

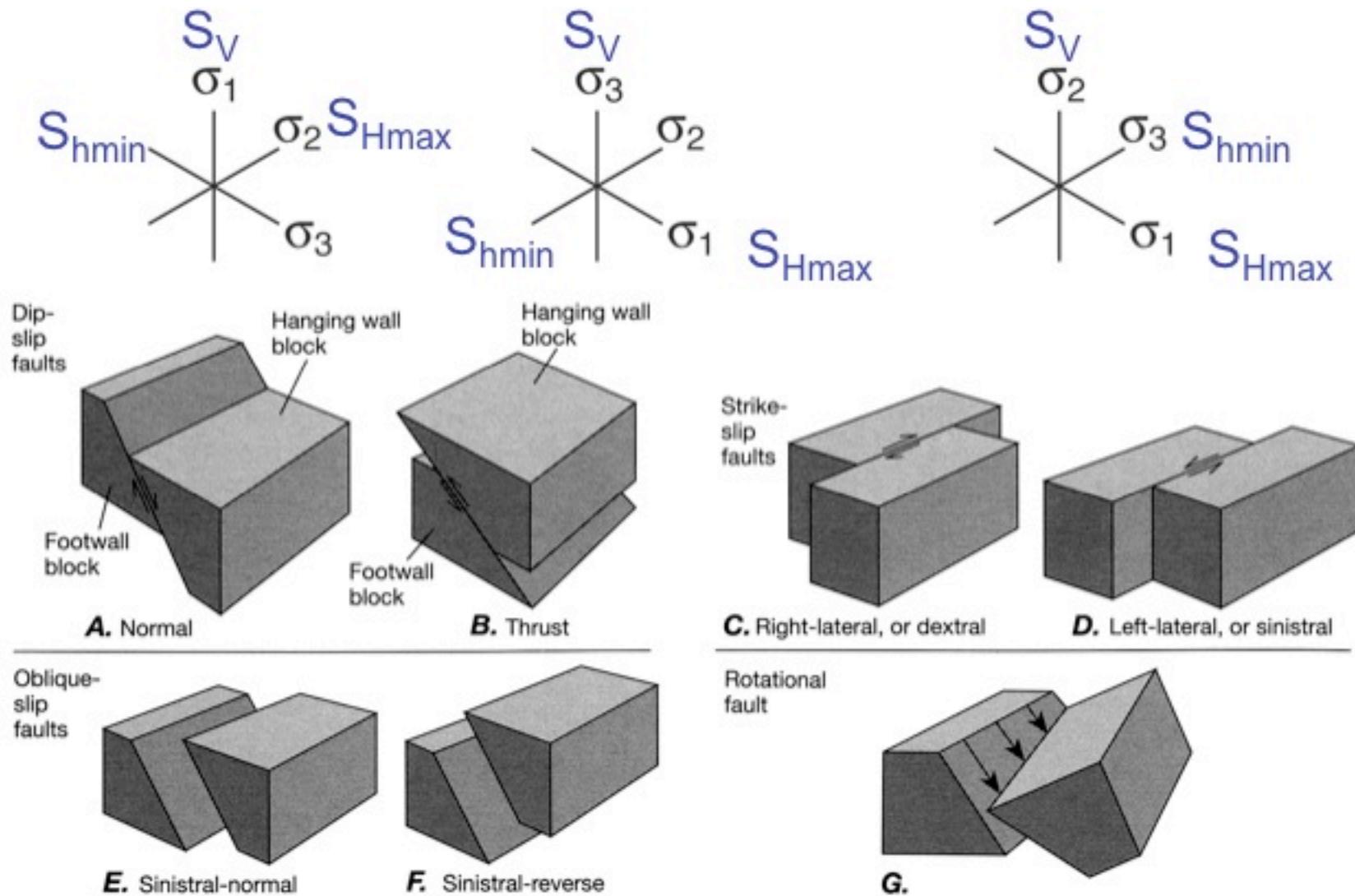


Tektonische Kräfte & Plattenbewegungen

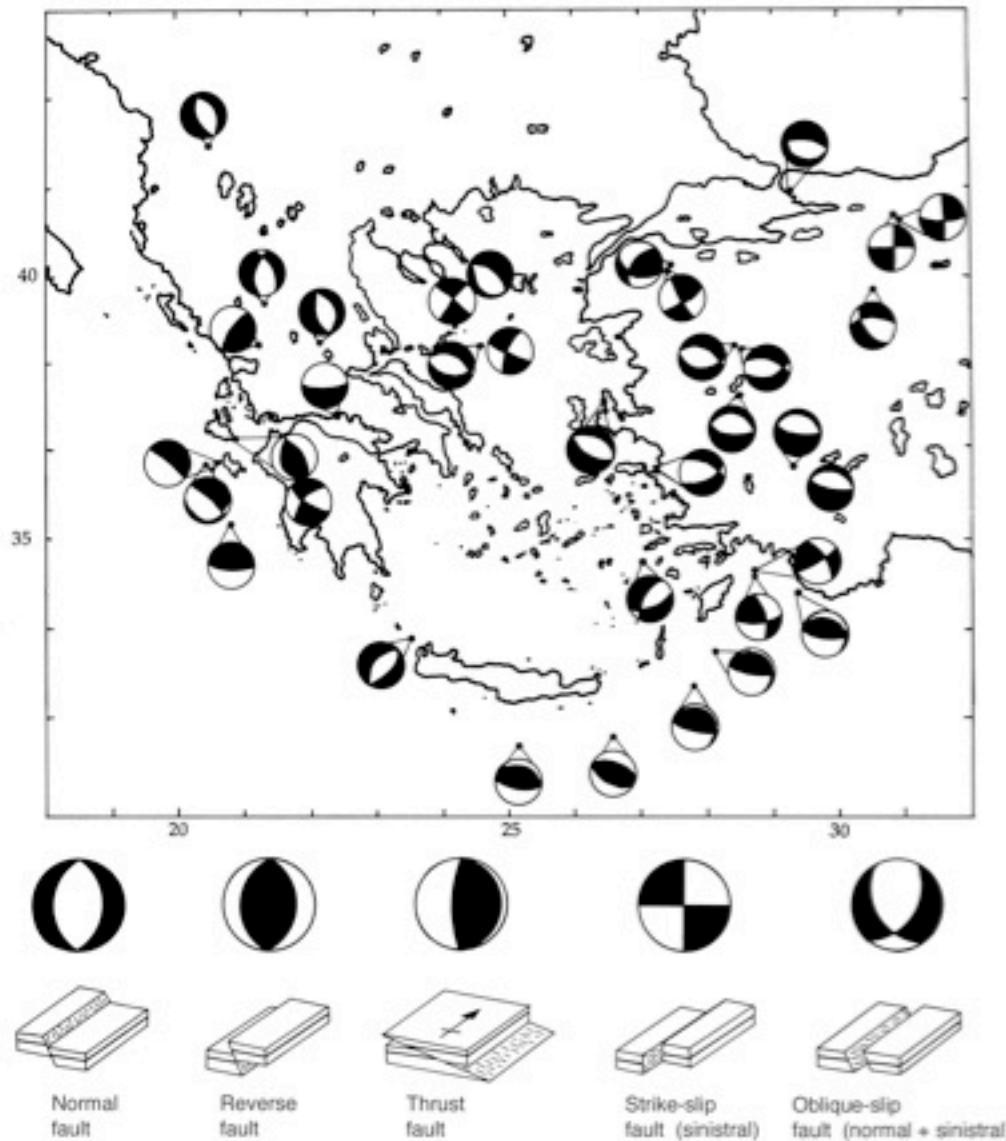
Topics: Jan 15, 2009

- (1) Indicators of tectonic stress fields
- (2) Tectonic stress-field provinces
- (3) Neotectonics in intraplate regions

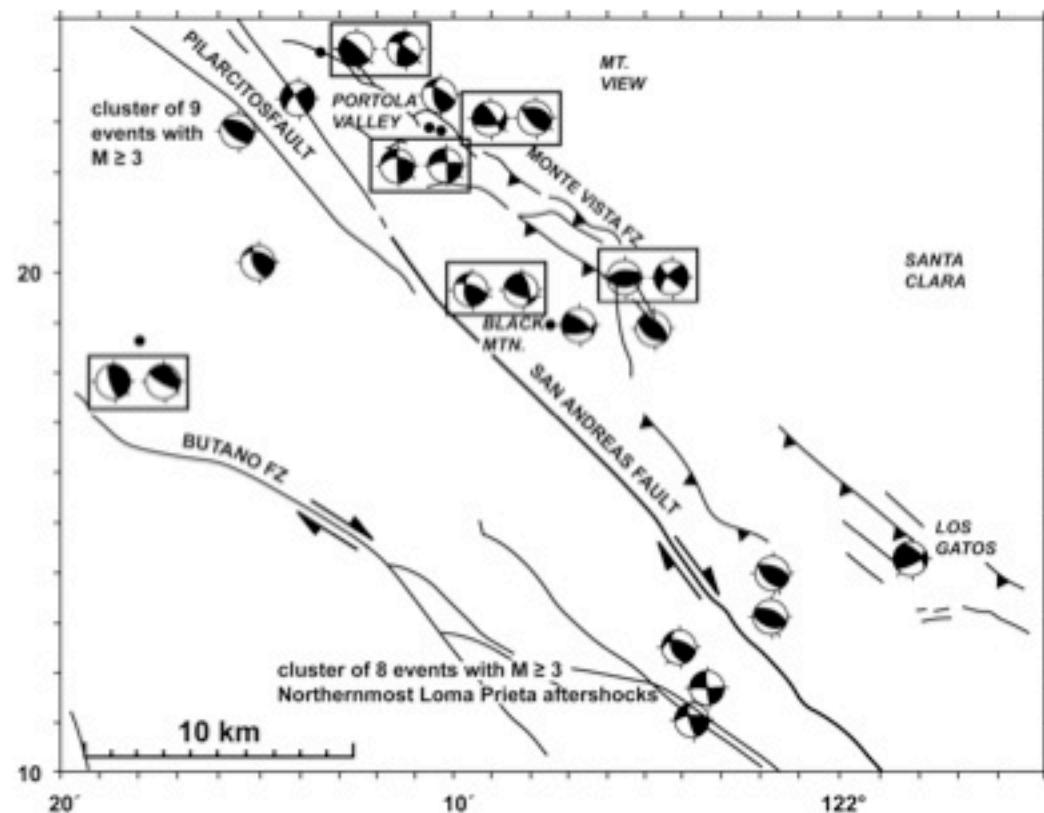
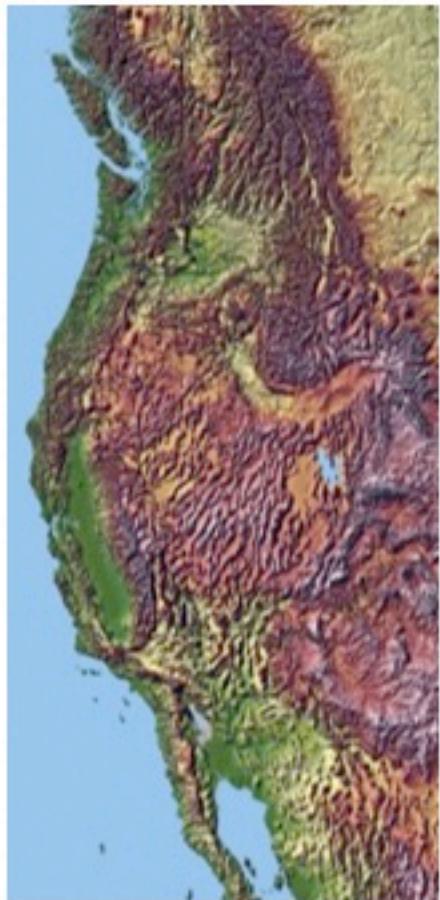
(1) Indicators of tectonic stress fields

