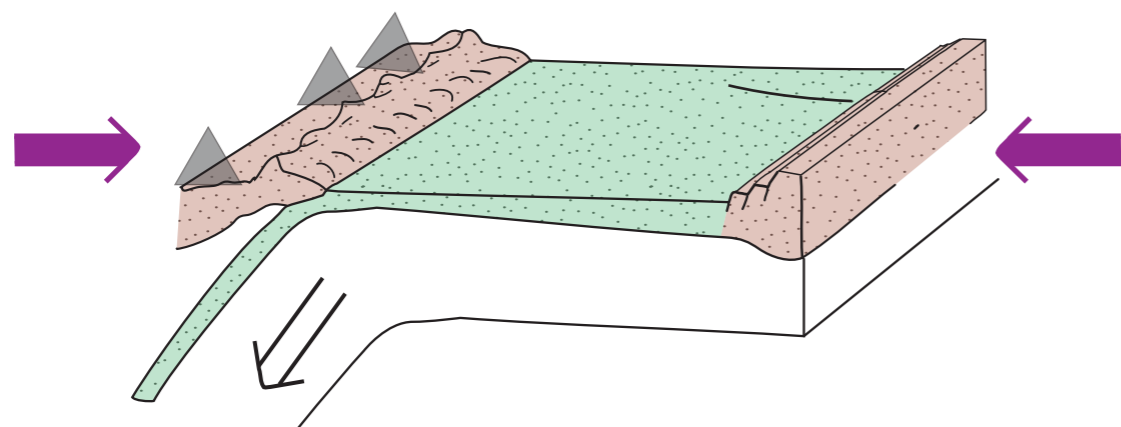
Rifting



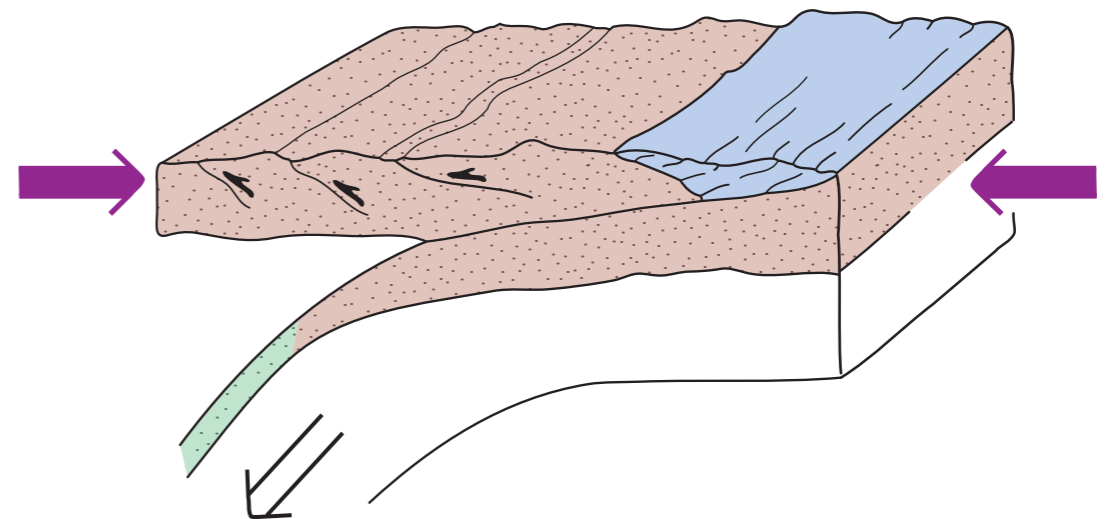
Ocean formation



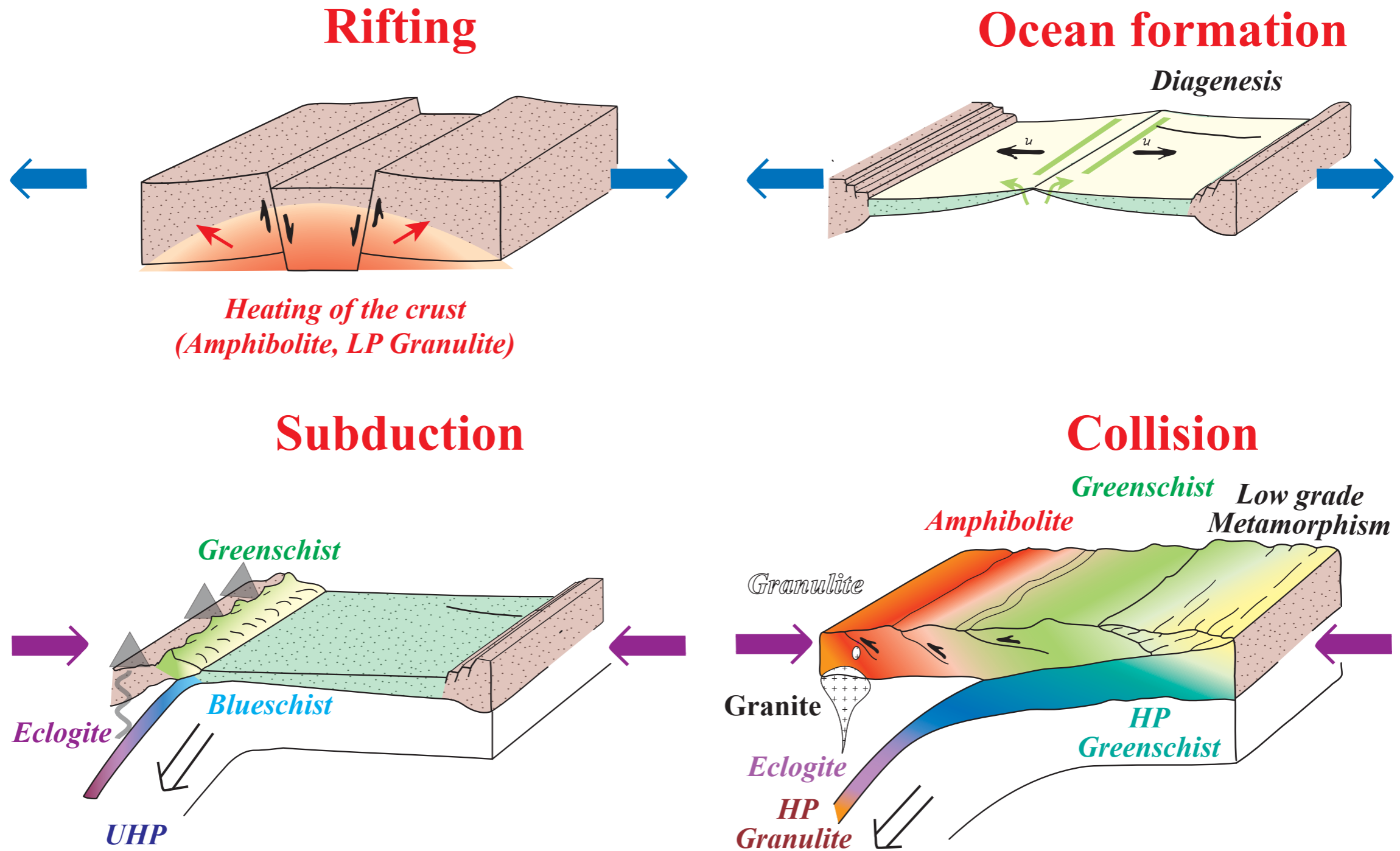
Subduction



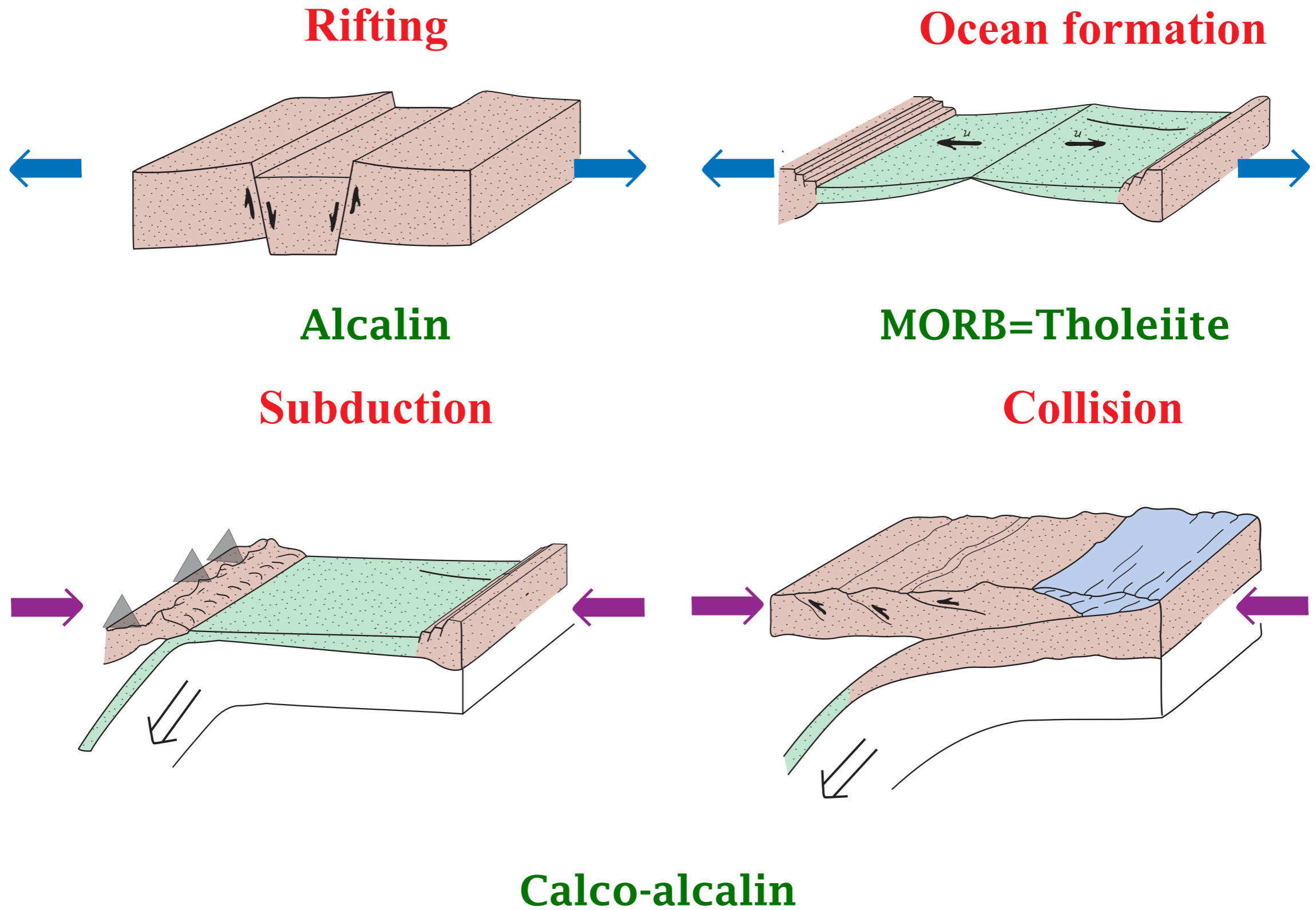
Collision



Wilson Cycle: Metamorphism

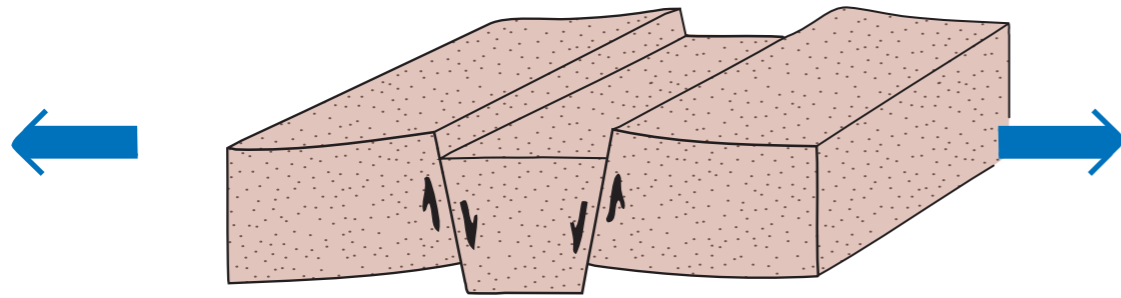


Wilson Cycle: *Volcanism*



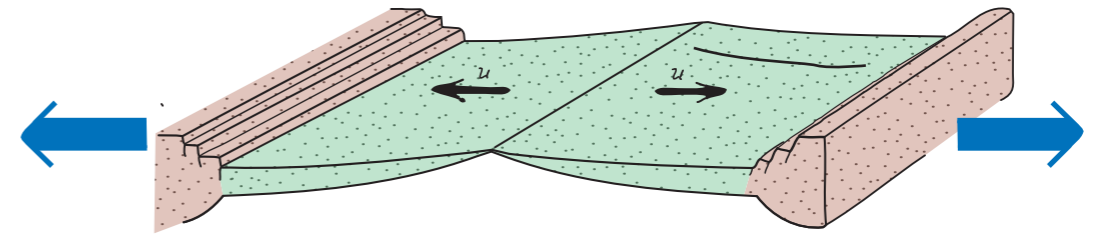
Wilson Cycle: *Sediment*

Rifting



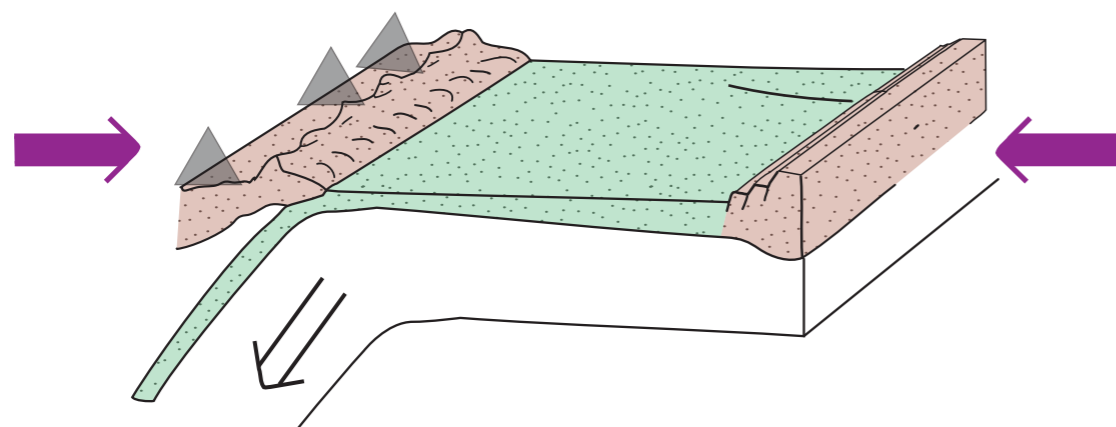