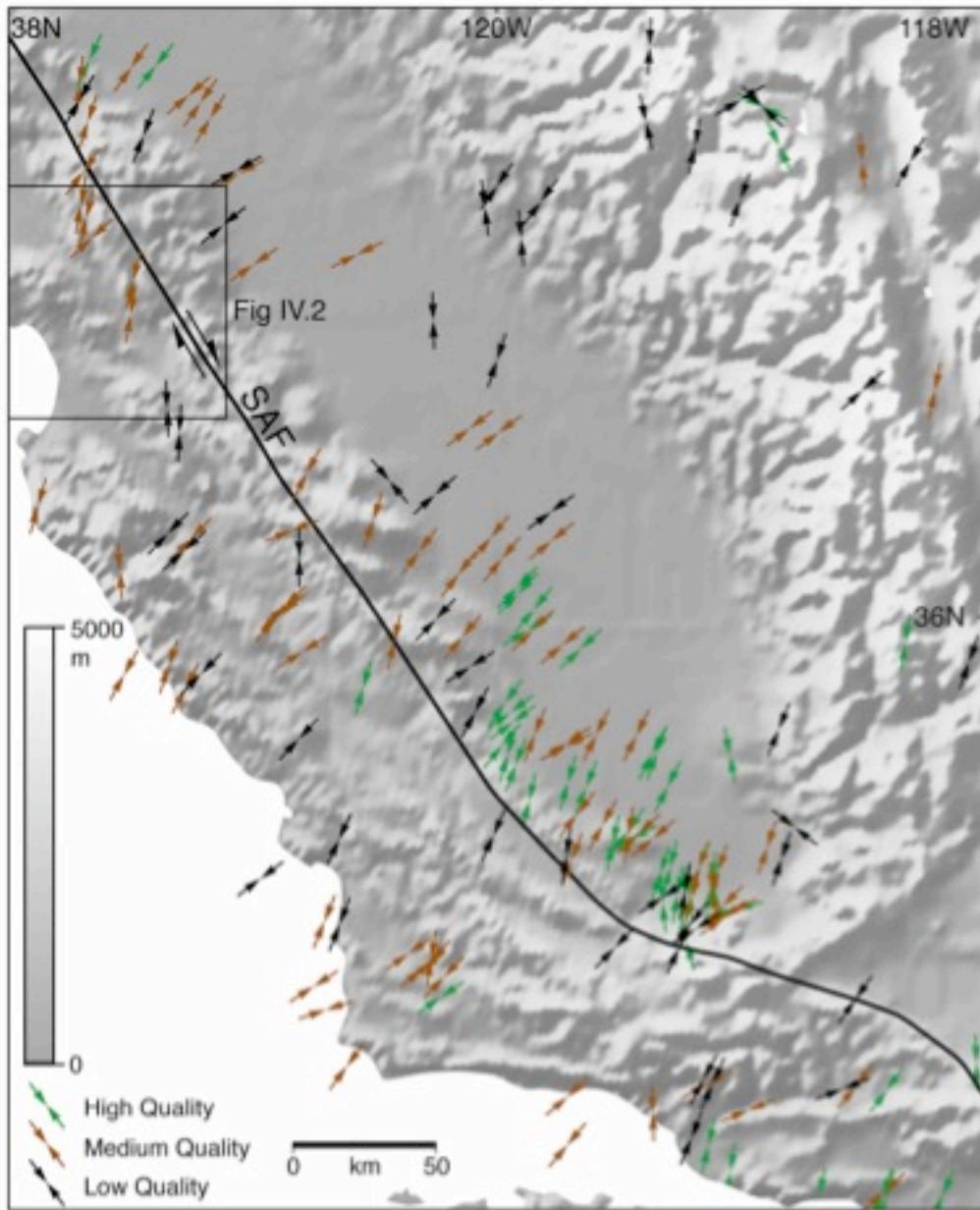


Earthquake focal mechanisms



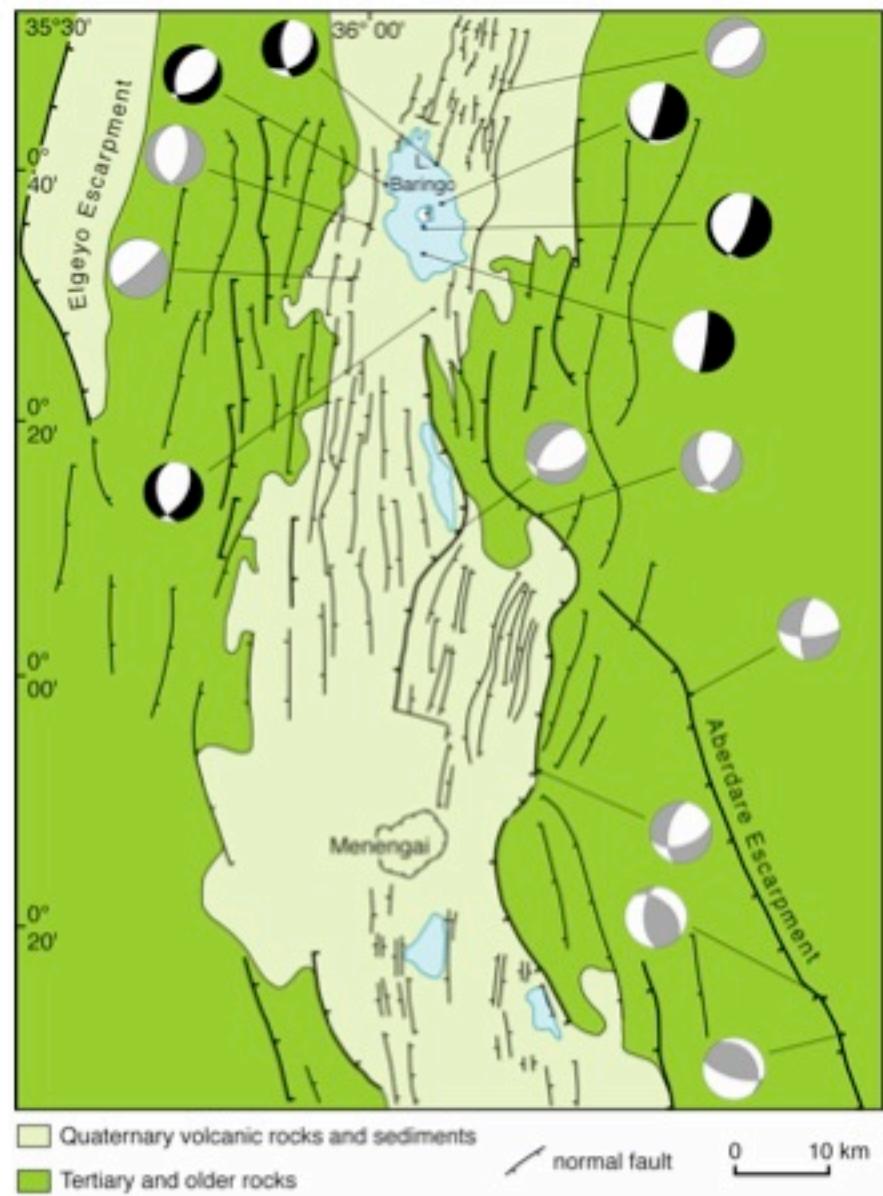
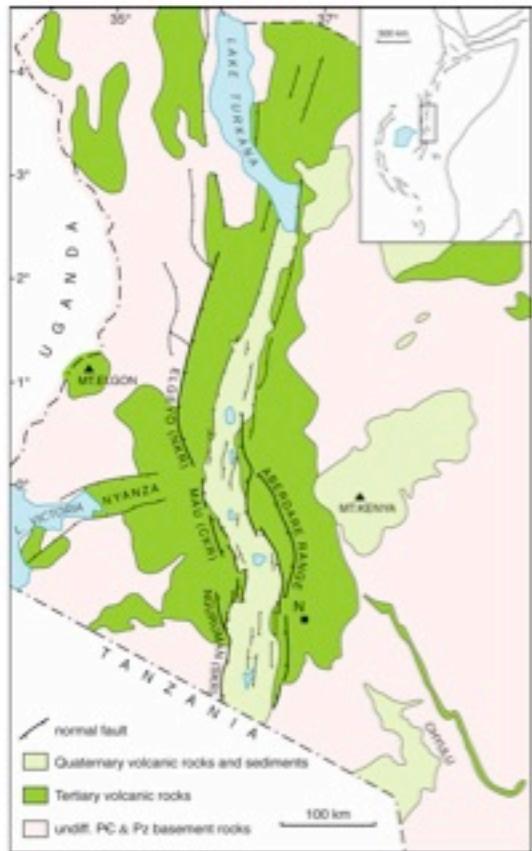
State of stress along the San Andreas Fault



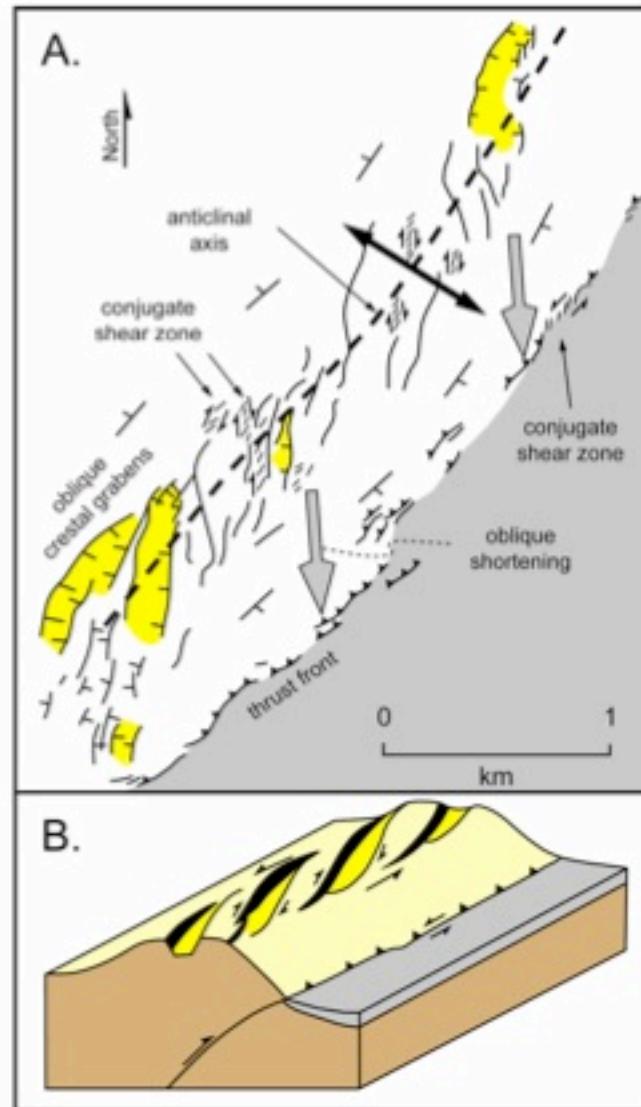
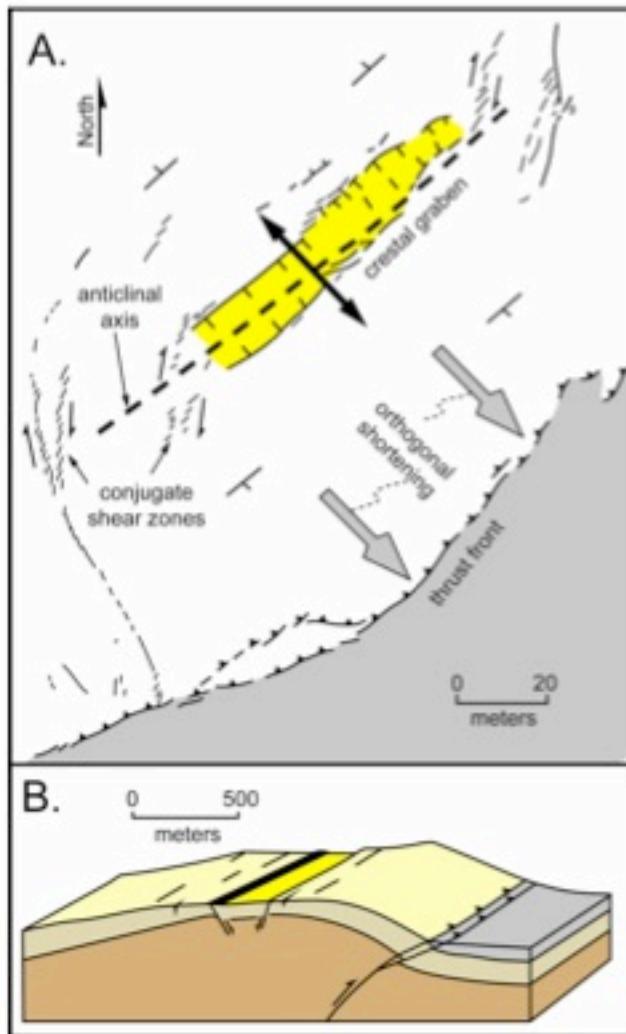


Different types of stress-field indicators in the immediate vicinity of the San Andreas Fault show a fault-normal orientation of $S_{H\max}$

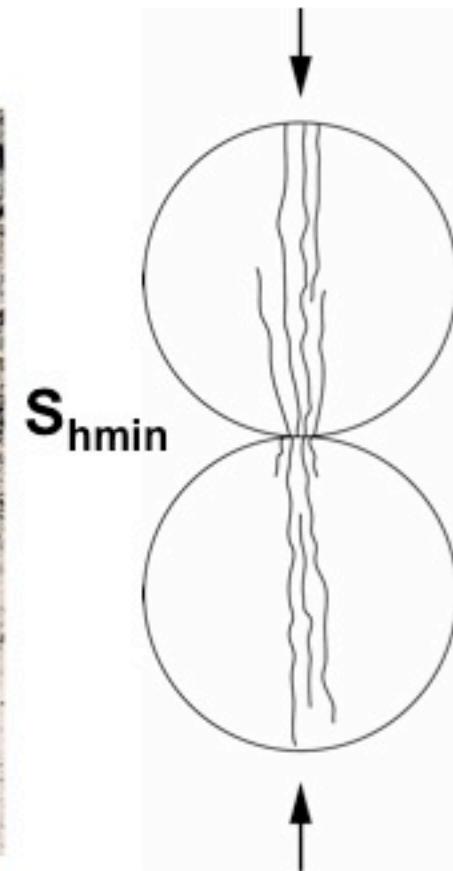
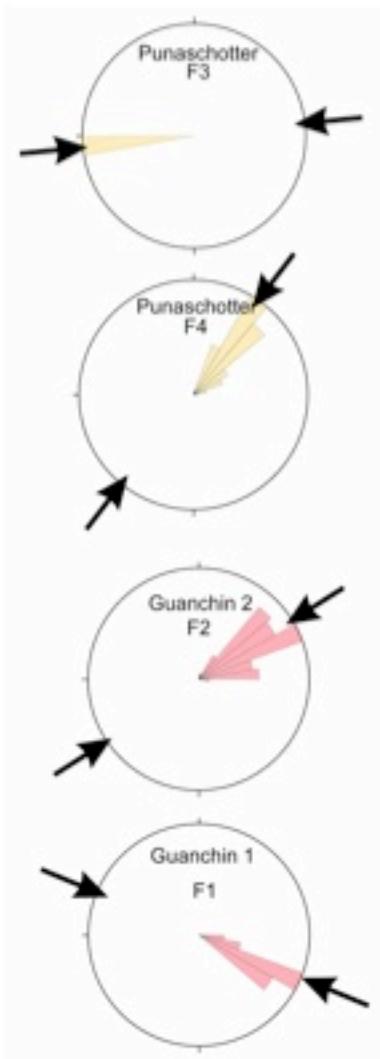
Diverging strikes of different normal fault generations and inferred S_{hmin} orientation obtained from focal mechanisms



Inferences based on structures: surface ruptures of the 1980 El Asnam earthquake



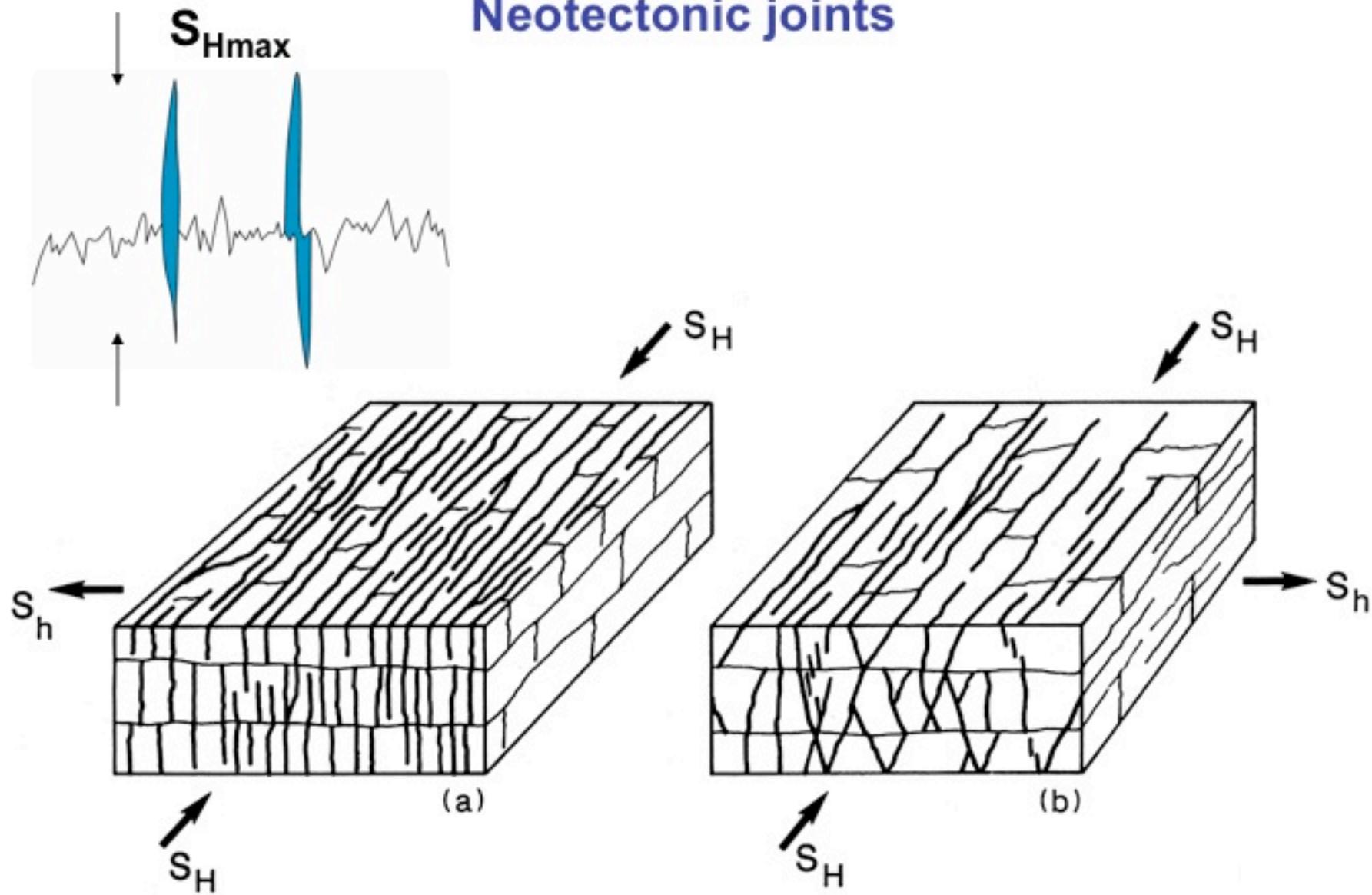
Transgranular rock fractures - inferring $S_{H\max}$



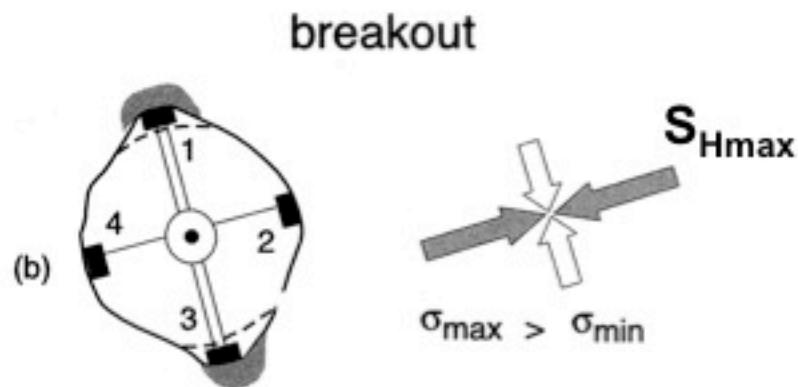
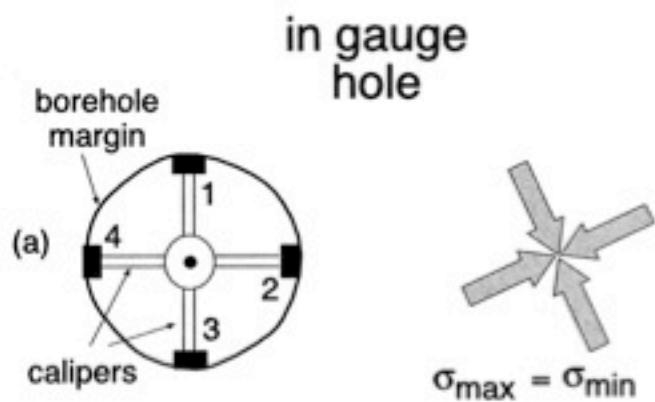
$S_{H\max}$