Evaporites, conglomerat

Ocean formation



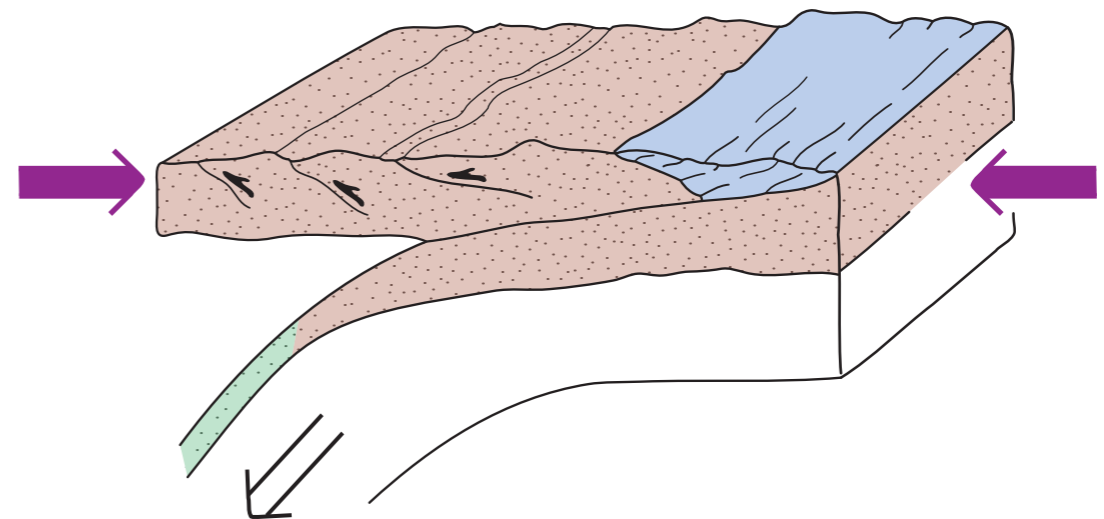
Radiolarite, carbonates

Subduction



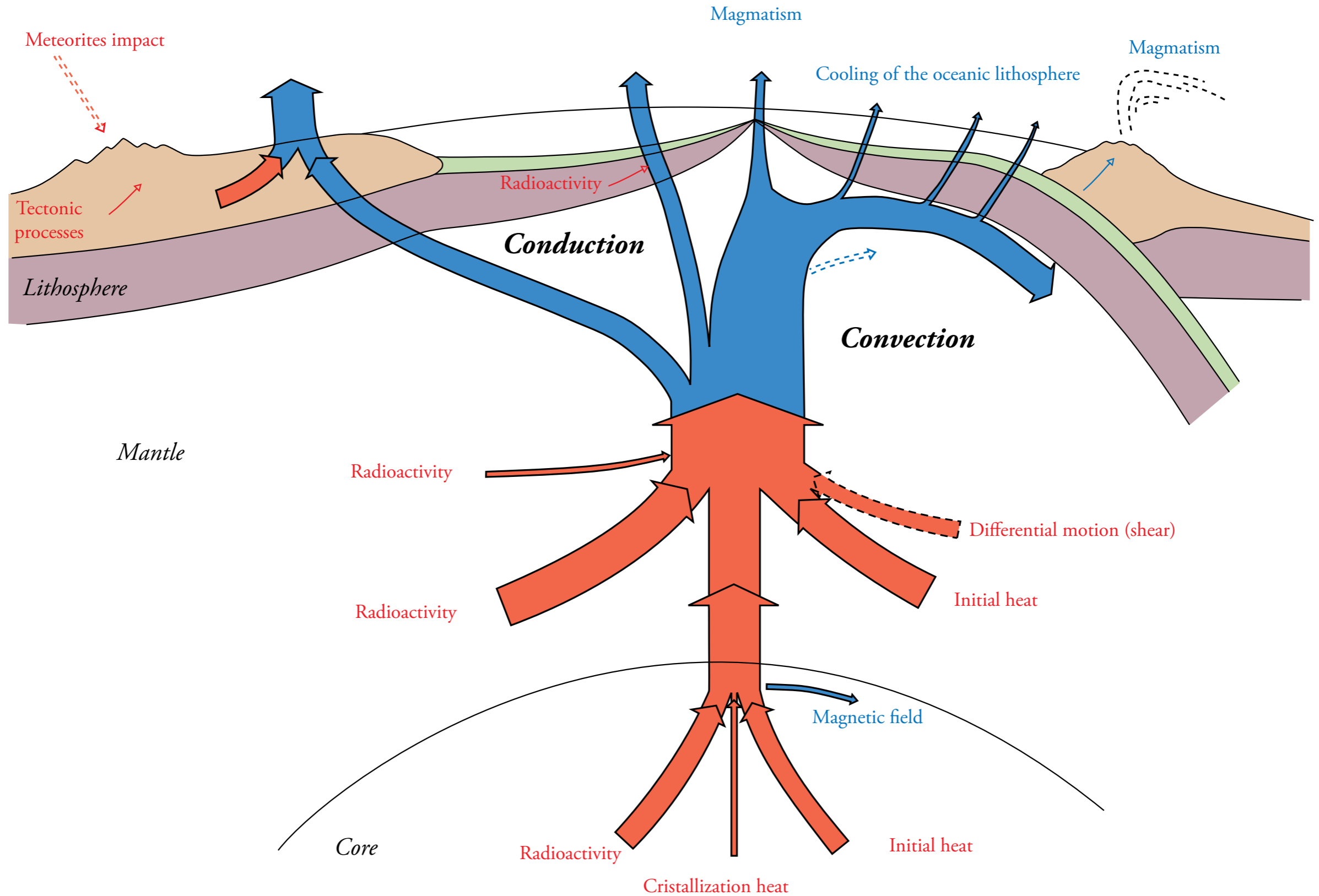
Flysch, turbidites

Collision

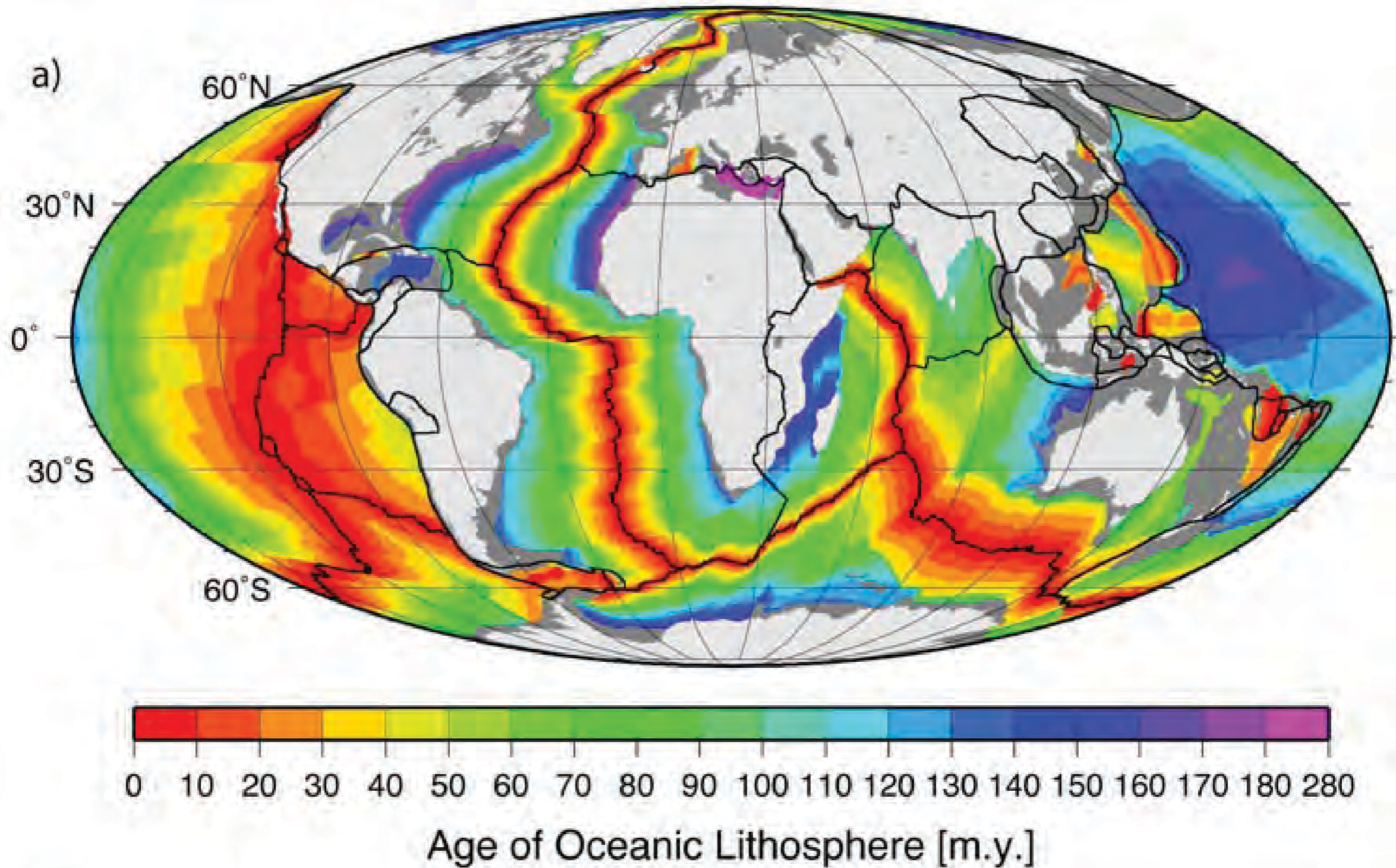


Molasse

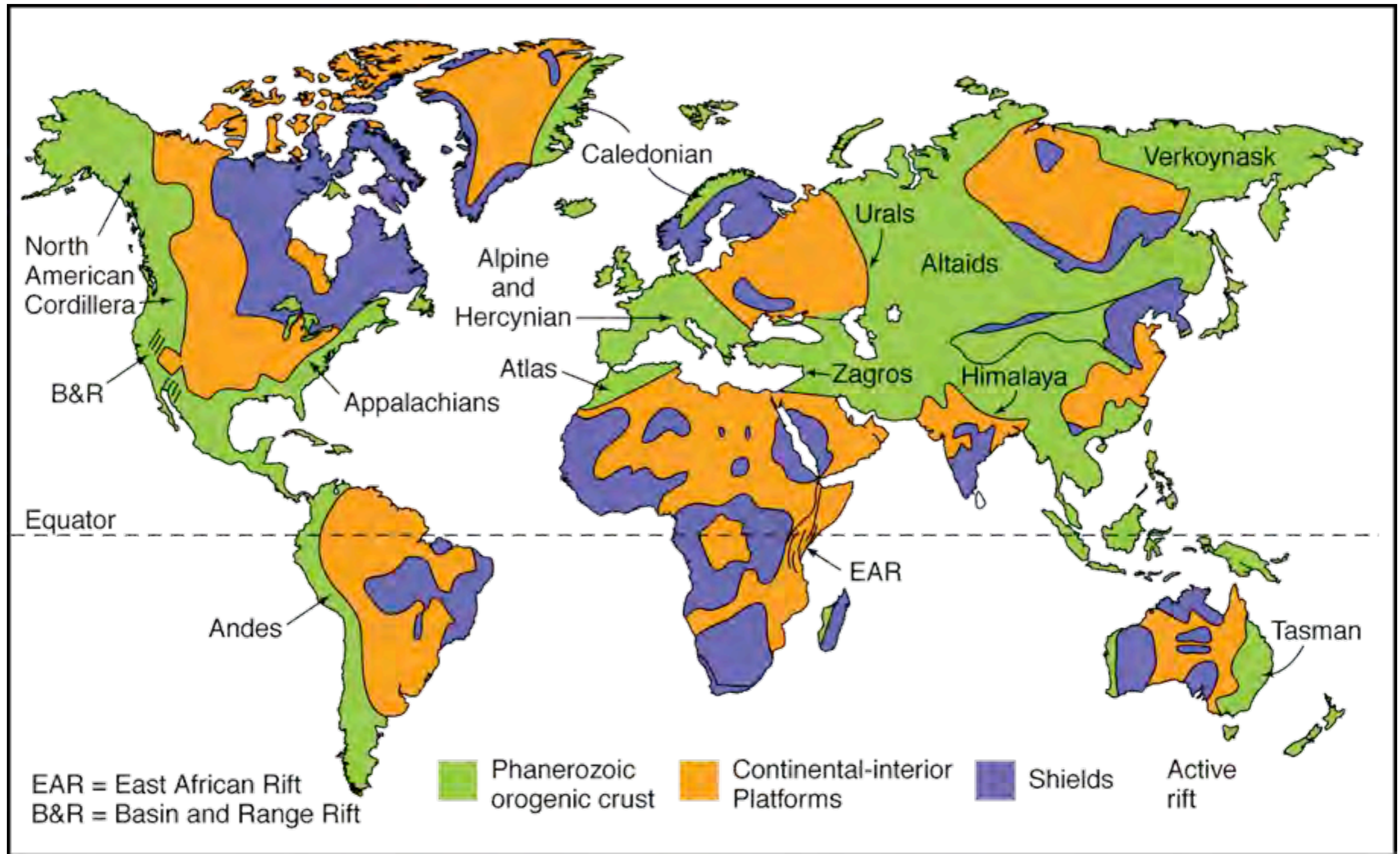
Earth heat flow balance



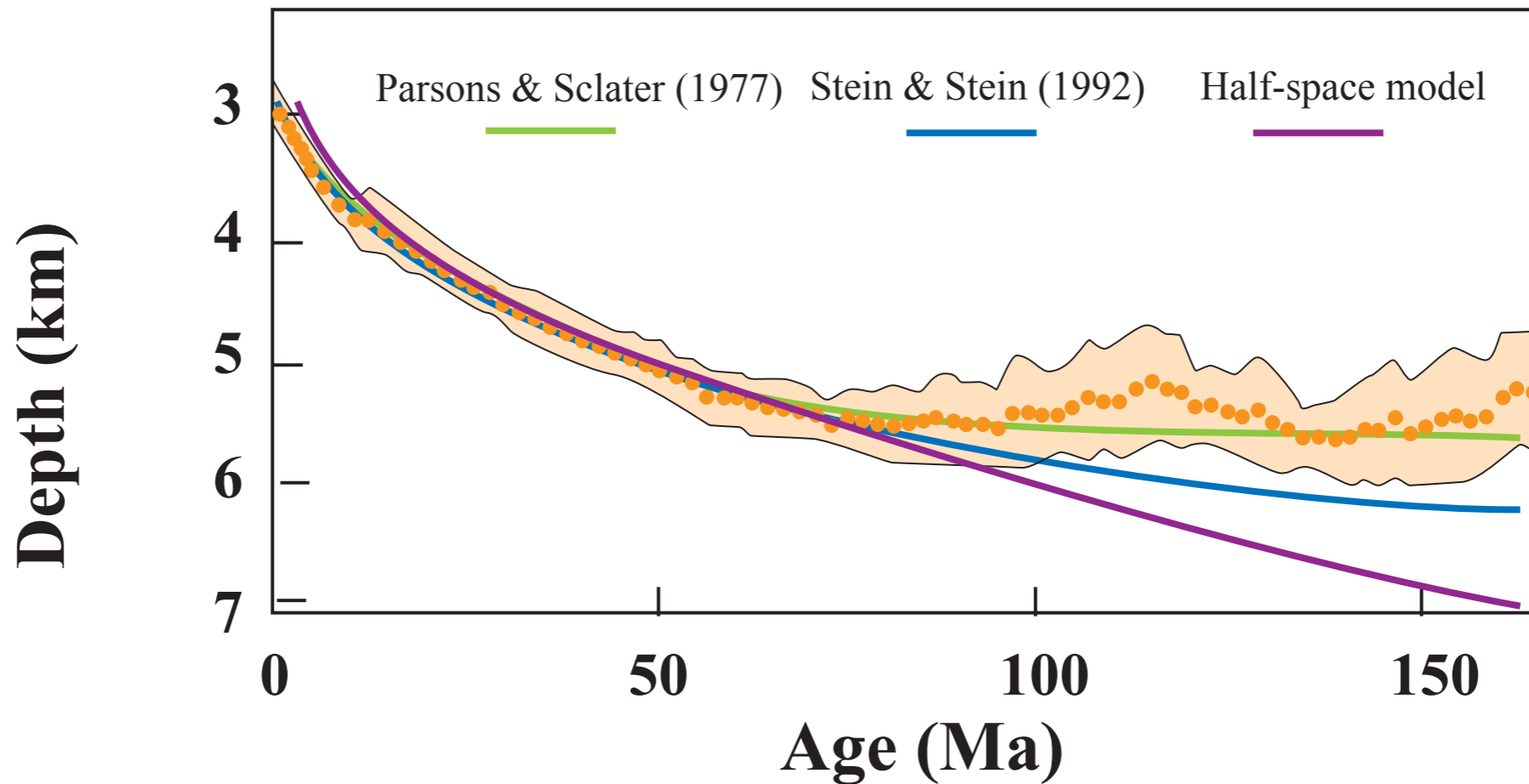
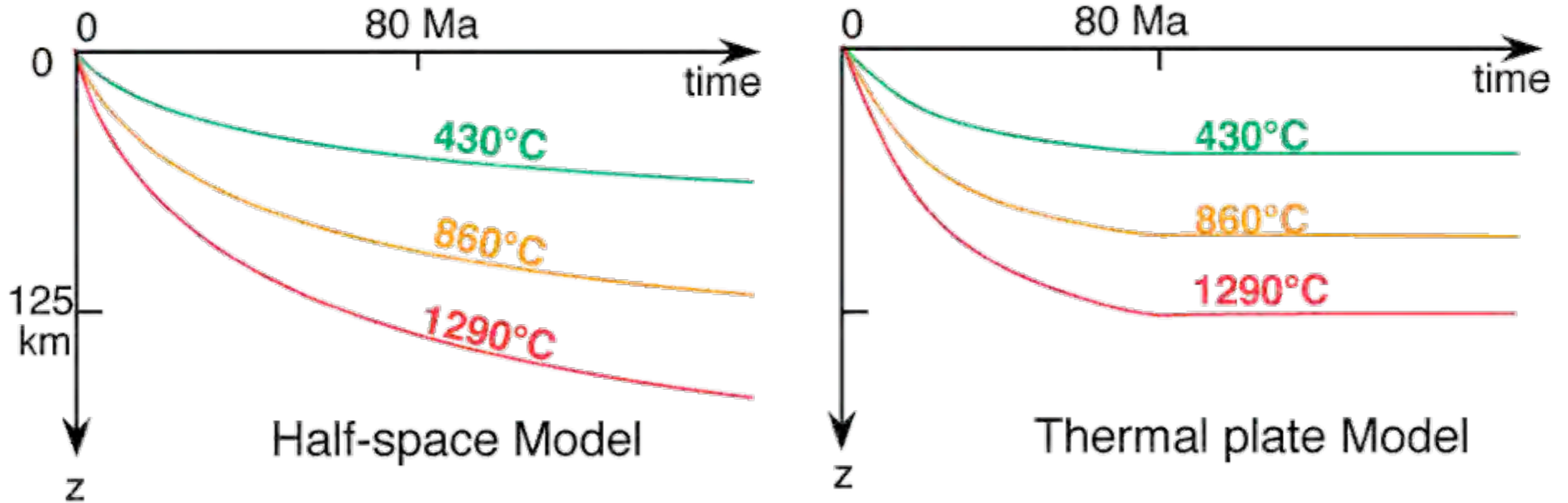
Oceanic crust ages



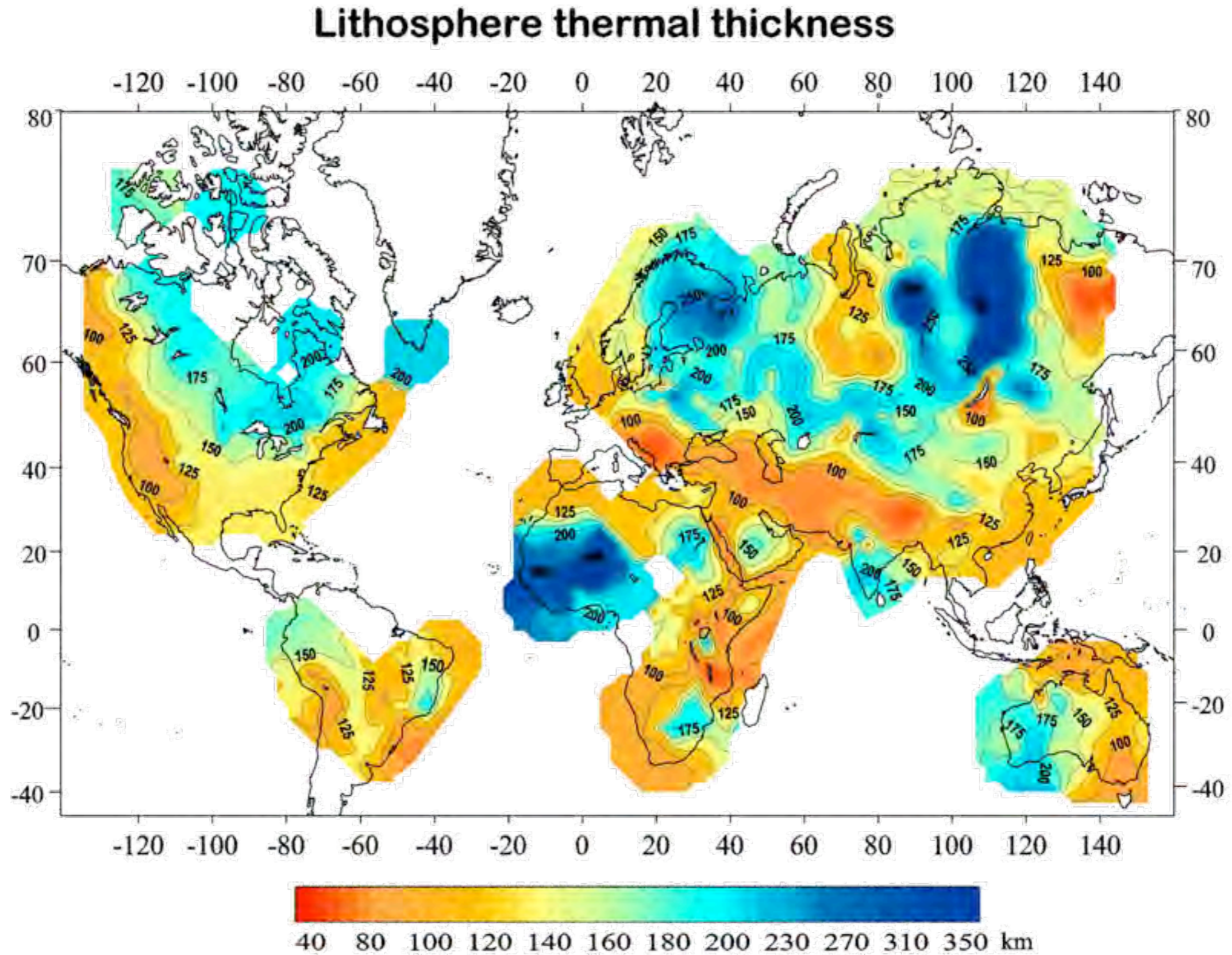
Continental crust ages



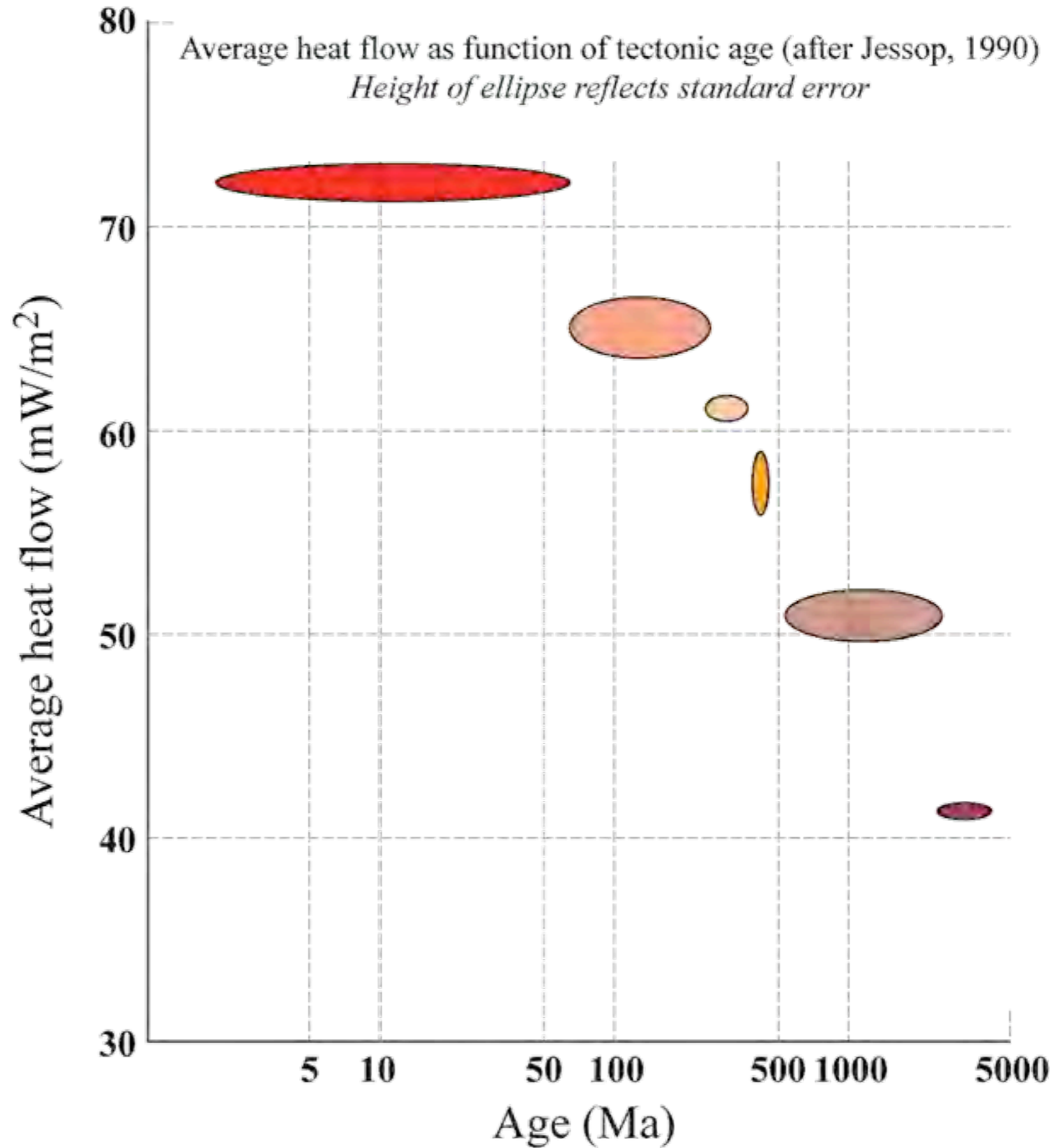
Oceanic lithosphere structure



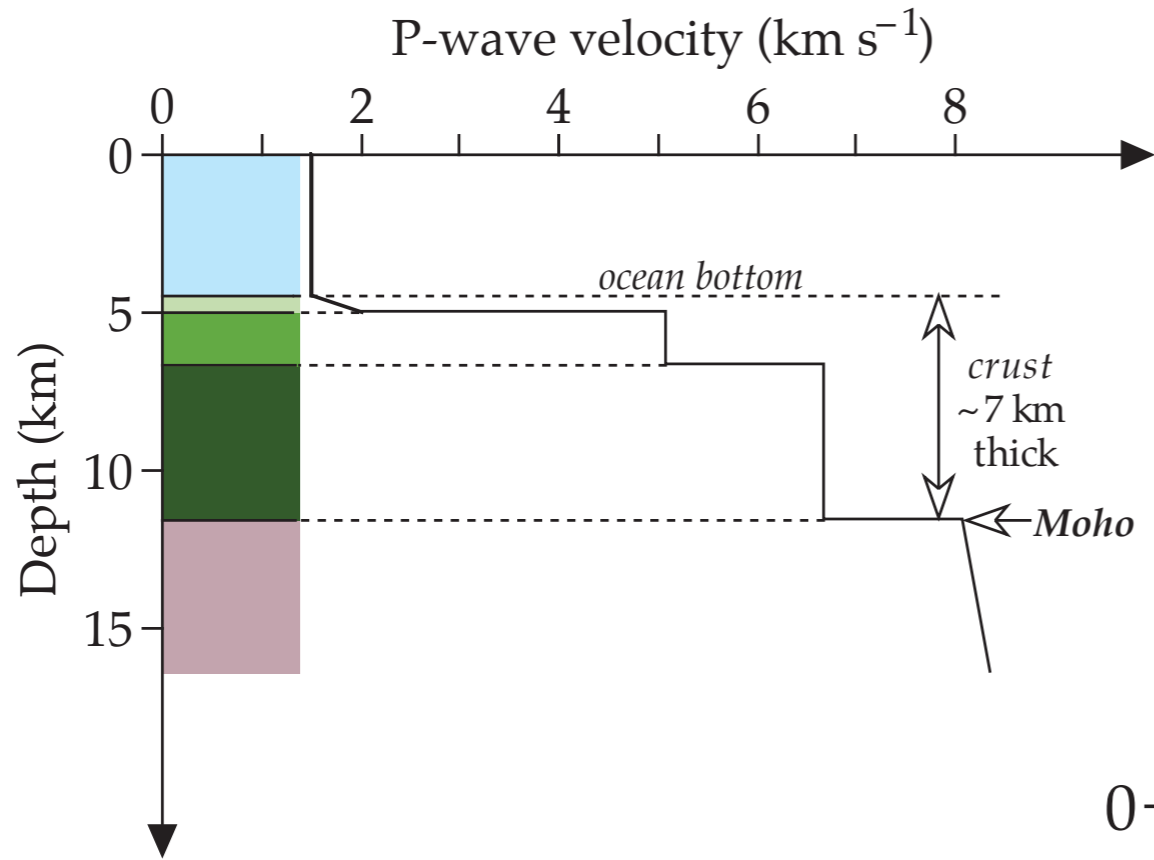
Continental lithosphere structure



Thermal evolution of the continental lithosphere

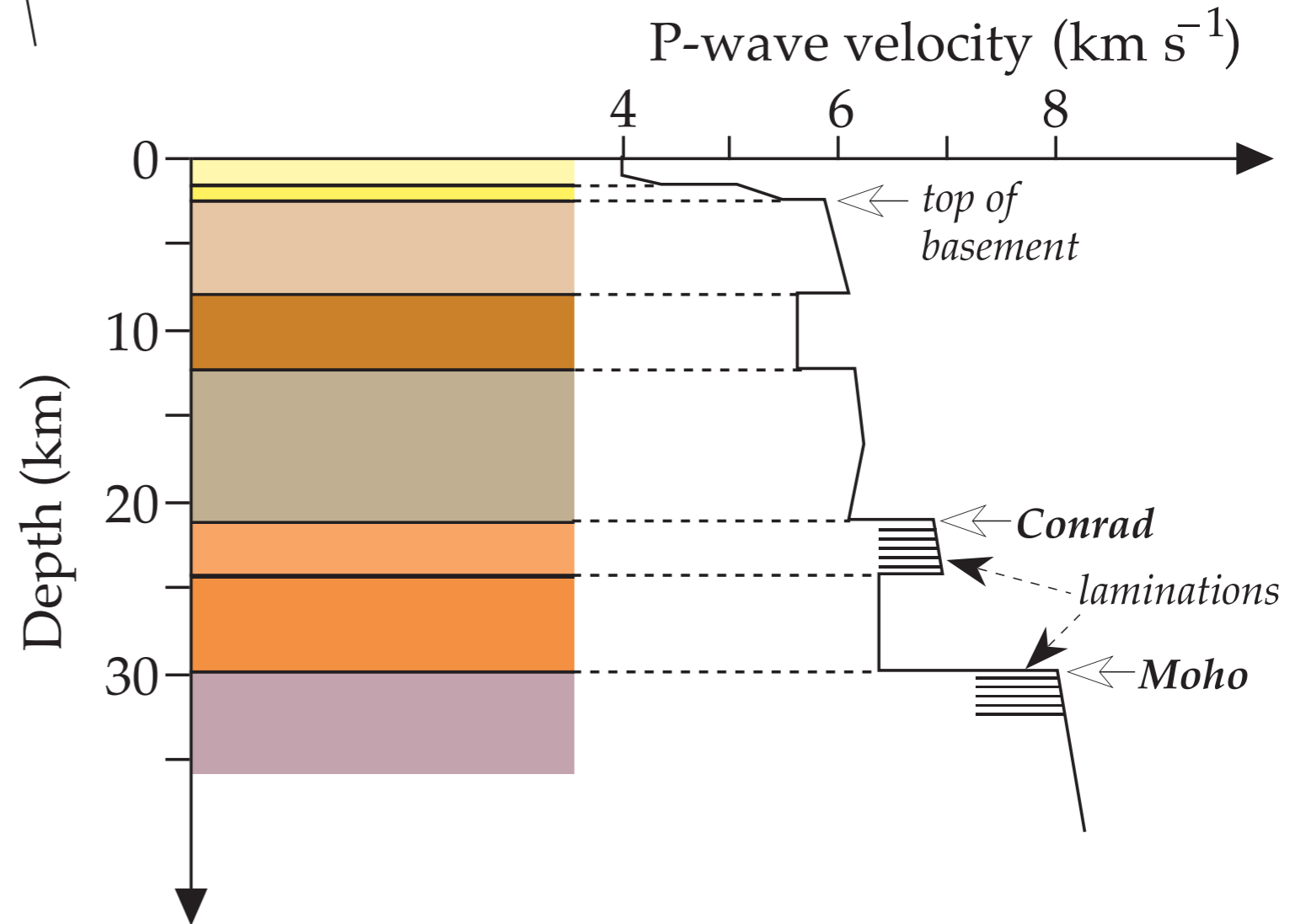


Oceanic vs continental crust



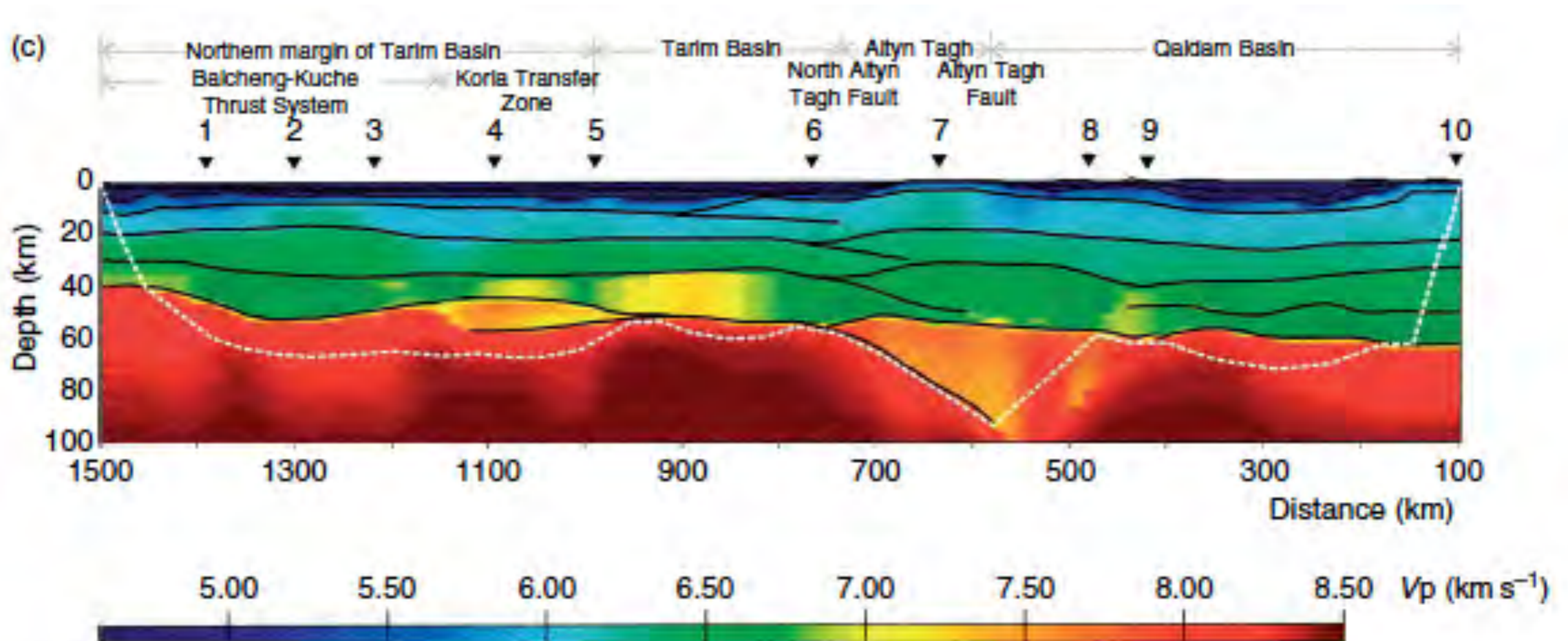
- ocean sea water
- Layer 1 oceanic sediments
- Layer 2 basalt
- Layer 3 gabbro
- upper mantle ultramafics

- near-surface low-velocity layer
 - Cenozoic sediments
 - Mesozoic & Paleozoic sediments
- zone of positive velocity gradient
 - upper crystalline basement