Stylolites

Neotectonic joints



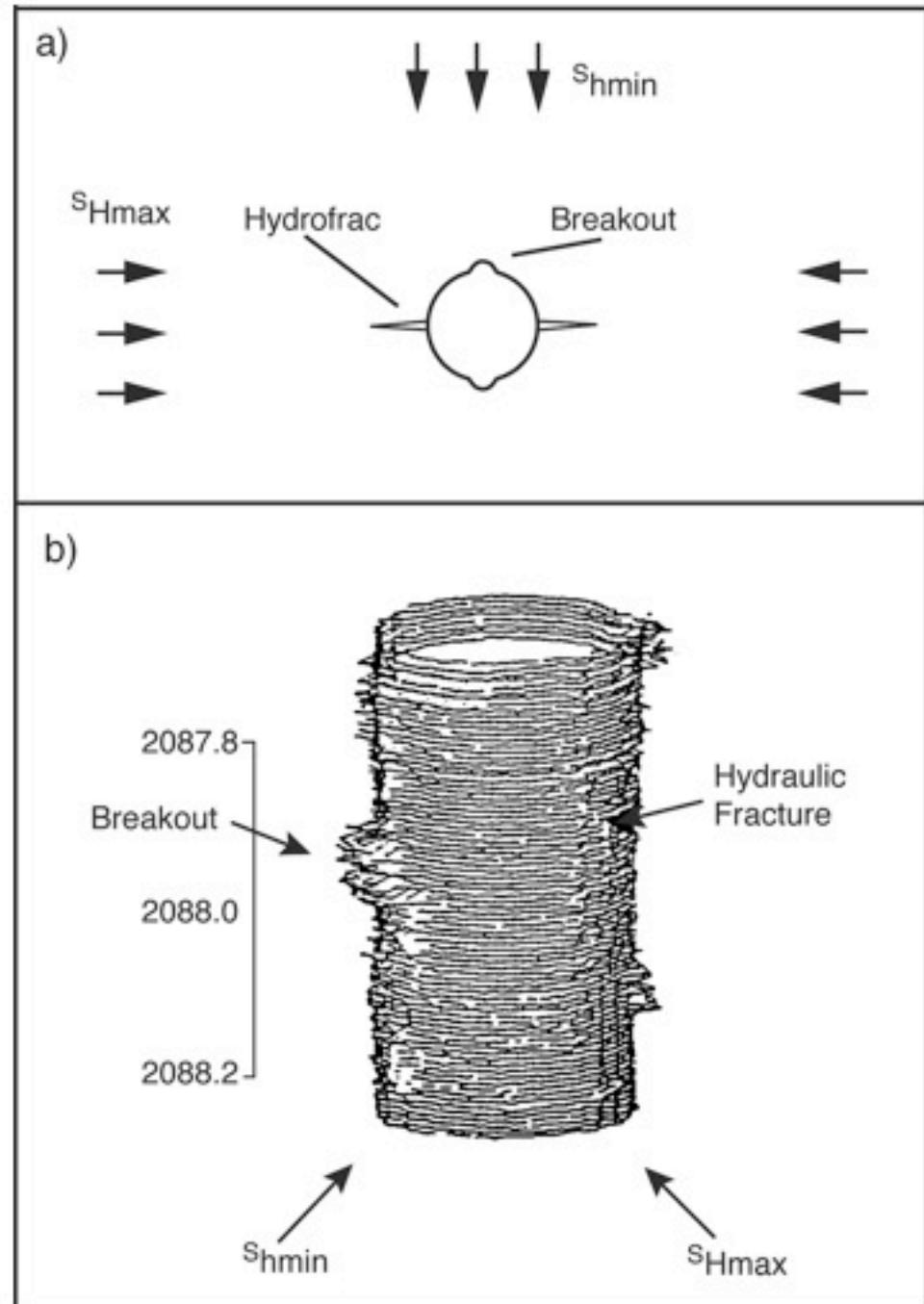
Borehole breakouts



Borehole breakouts and hydrofracs

Hydrofracs = $S_{h\max}$
Fracture perpendicular to $S_{h\min}$

Breakouts = $S_{h\min}$
Shear fractures oblique to $S_{h\max}$

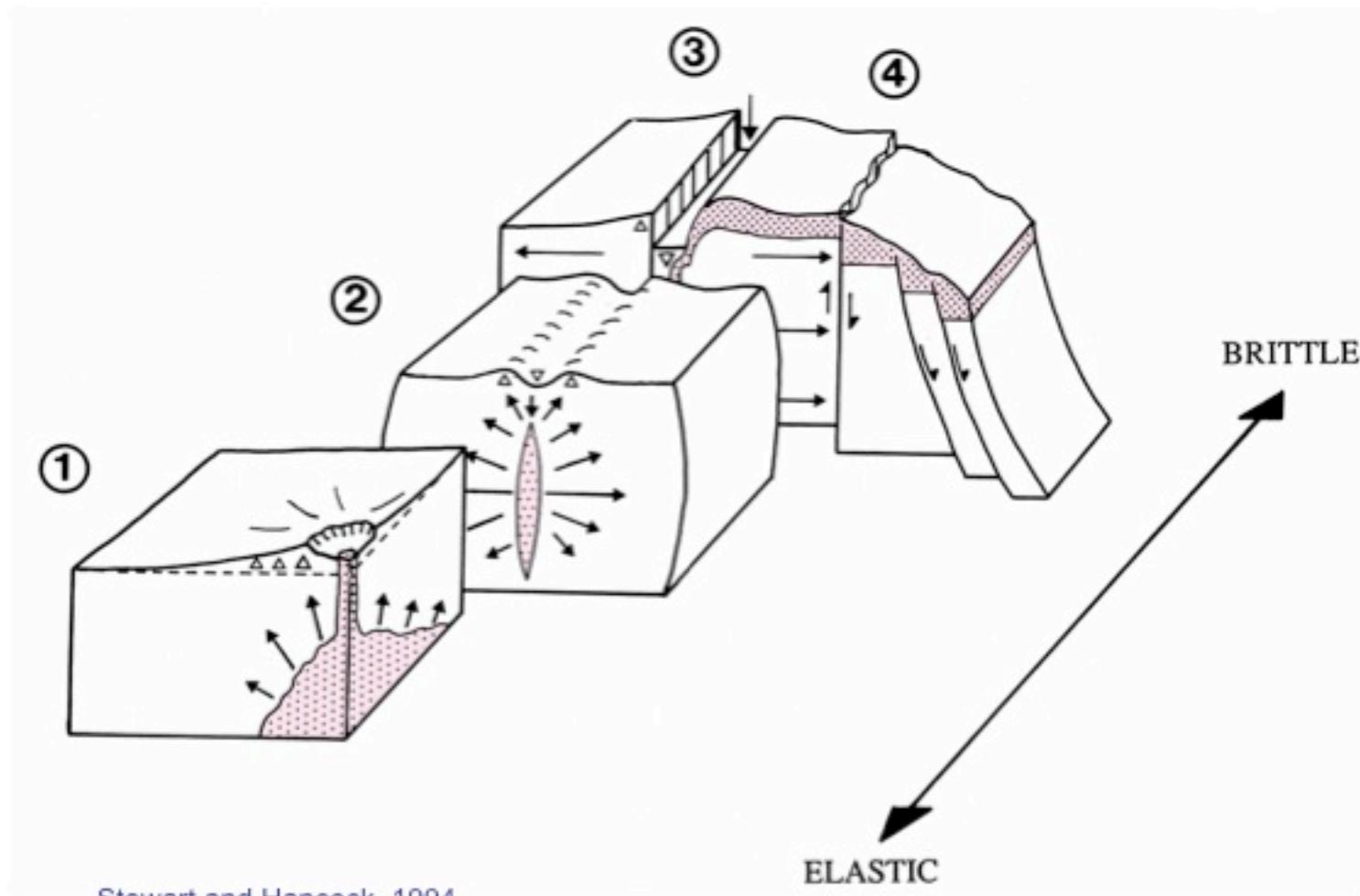


Neptunic dykes

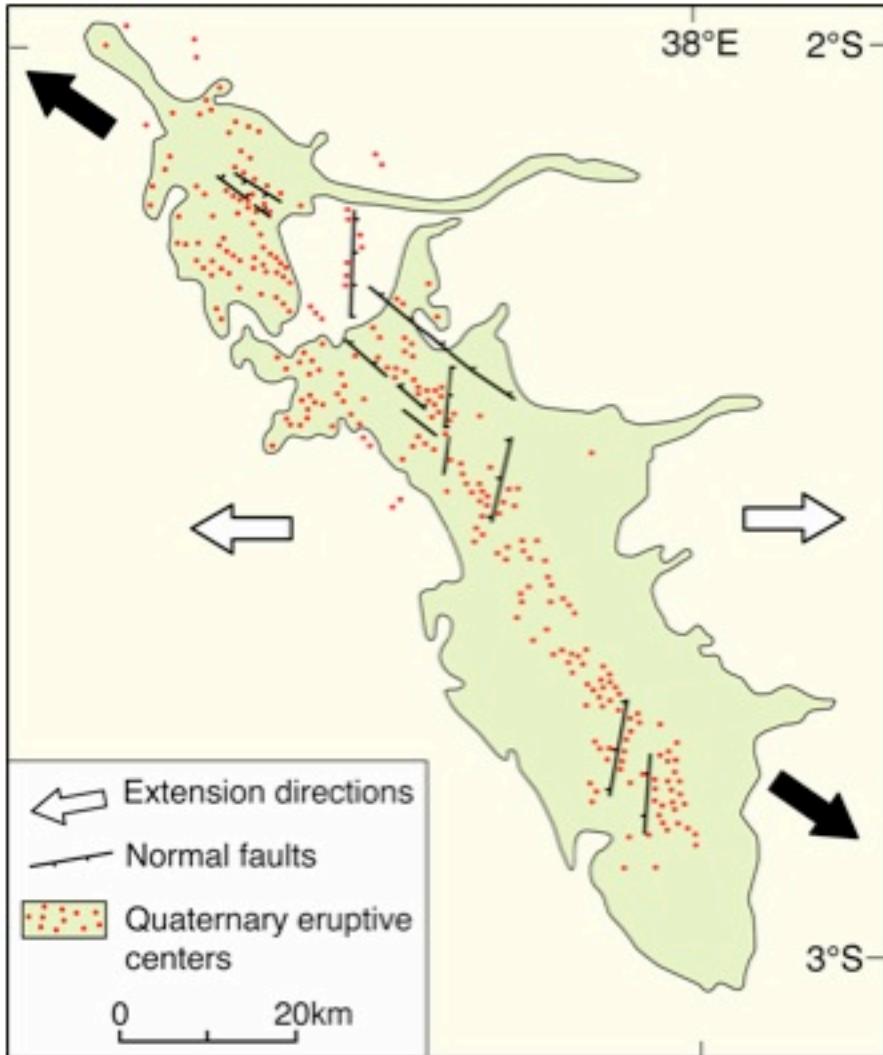
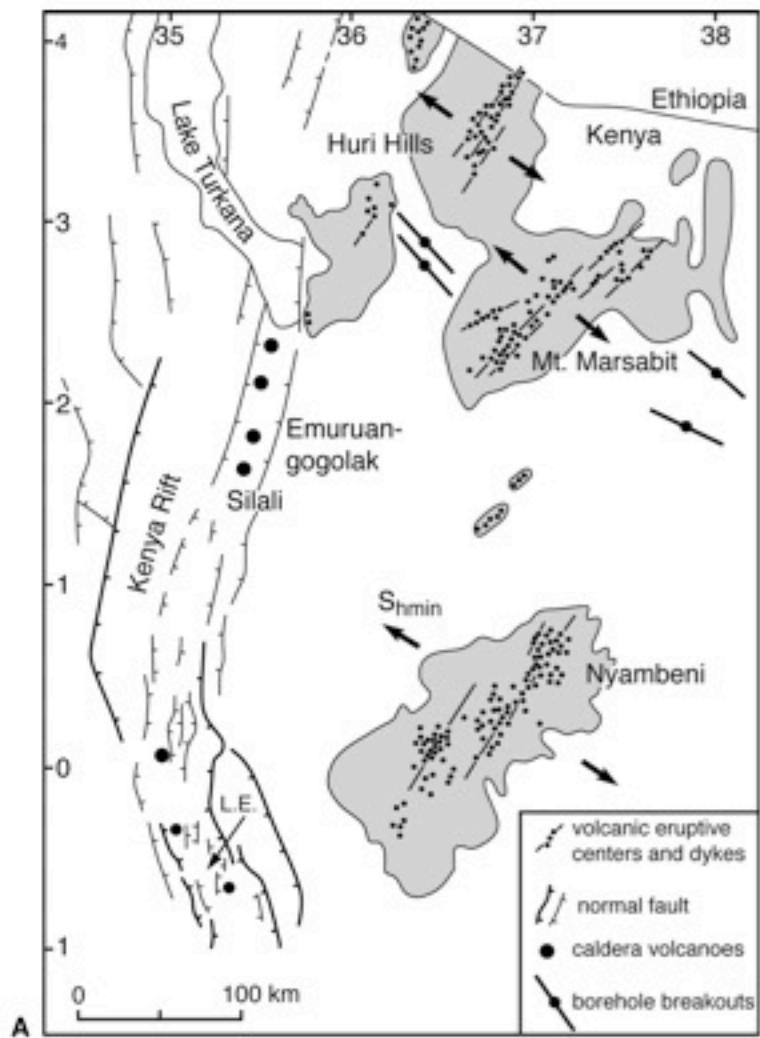
$$P_f = r_f \cdot g \cdot D > S_v = r_{\text{sed}} \cdot g \cdot D$$