- sialic low-velocity layer
 - granitic layer
- middle crustal layer
 - Gneiss
- high-velocity tooth
 - amphibolites
- lower crustal layer
 - granulites
- uppermost mantle
 - ultramafics



Example of continental crust

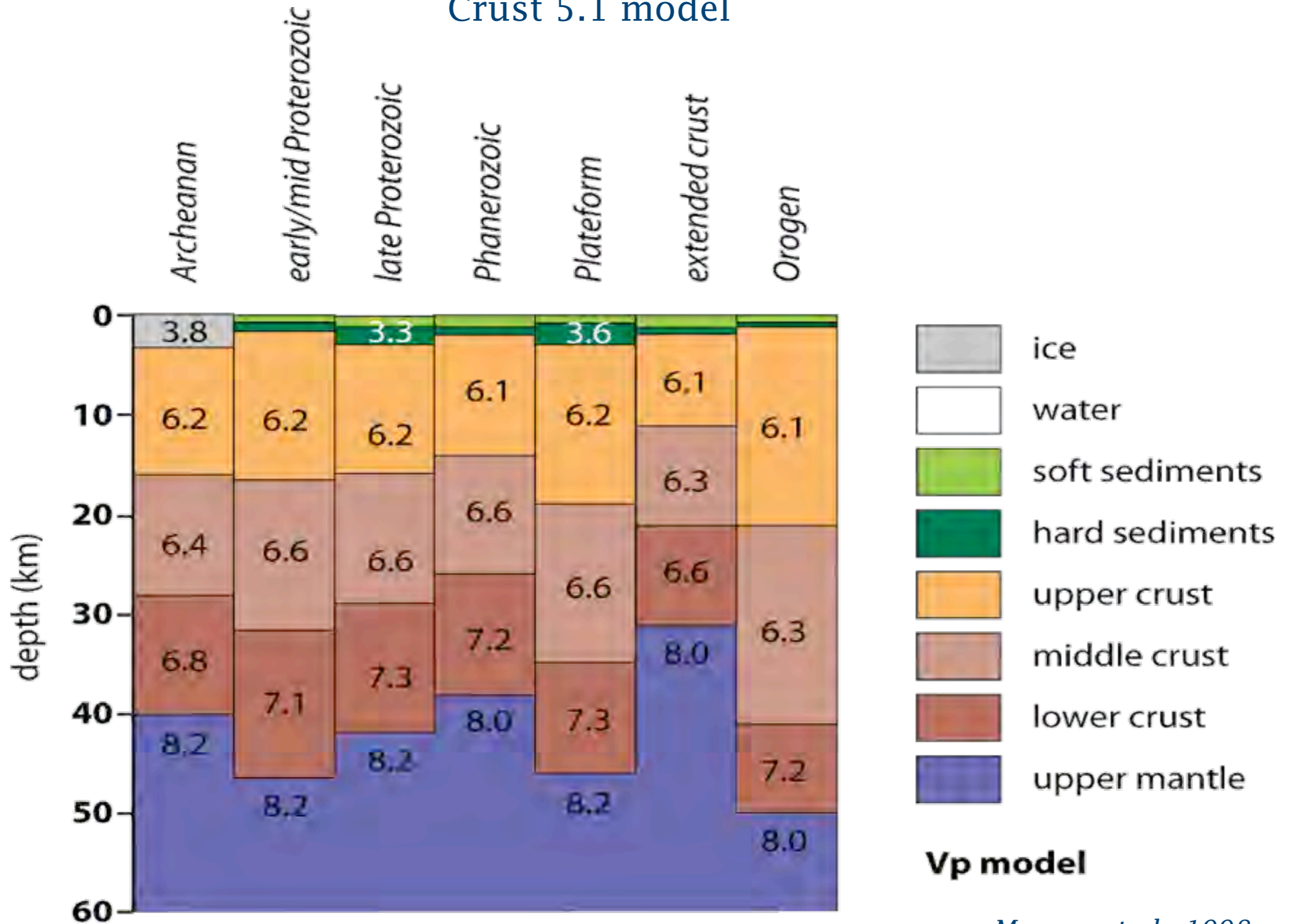
Crustal structure across the Altyn Tagh Range at the northern margin of the Tibetan Plateau



Modeling results along the seismic refraction/wide-angle reflection profile at the northern margin of the Tibetan Plateau. Crustal and uppermost mantle seismic velocity (V_p) model. Both V_p and V_s can be estimated for the crust from these data. P_n velocity ranges from 7.8 to 8.2 km s^{-1} .

Different types of continental crust

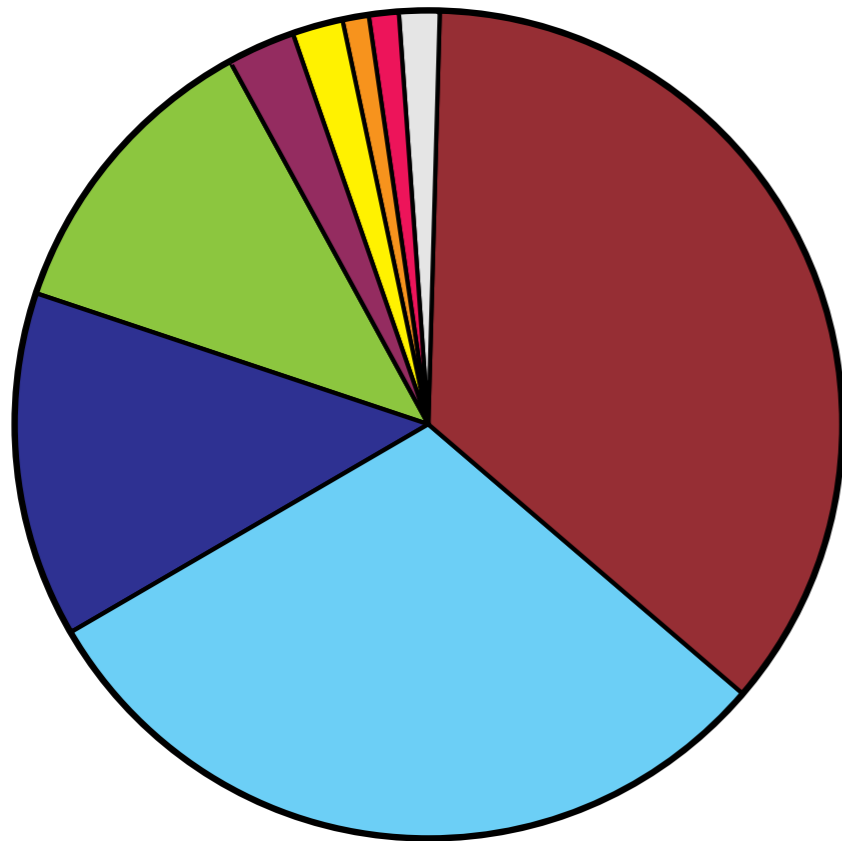
Crust 5.1 model



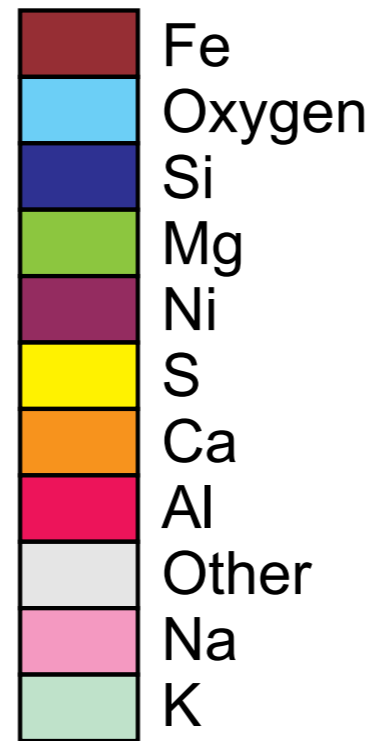
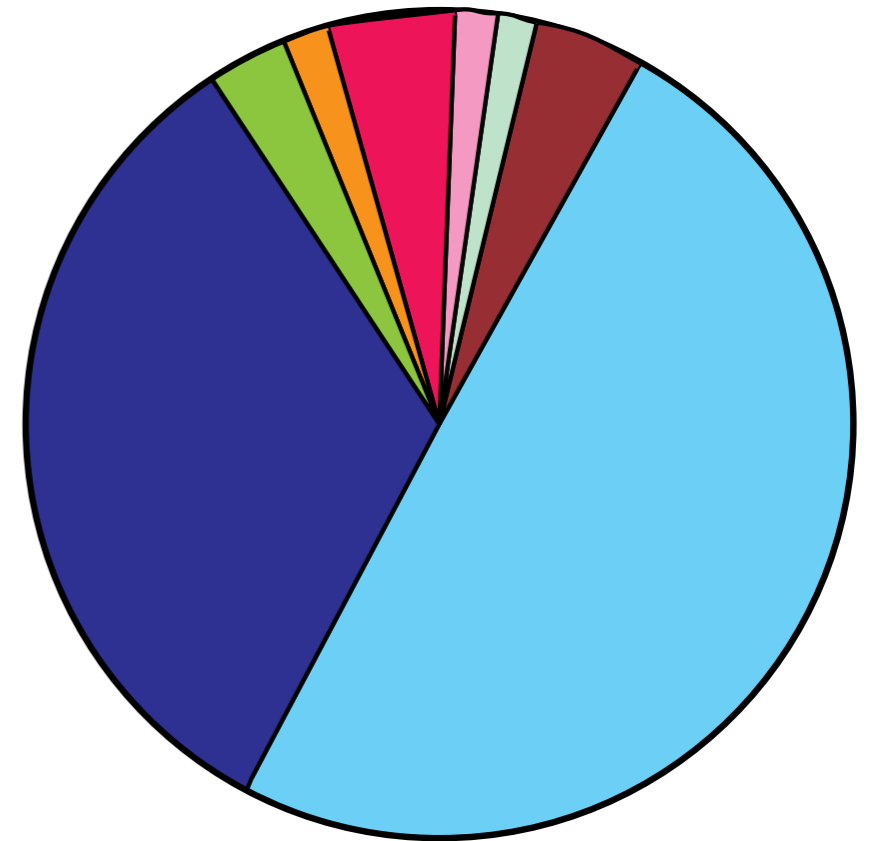
Mooney et al., 1998

Chemical composition of the continental crust

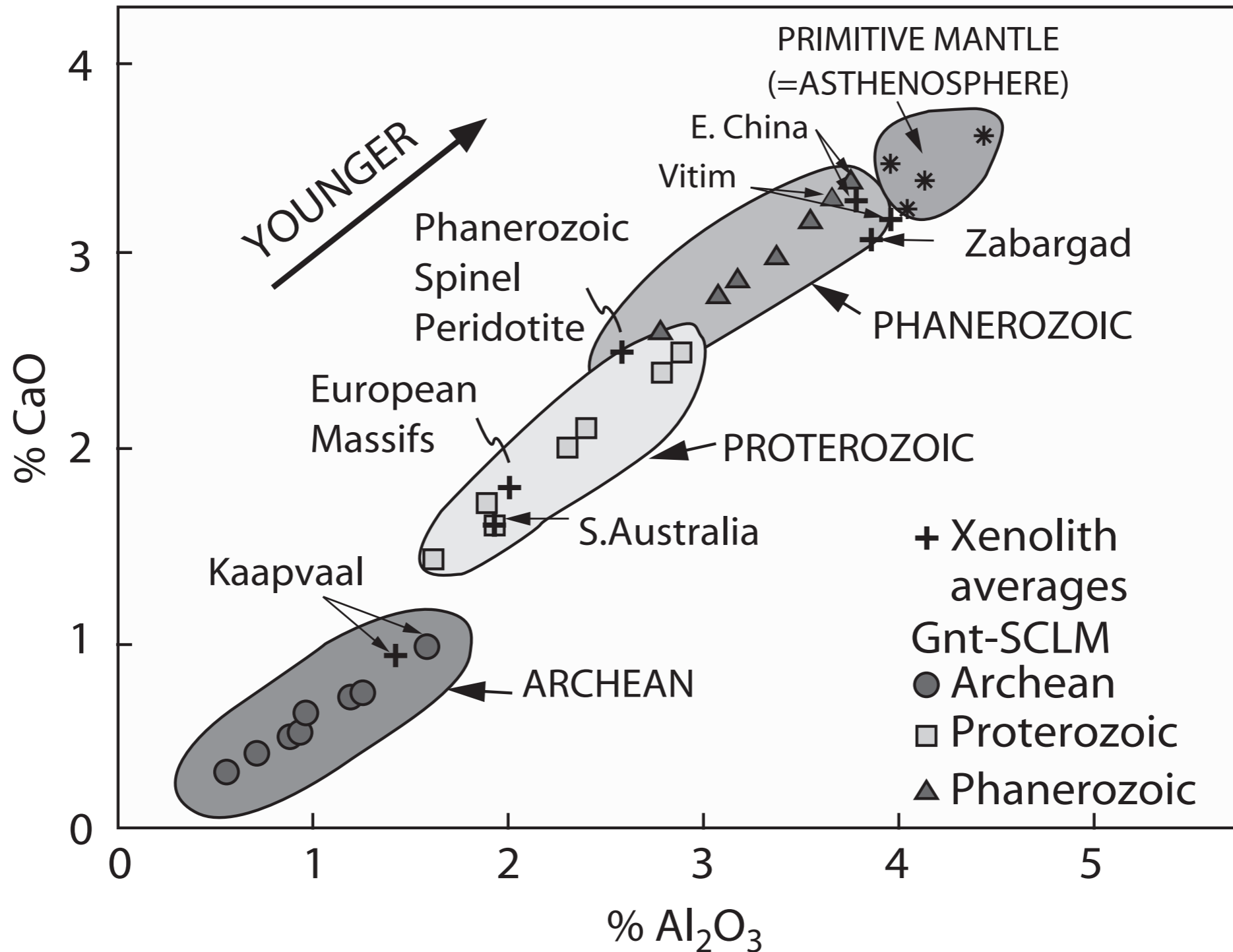
Whole Earth



Earth's crust

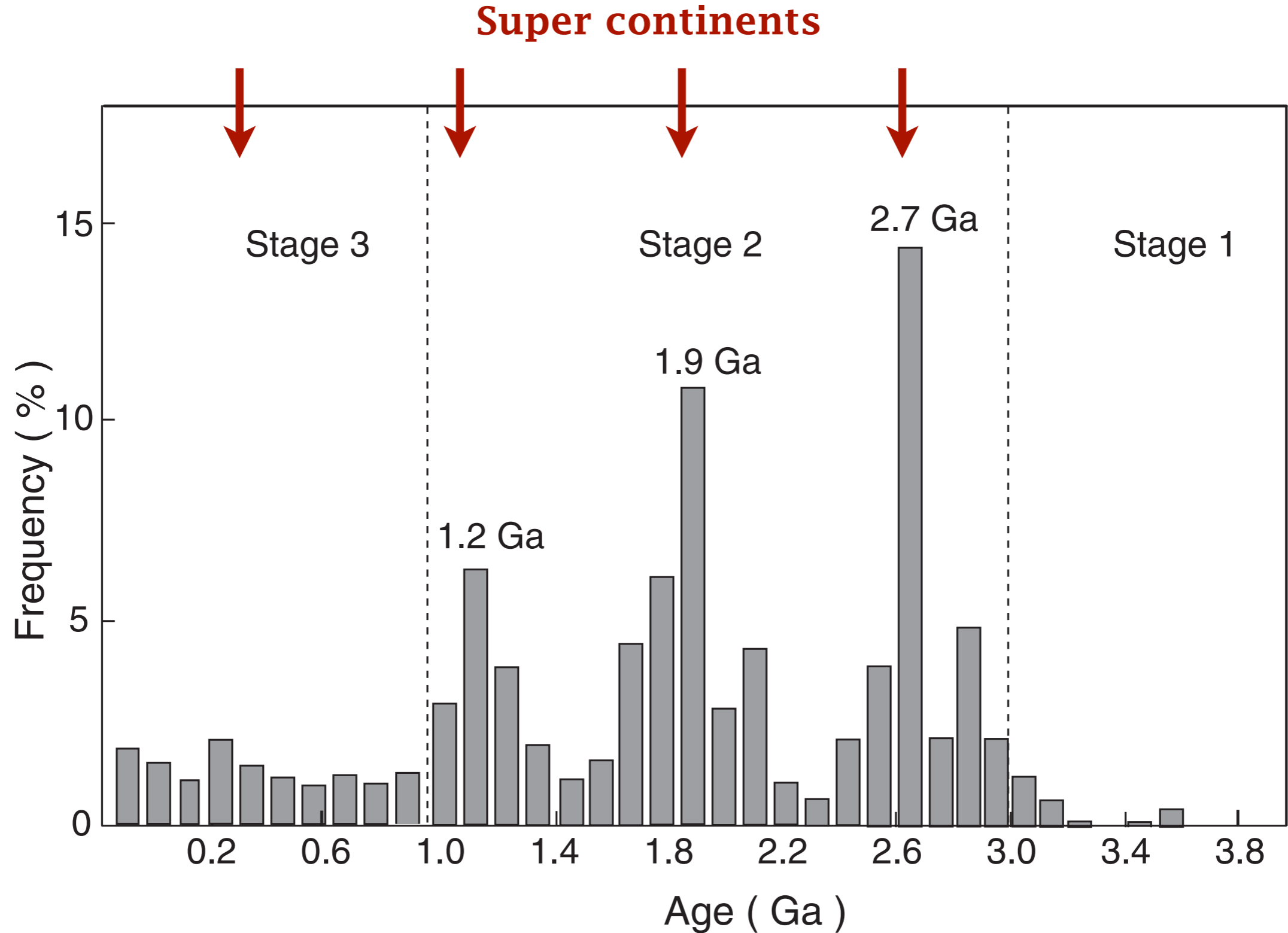


Chemical composition of the continental crust



CaO-Al₂O₃ plot showing the range of subcontinental lithospheric mantle (SCLM) compositions for selected cratons that have been matched with ages of the youngest tectonothermal events in the overlying crust (after O'Reilly et al., 2001)