Dyke swarms, volcanic alignments & caldera elongation

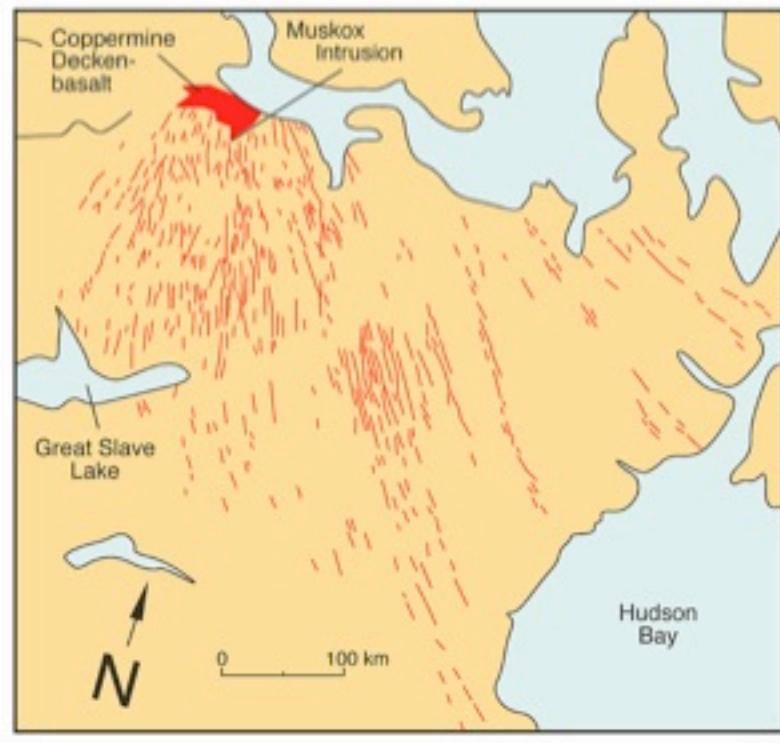


Volcanic chains

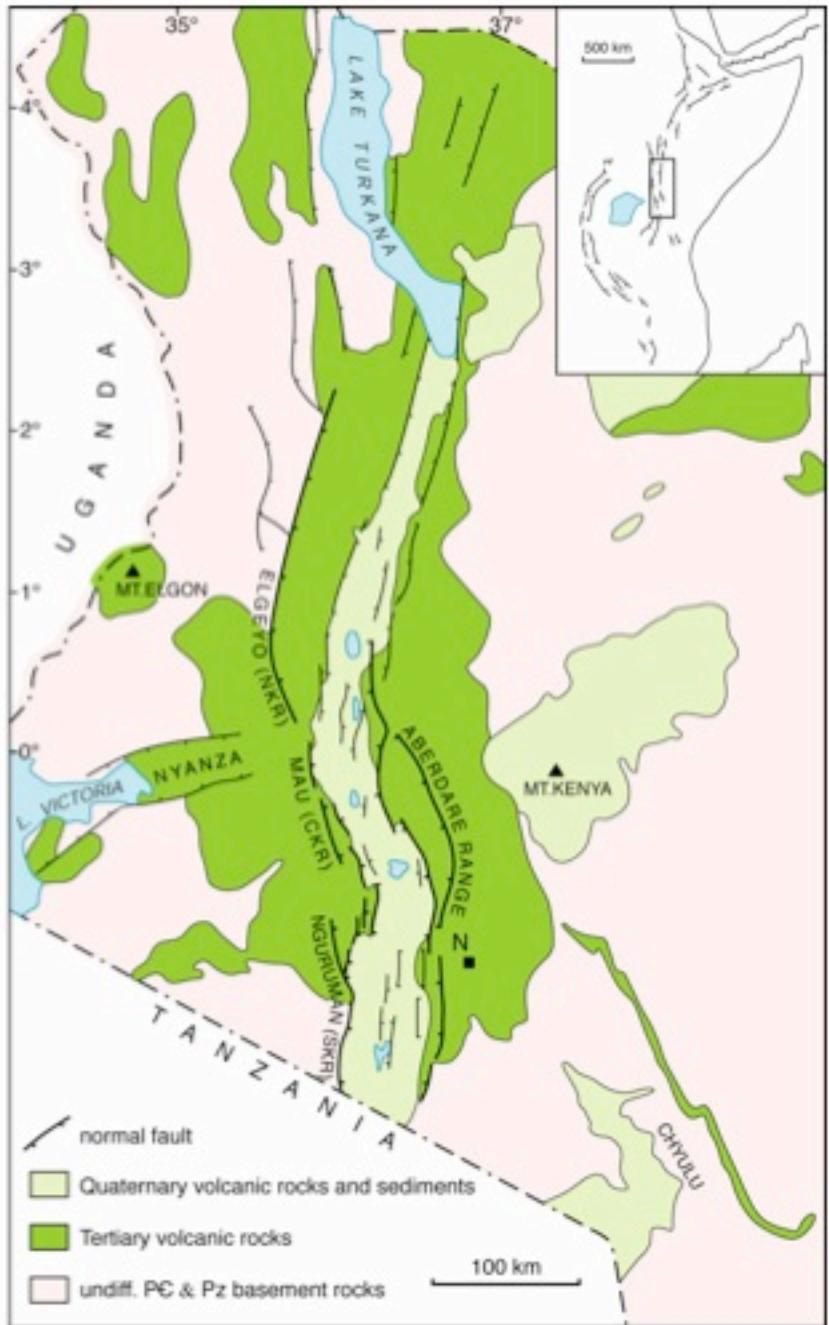


Haug and Strecker, 1994; Bosworth and Strecker, 1998;

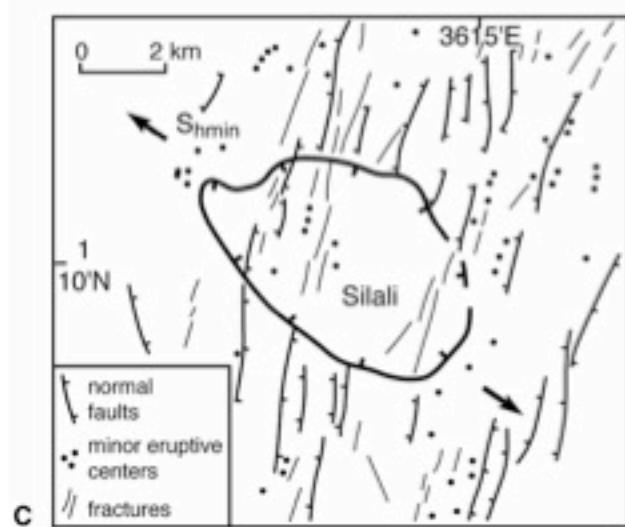
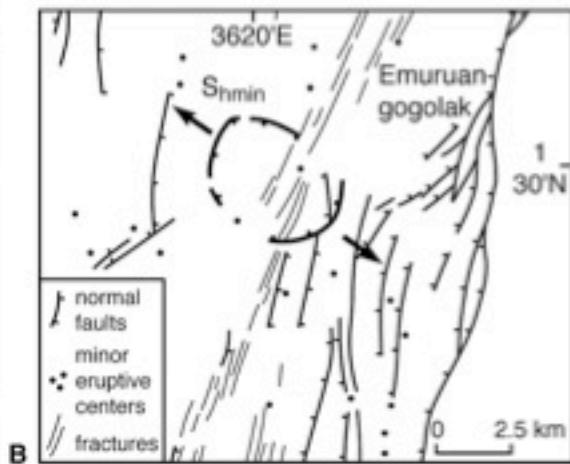
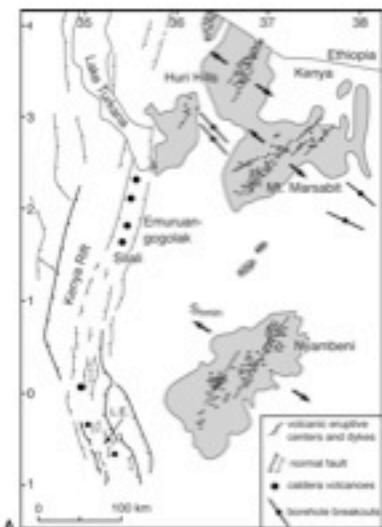
Dyke swarms

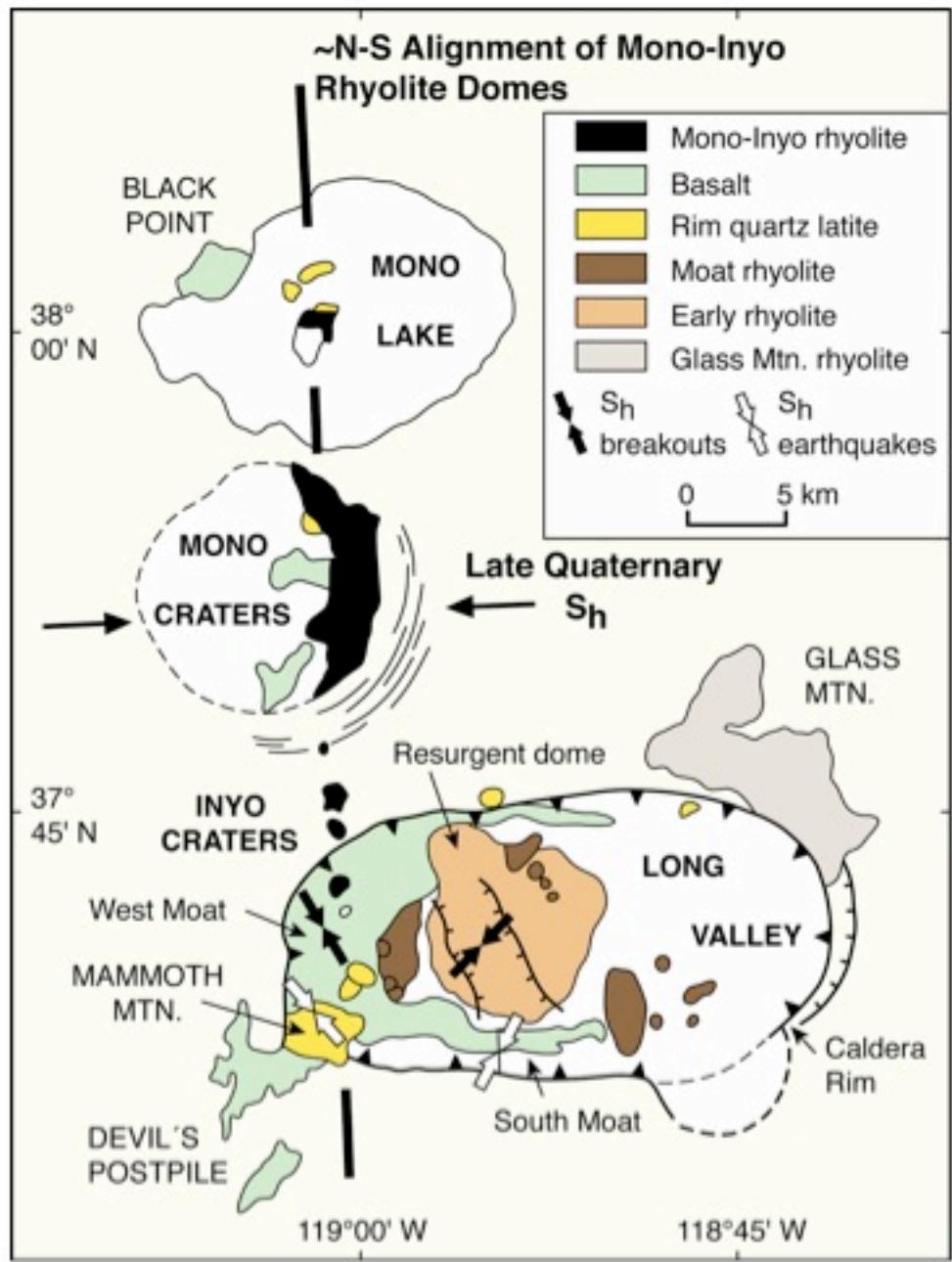
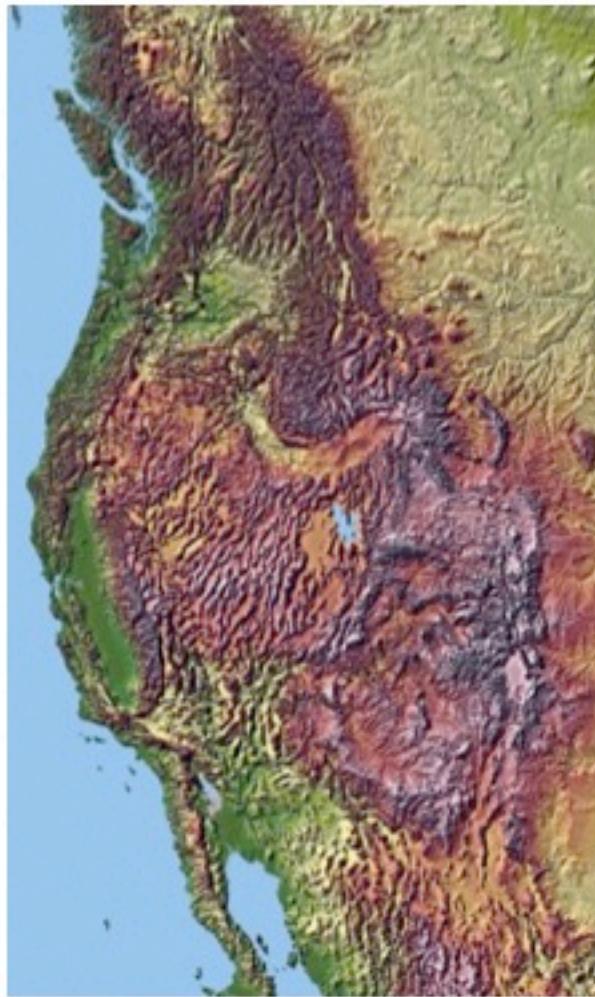


Halls and Fahrig, 1987



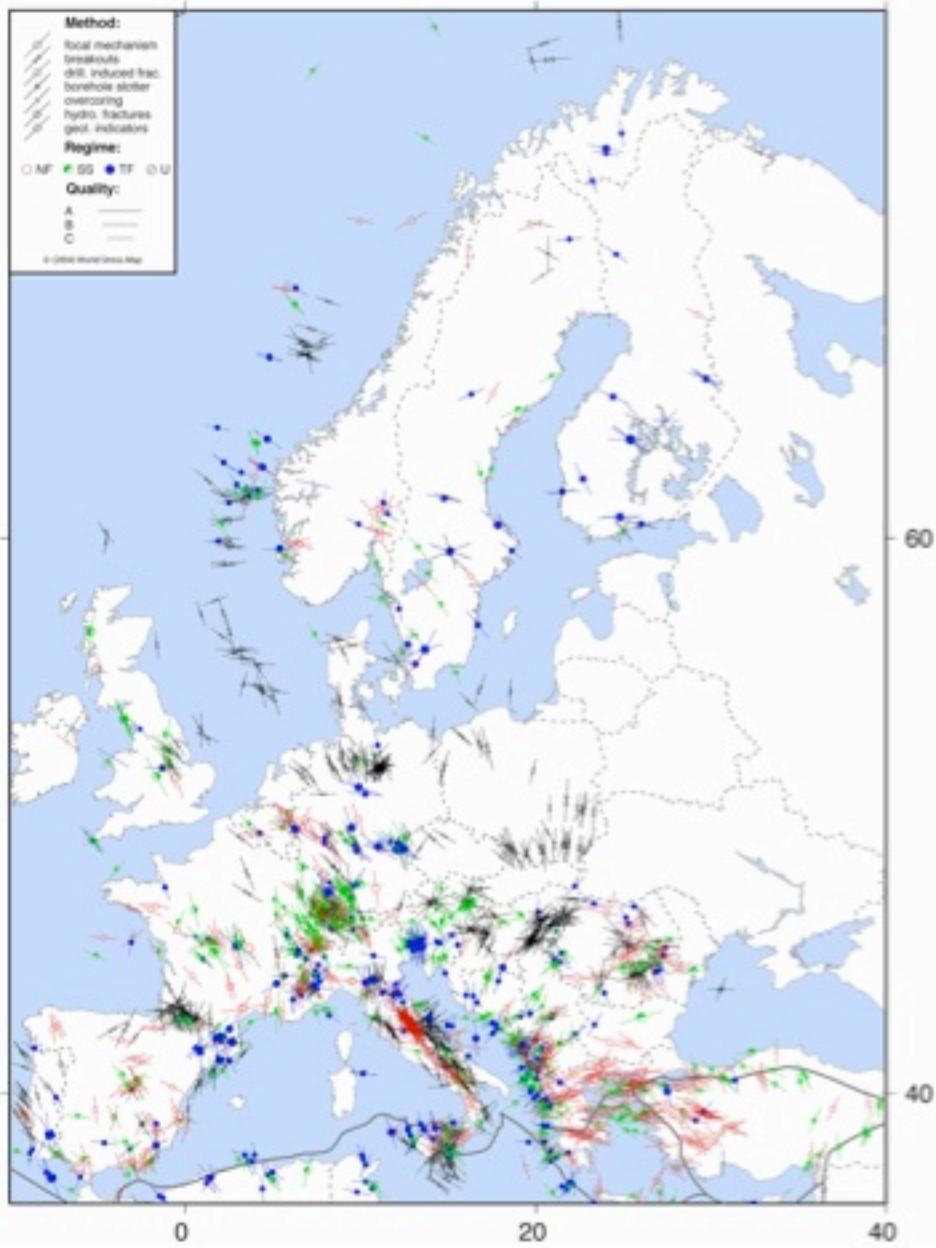
Caldera elongation



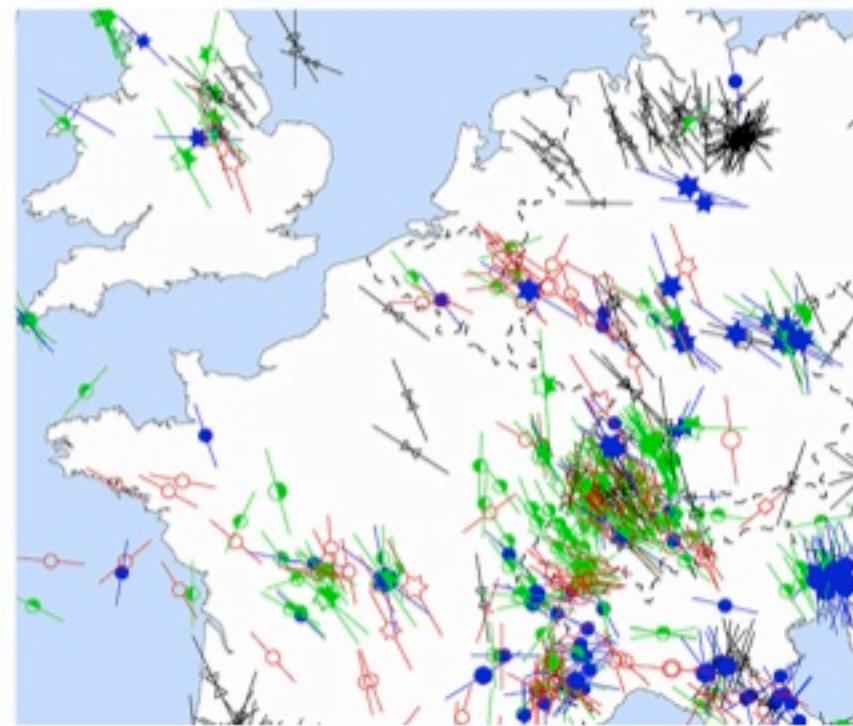


Bosworth et al., 2003

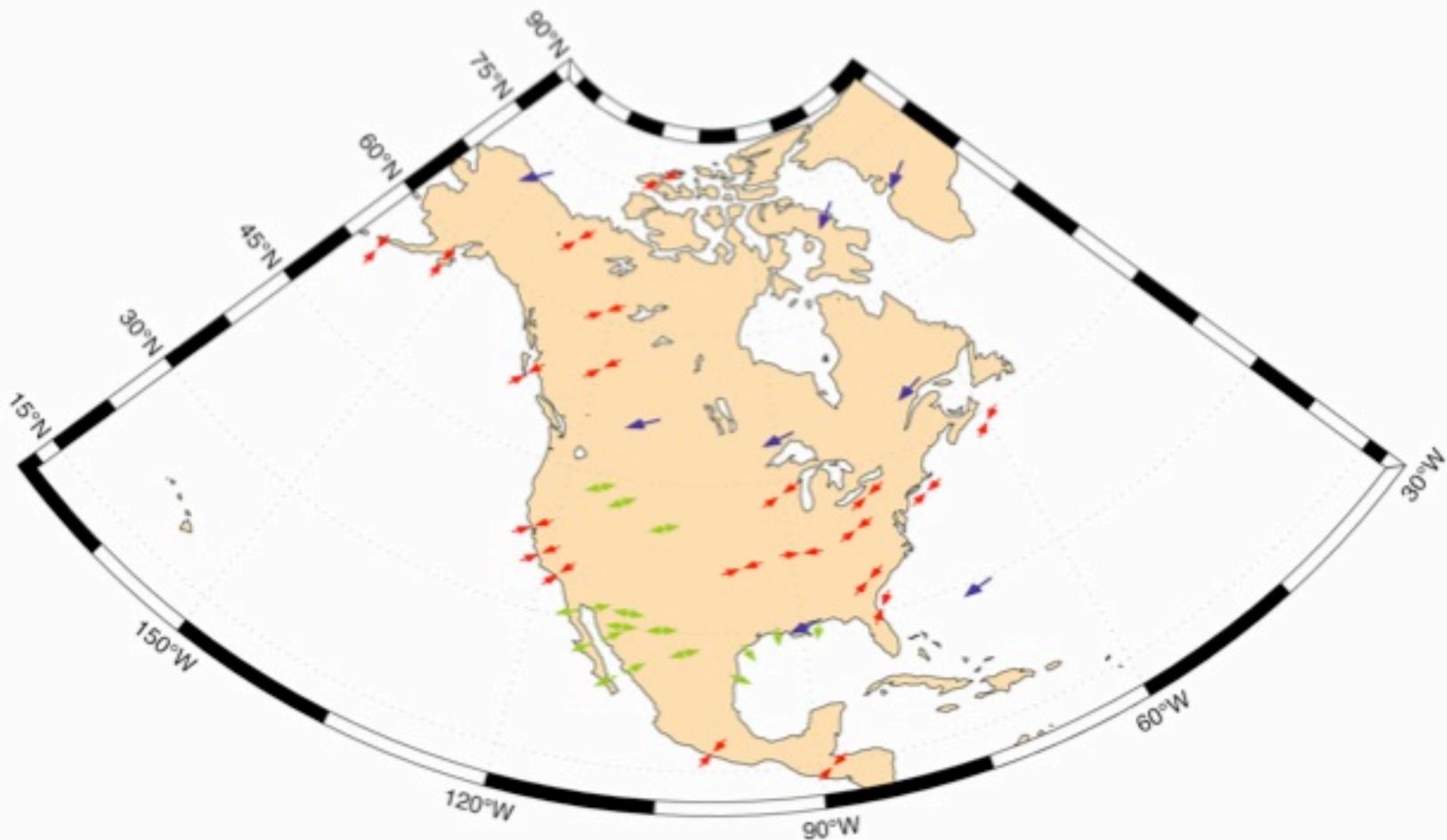