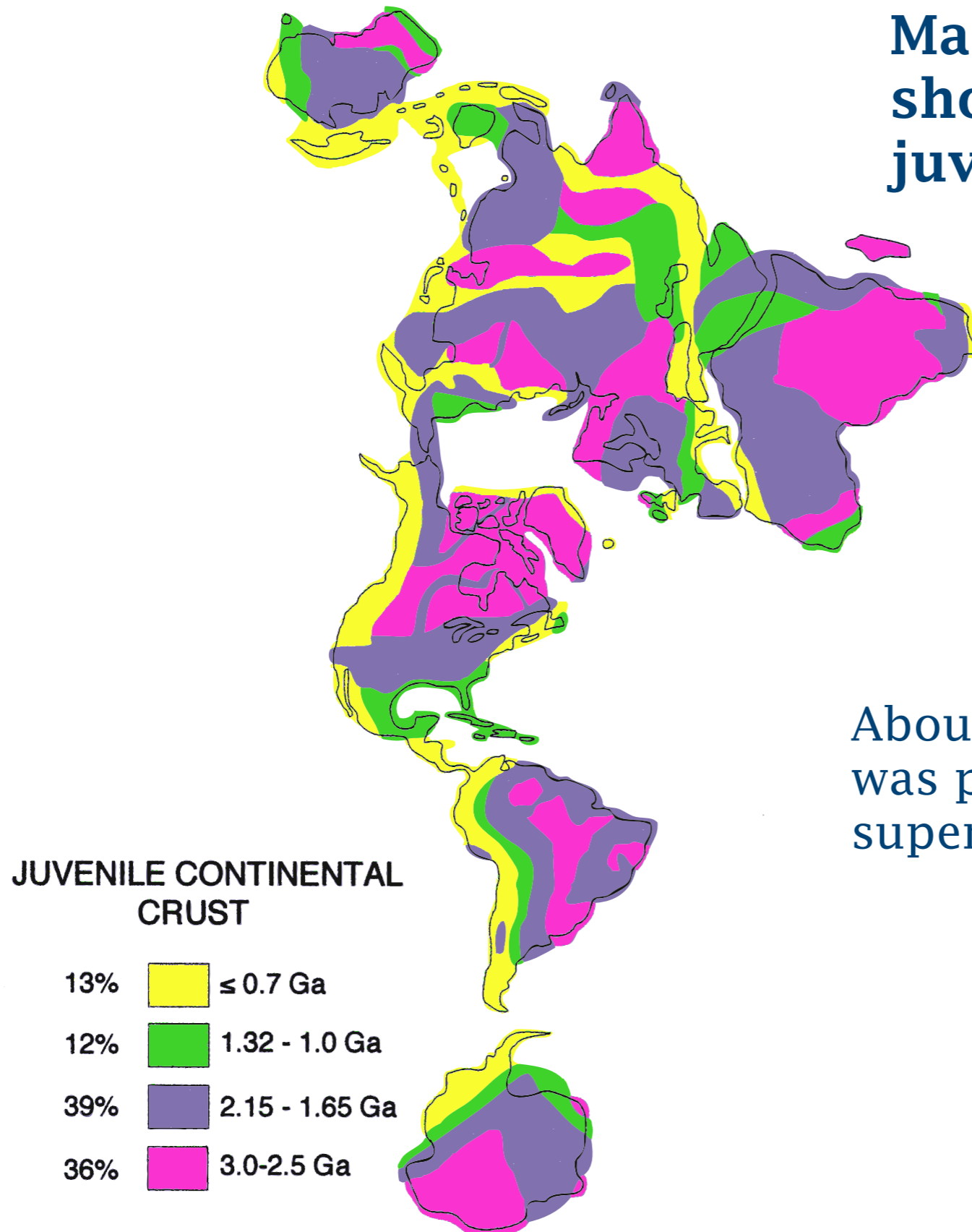
Age of the continental crust



Plot showing the distribution of U-Pb zircon ages in continental crust
(after Condie, 1998)

Spatial distribution of the juvenile crust

Map of the continents showing the distribution of juvenile continental crust



About 75% of the continental crust was produced during the first two super event cycles.

Condie, 1998

Age of the Earth

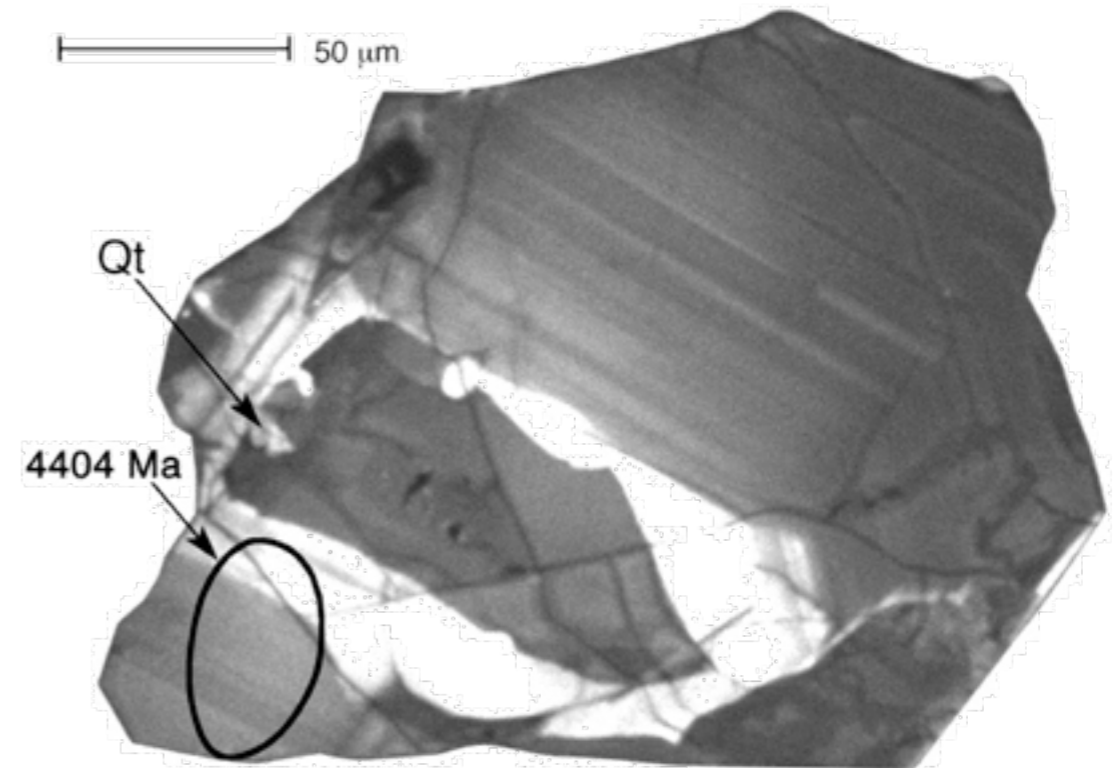
Age of Earth inferred to be 4.55 billion years old (Ga) based on radiometric age determination of meteorites such as Allende.



Age of the continental crust



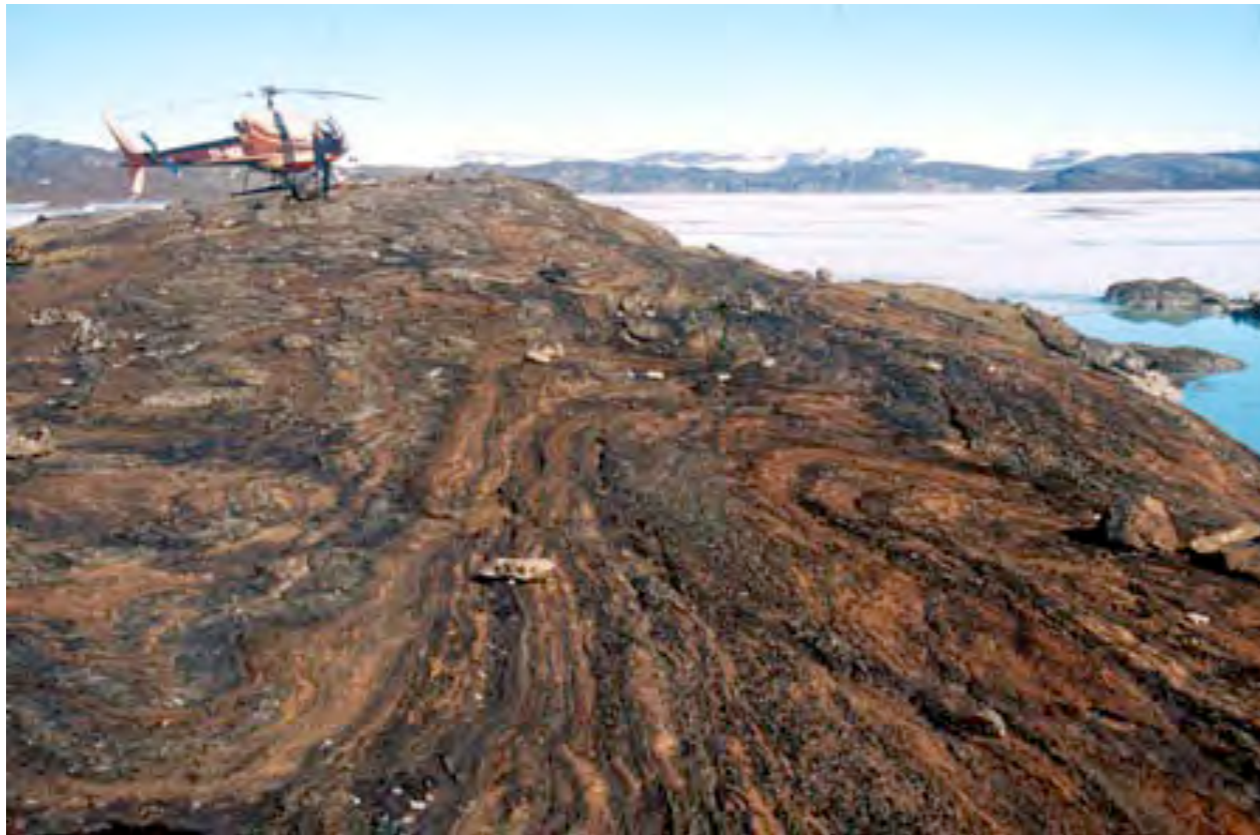
Red quartzite and conglomerate,
Jack Hills, Western Australia
(Peck et al., 2000, 2001)



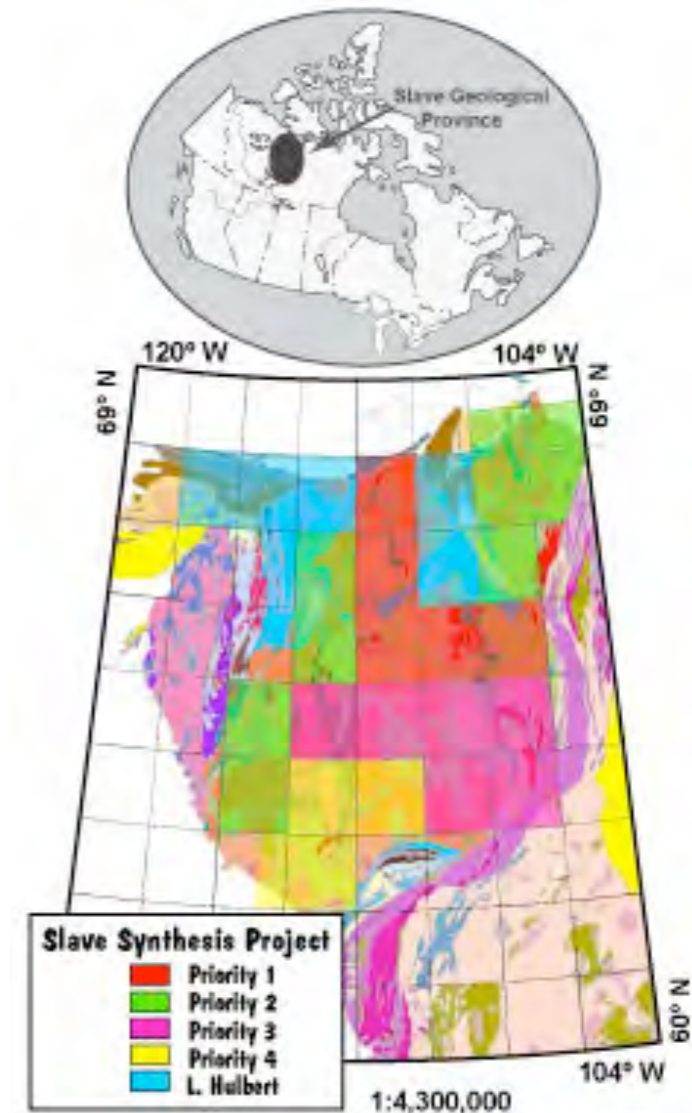
Zircon crystal (Peck et al,
2000)

Oldest dated mineral is zircon crystal 4.4 billion years old
found in Western Australia