(2) Tectonic stress - field provinces

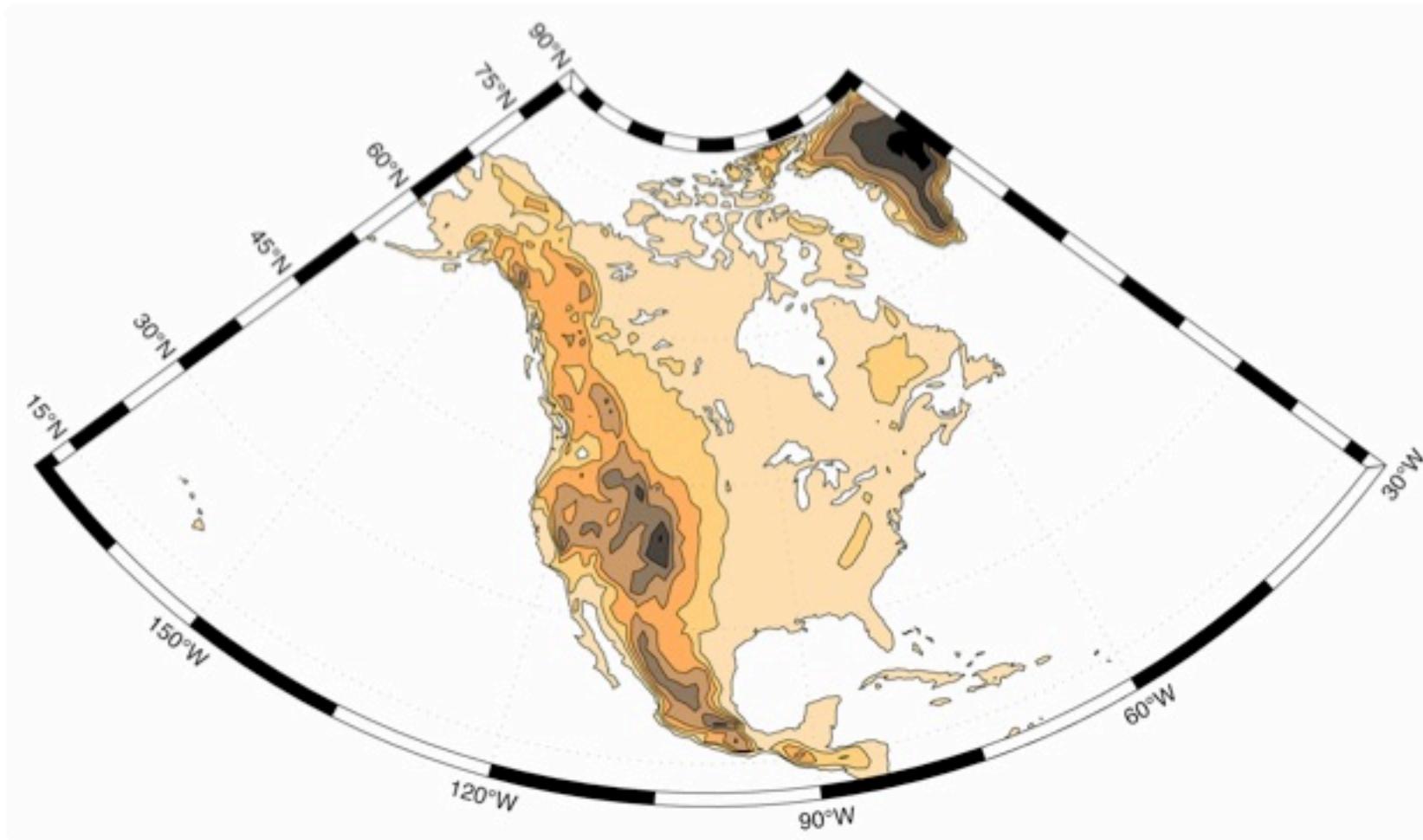


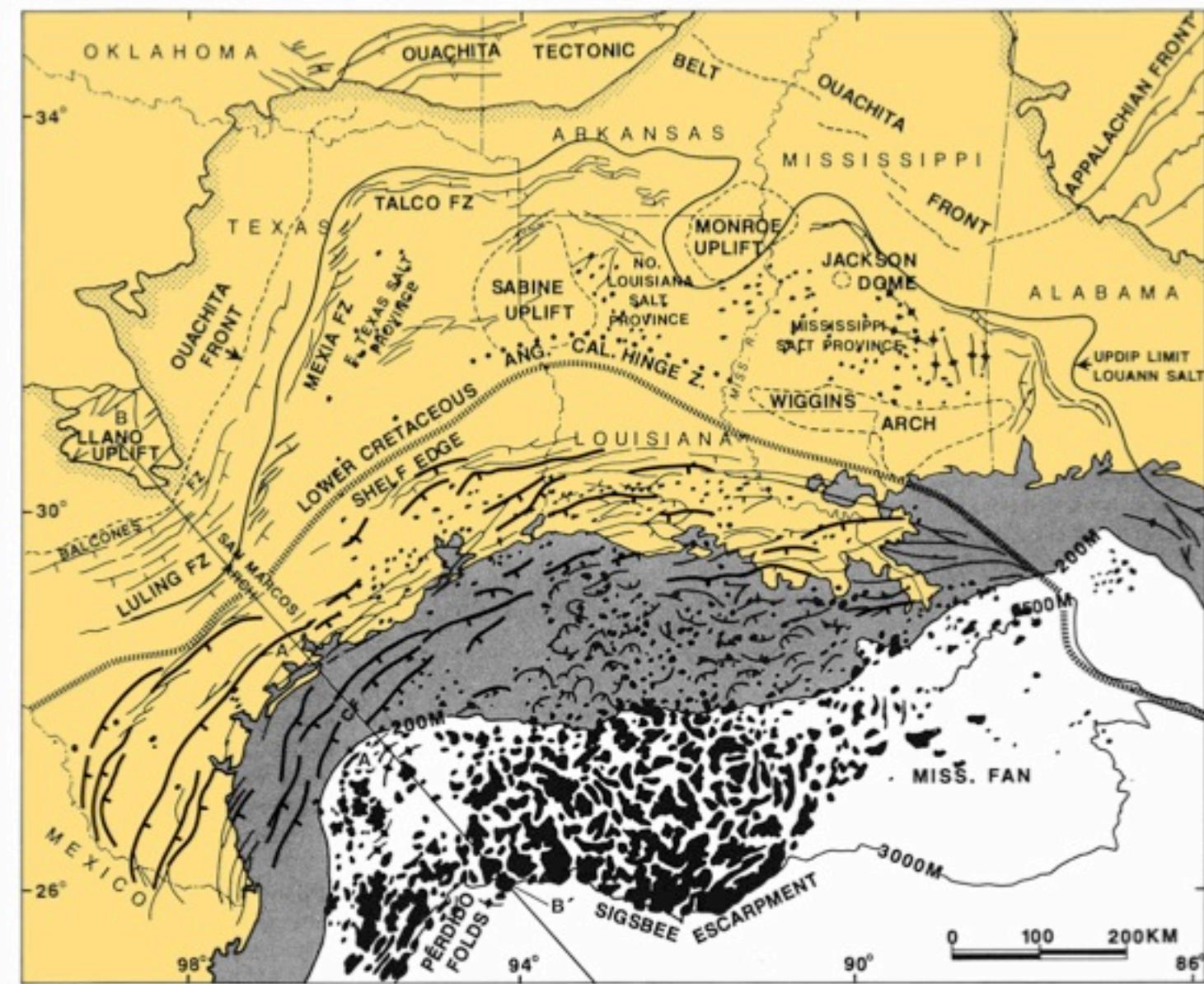
Europe

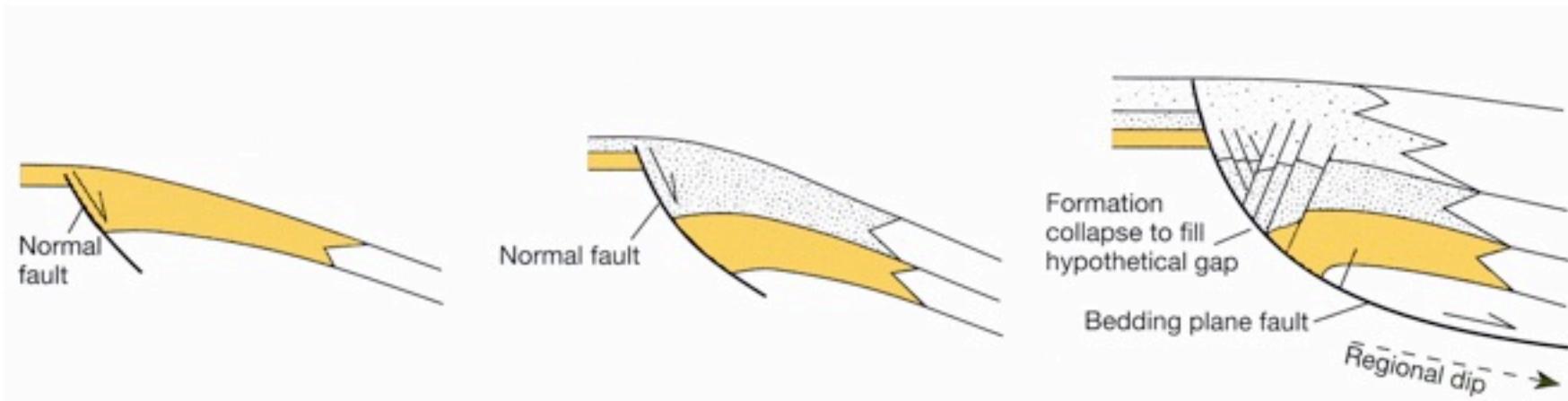
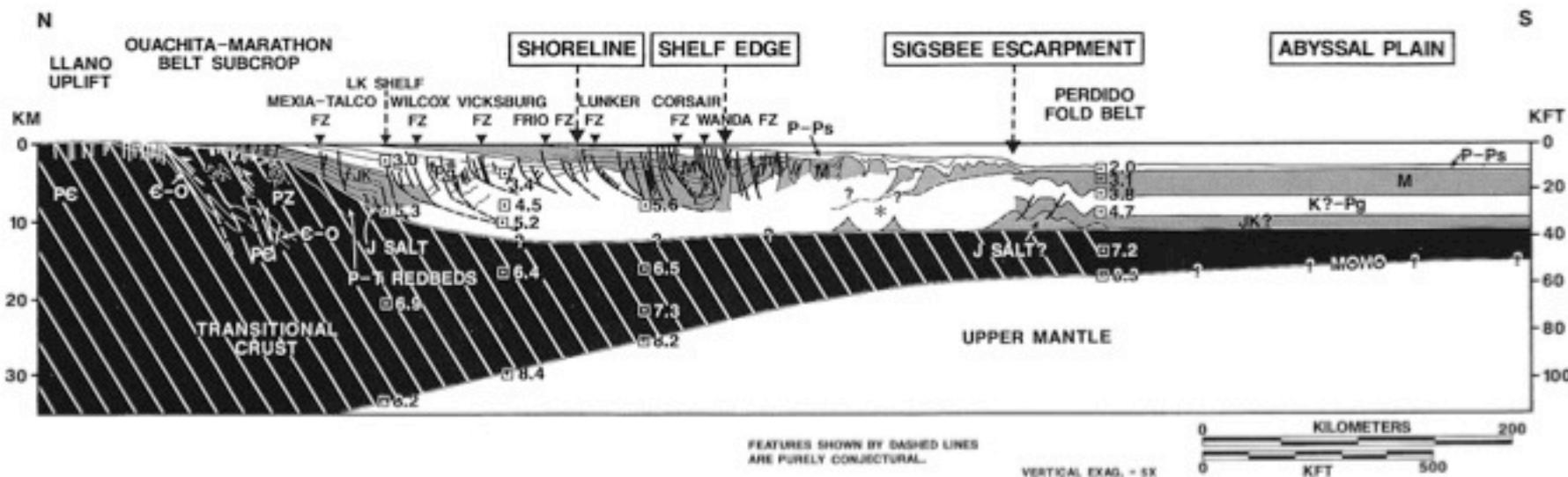