Age of the continental crust



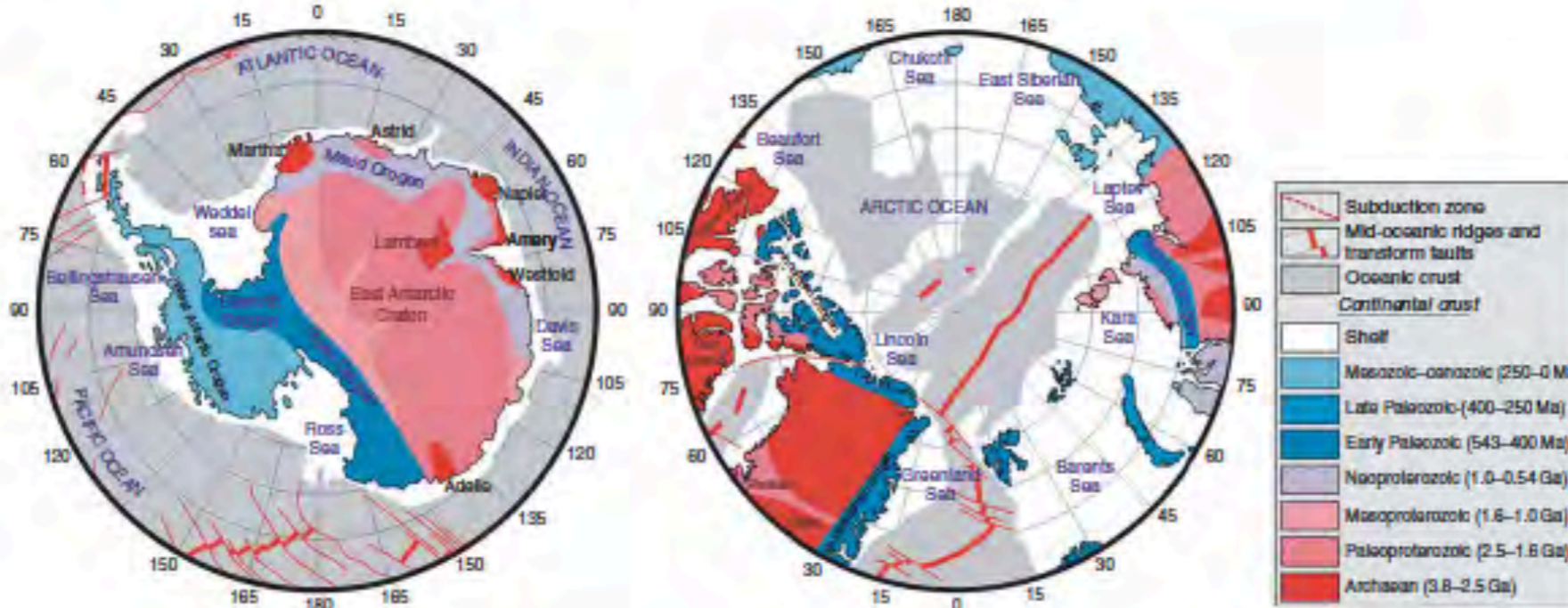
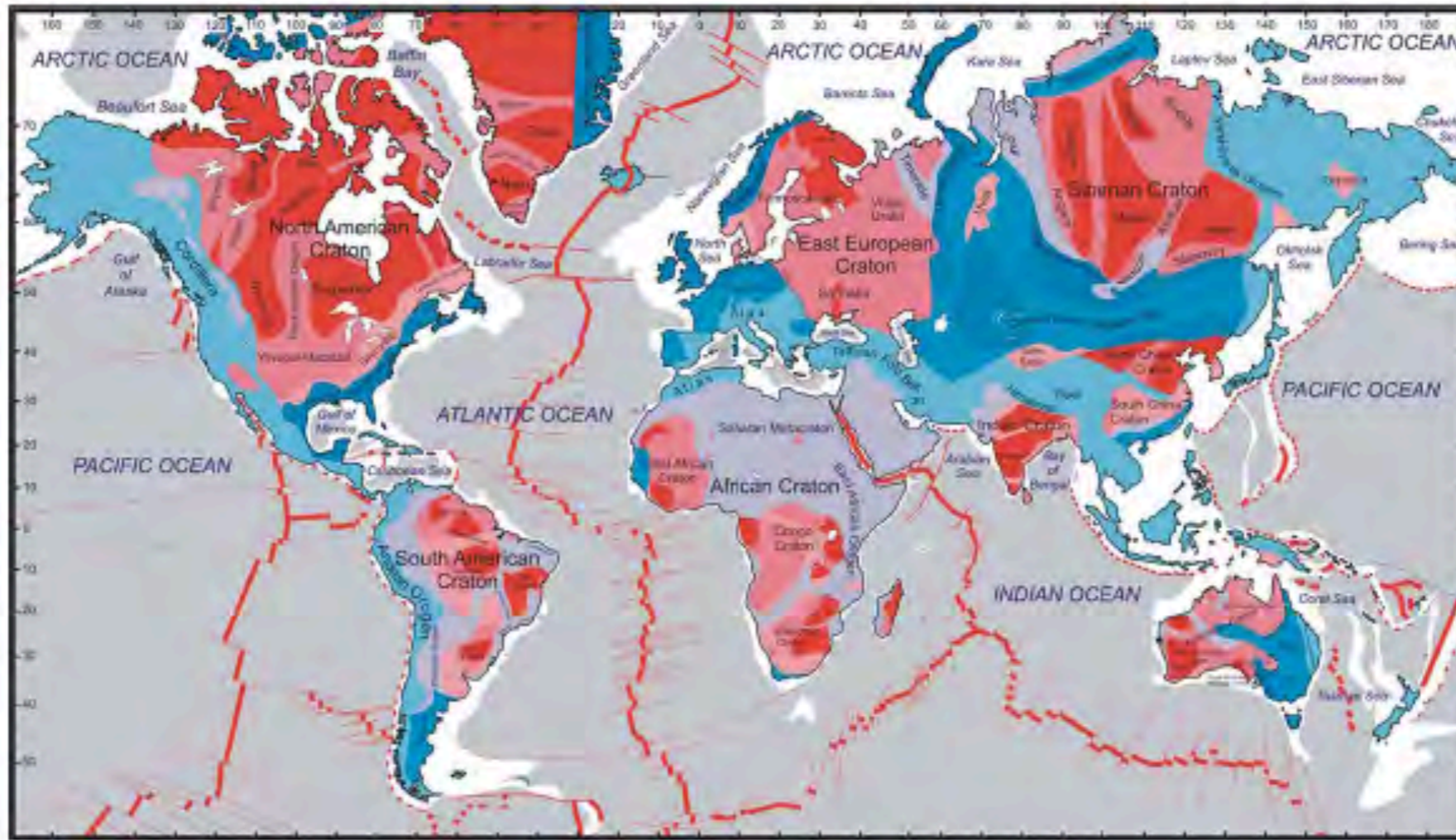
Isua Greenstone Belt, West Greenland
3.8 Ga (David Green, Denison Univ.)



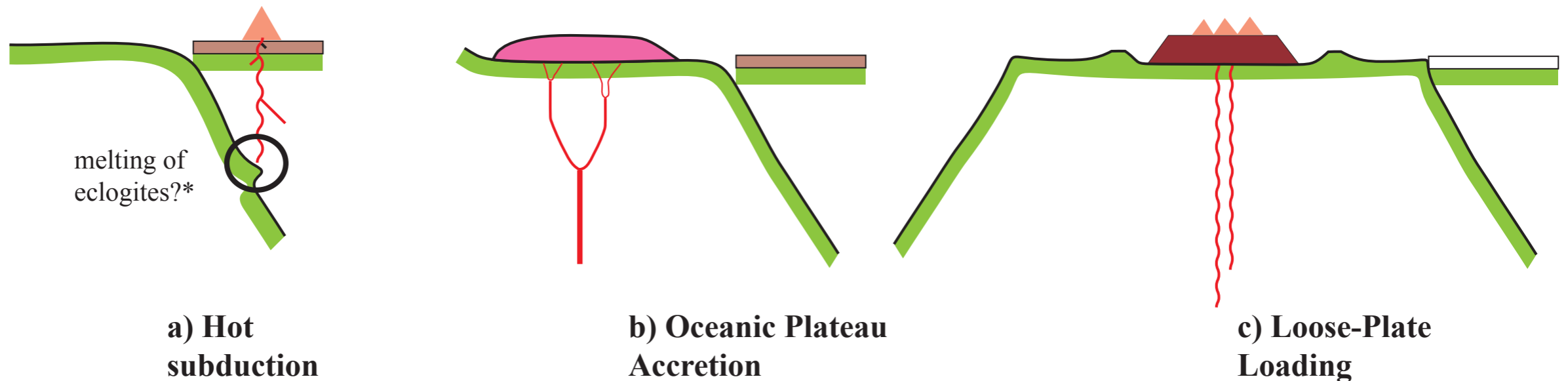
Acasta Gneiss, 4.04 Ga

Oldest dated rock is ~ 4 billion years old

Age of the continental crust



Models of continental protolith



a) hot subduction or subduction of mid-ocean ridges and young oceanic lithosphere makes production of felsic and intermediate liquids by melting of the hydrothermally altered oceanic crust possible. Contribution from the mantle wedge, the subcontinental lithosphere, and the overlying crust may be geochemically identified. The protolith is felsic to intermediate.

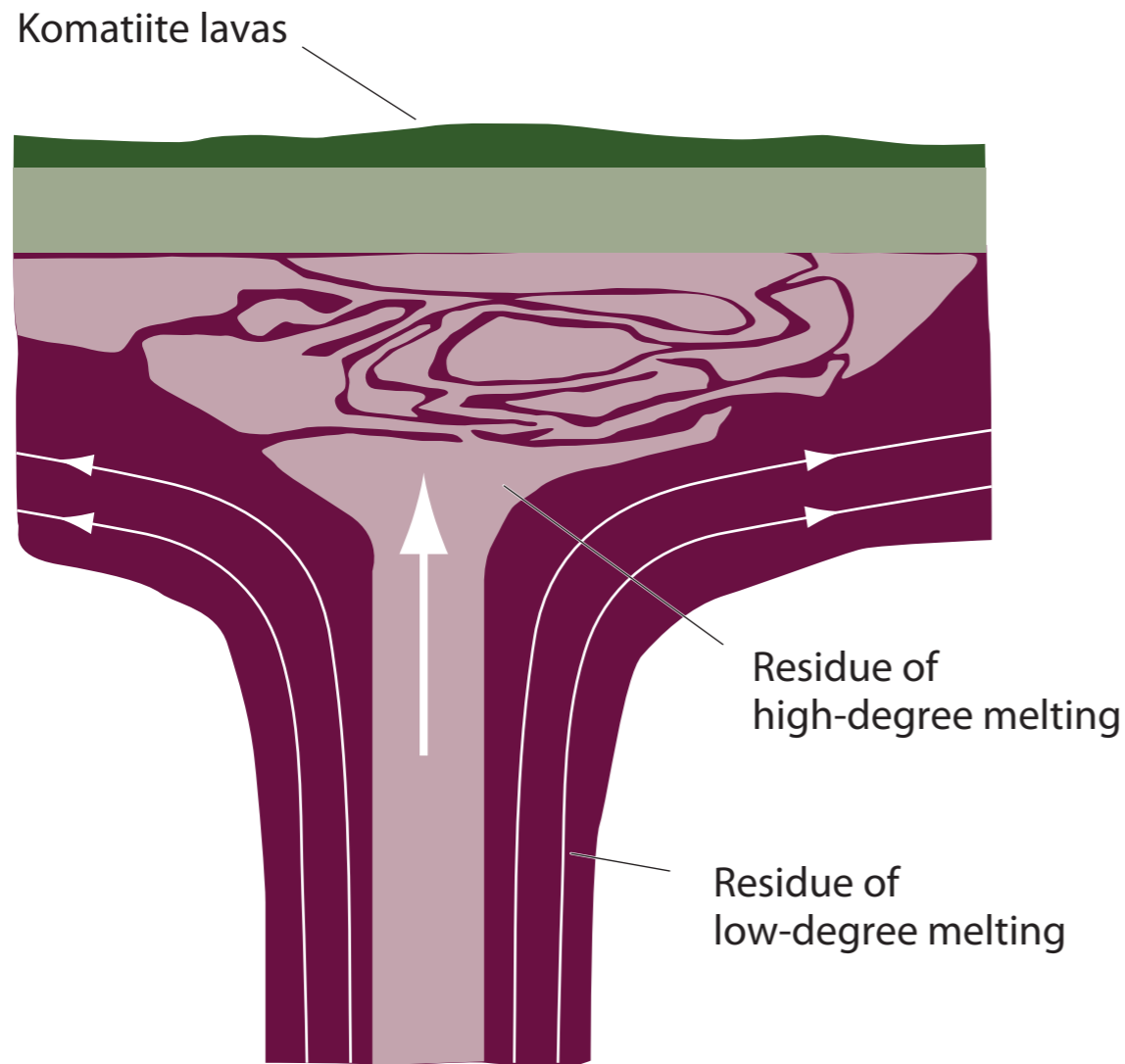
** Archean TTG could be produced by melting of hydrous eclogites (Rapp et al., 2003)*

b) accretion of oceanic plateaus created during superplume events (plume head).

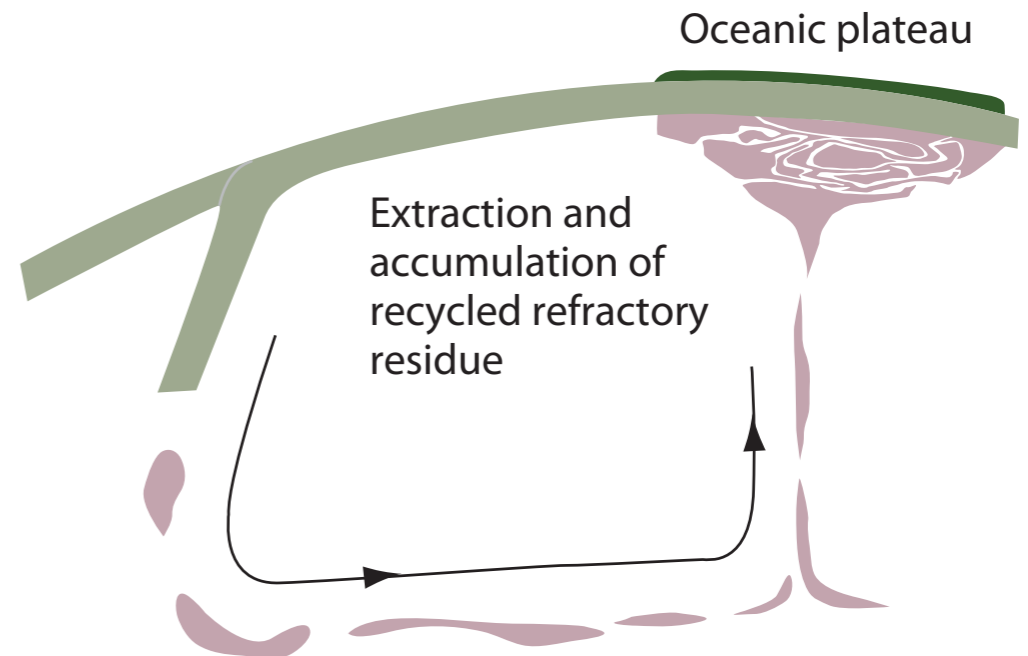
c) protracted loading by plume magma of a loose plate (not entrained by sinking lithosphere). In the last two cases, the protolith is basaltic. Crust is thickened and reaches critical buoyancy.

(after Albarède, 1998)

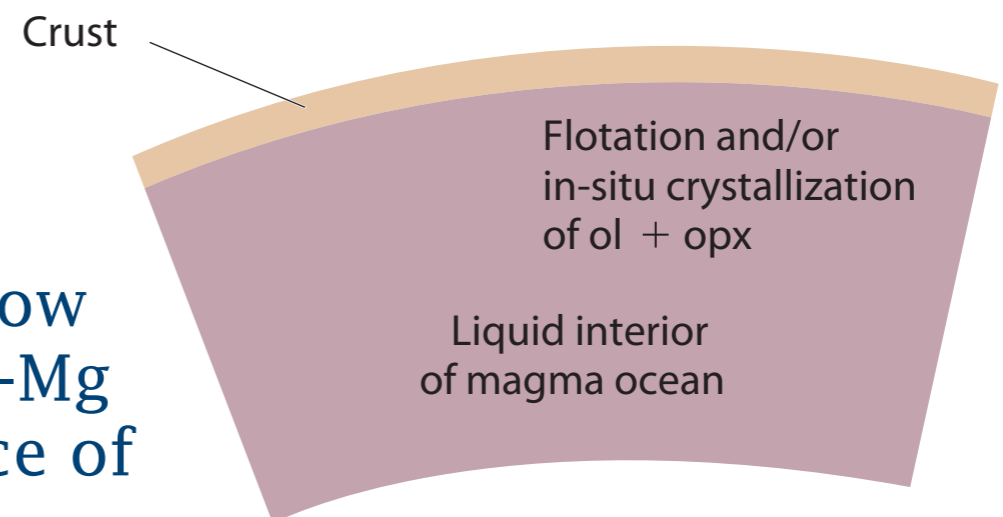
(a) Model 1: Segregation of residue from an upwelling mantle plume



(b) Model 2: Segregation of recycled refractory residue



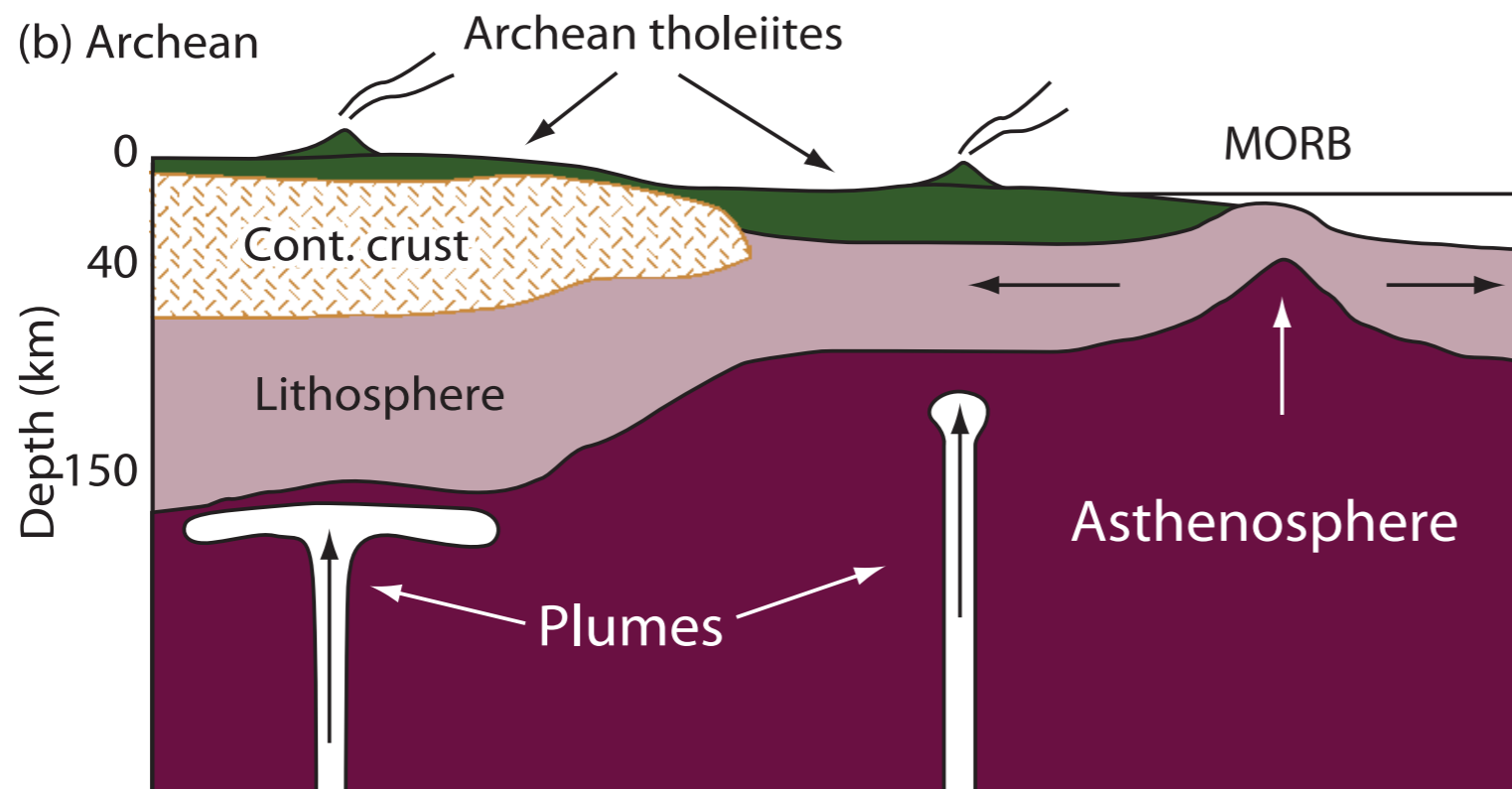
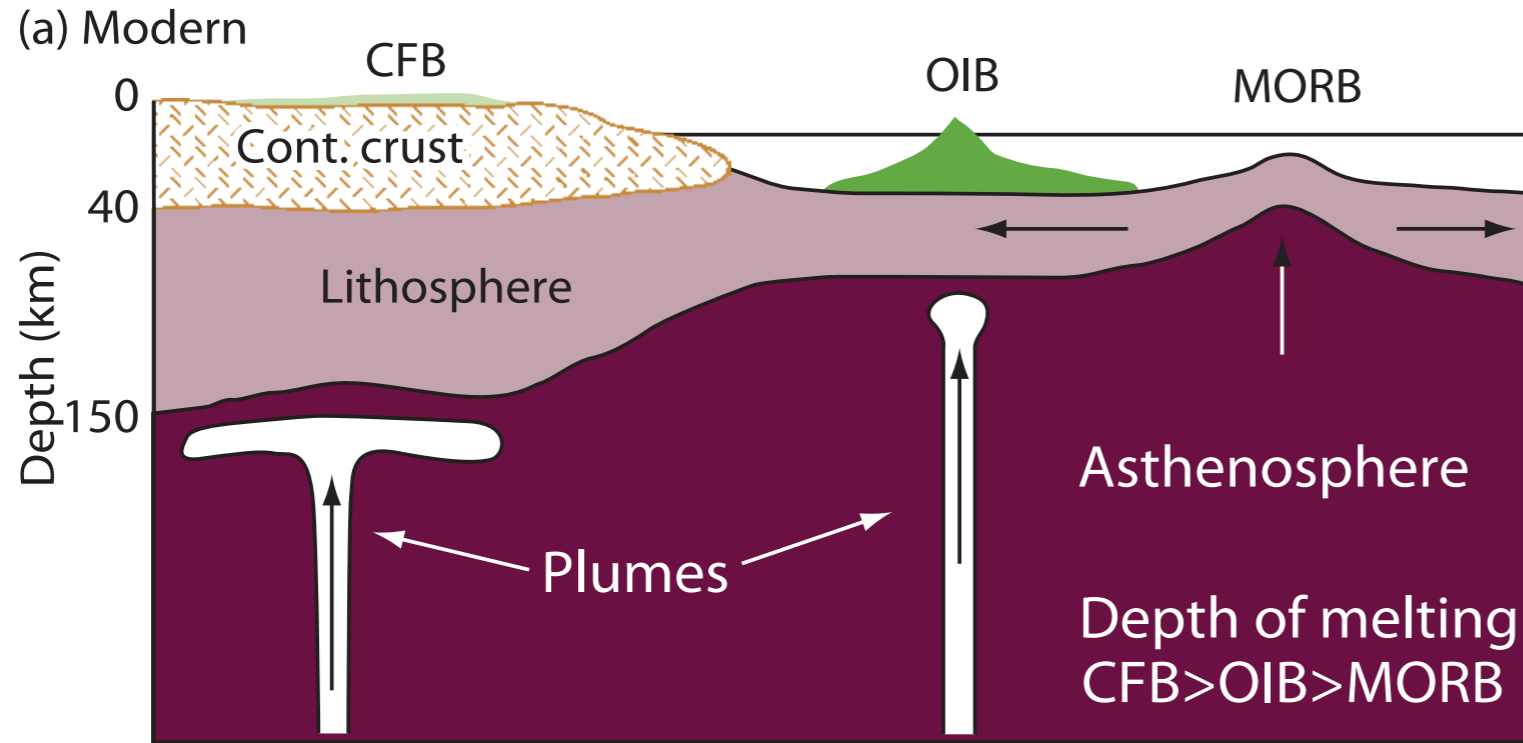
(c) Model 3: Preservation of remnants of the crust of a magma ocean



Three possible mechanisms that could allow the segregation and accumulation of high-Mg olivine and orthopyroxene near the surface of the Earth in the early Earth

after Arndt et al., 2002

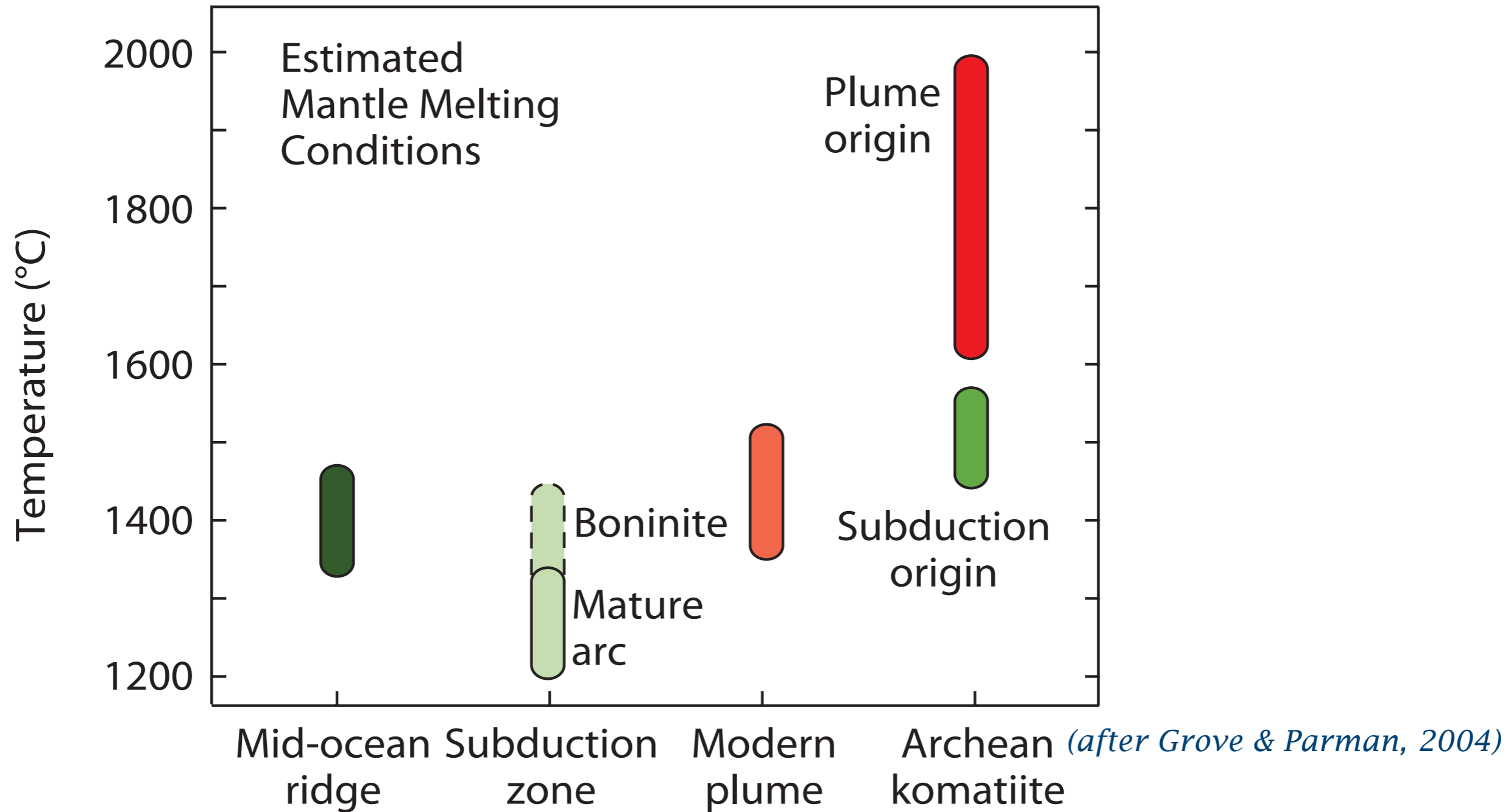
Model of komatiitic and tholeiitic basalt formation involving mantle plumes



Model shows the influence of lithospheric thickness on depth of melting where CFB is continental flood basalt, OIB oceanic island basalt, and MORB mid-ocean ridge basalt

(after Arndt et al., 1997)

Melting temperature vs age



The range of mantle melt generation temperatures estimated for various modern tectonic settings compared to temperatures inferred for komatiite melt generation by a plume model and a subduction model