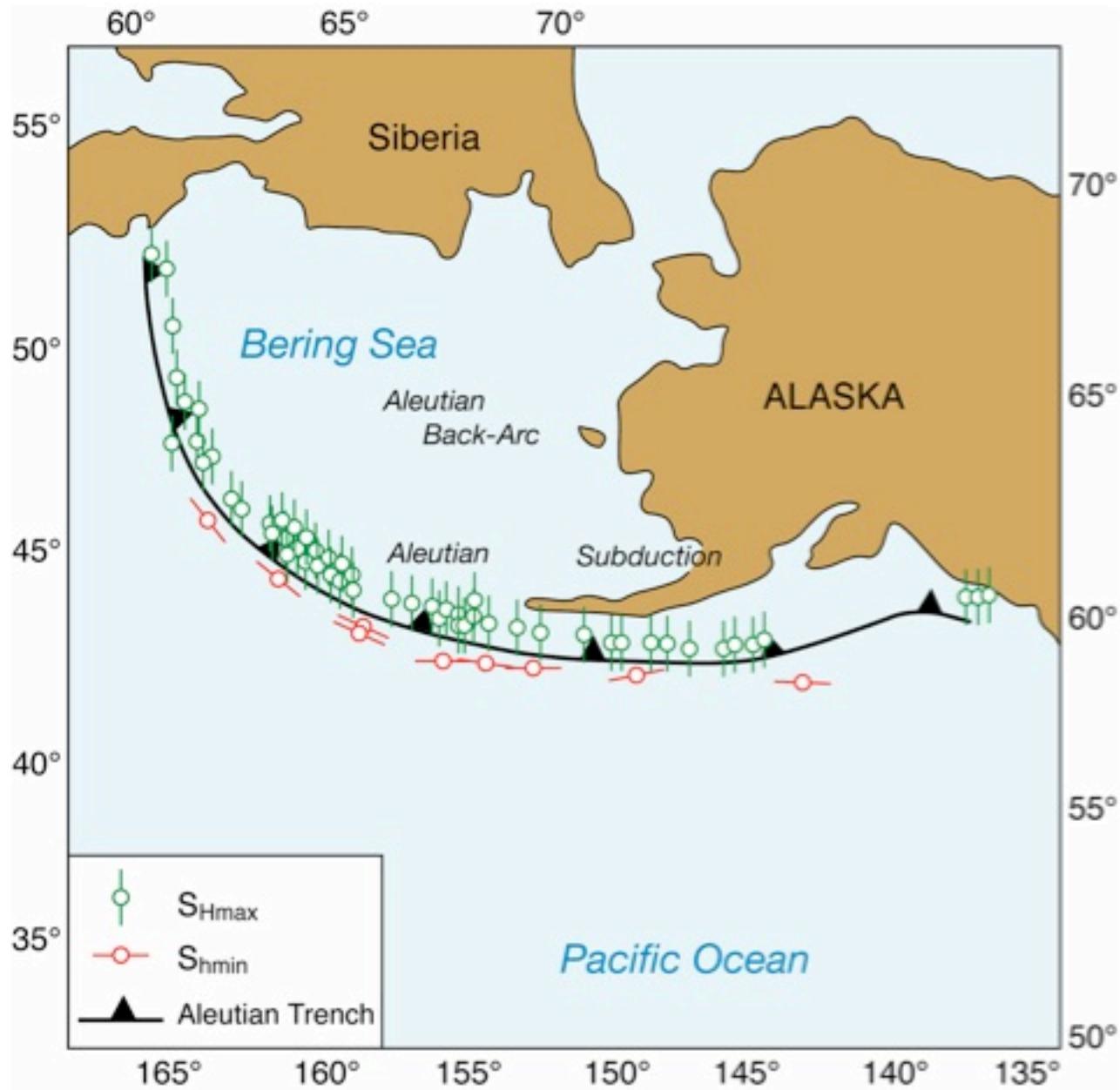
North America

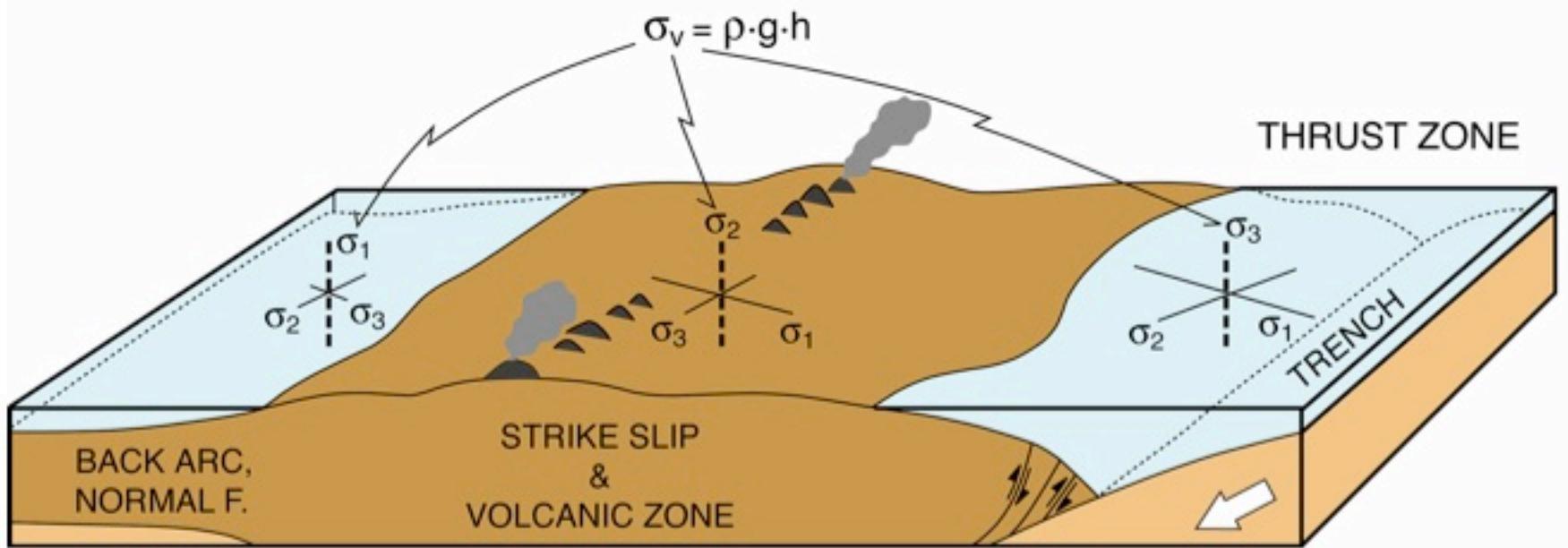




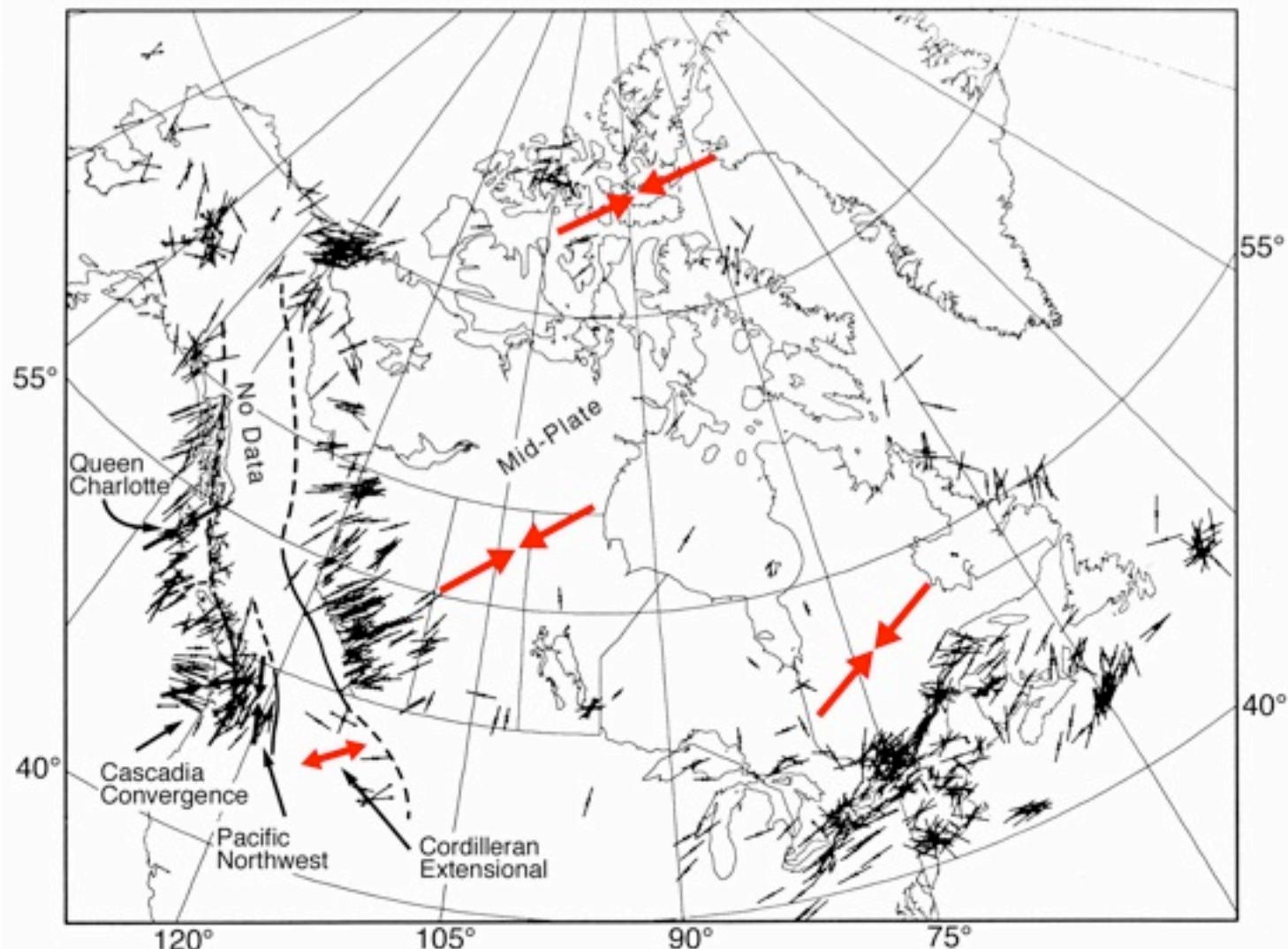




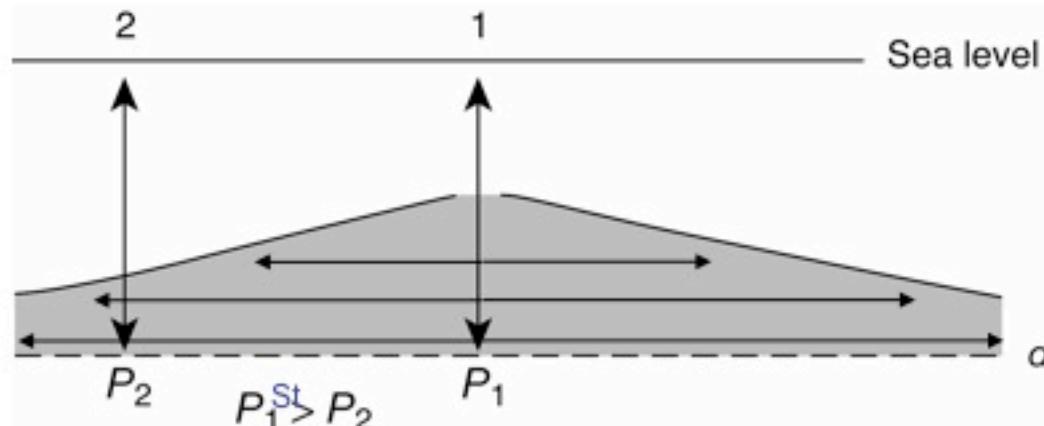




Nakamura and Uyeda, 1980



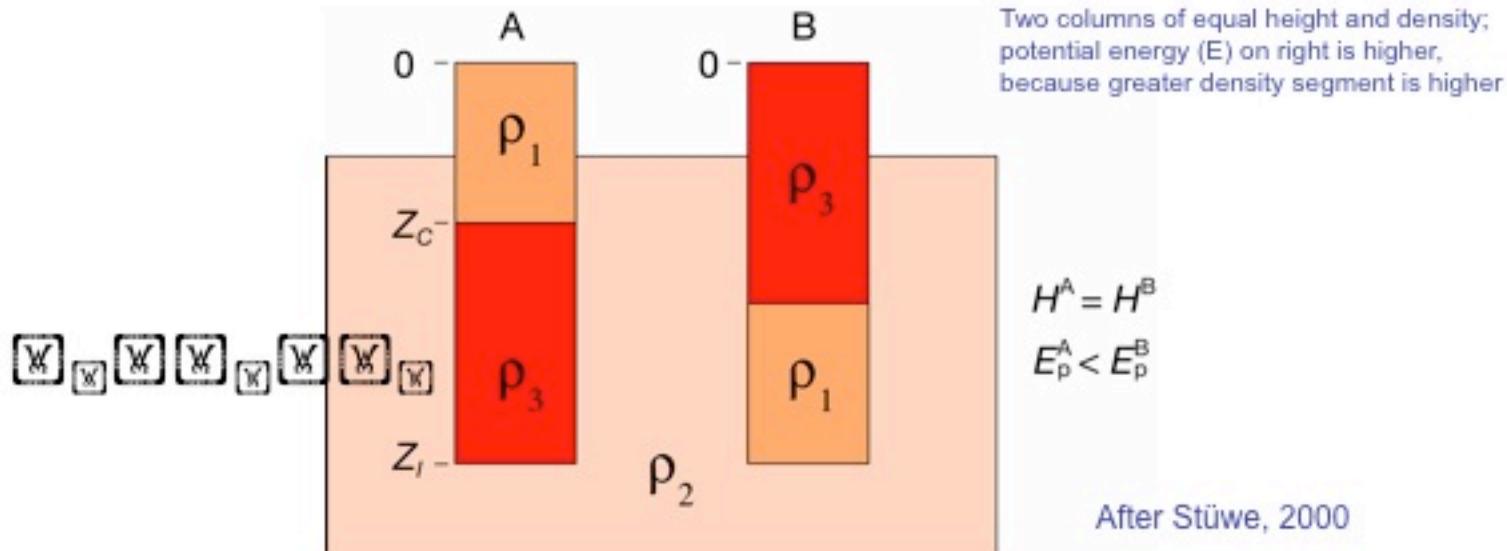
Ridge push



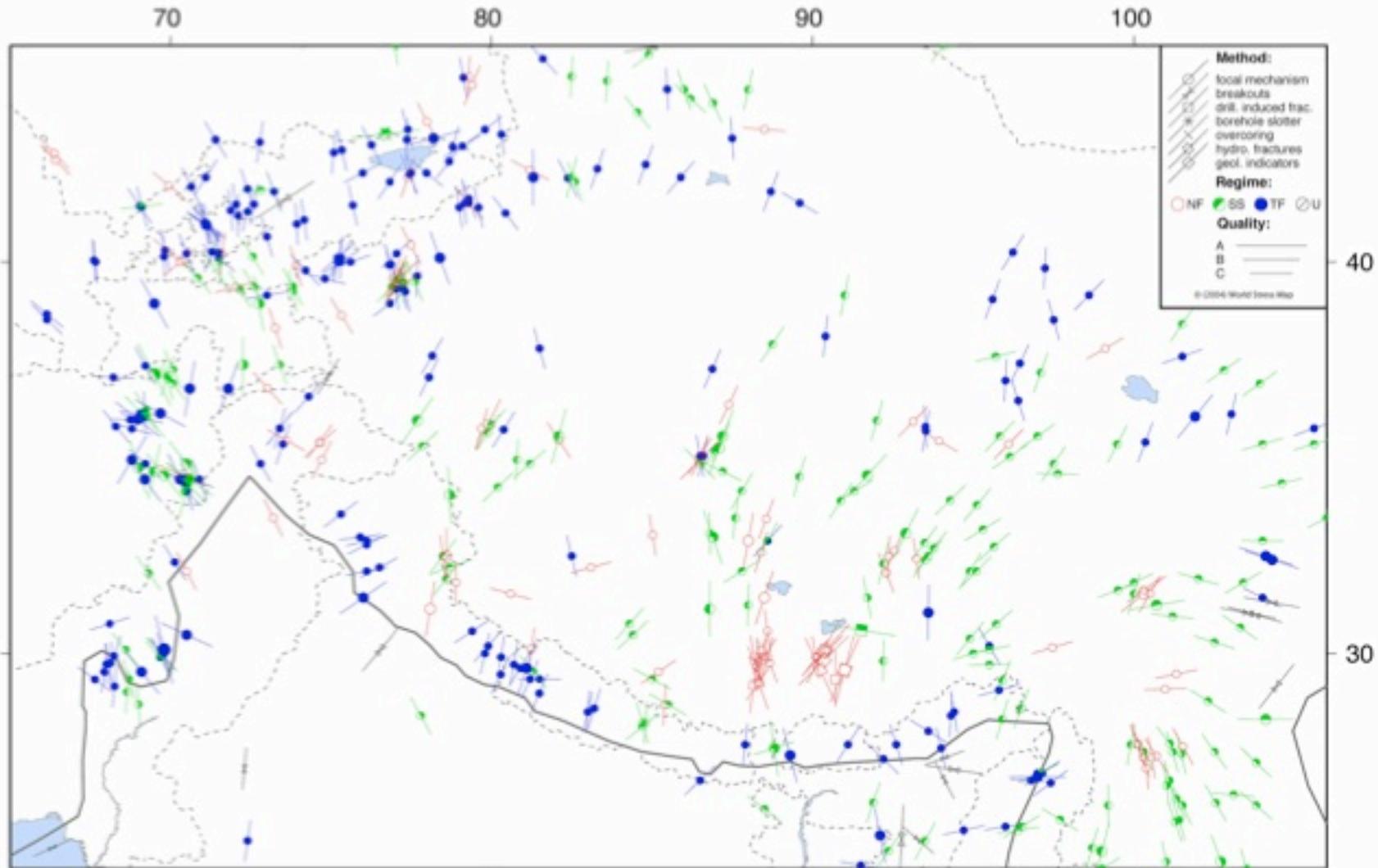
P_1 = Pressure at depth d below locality 1

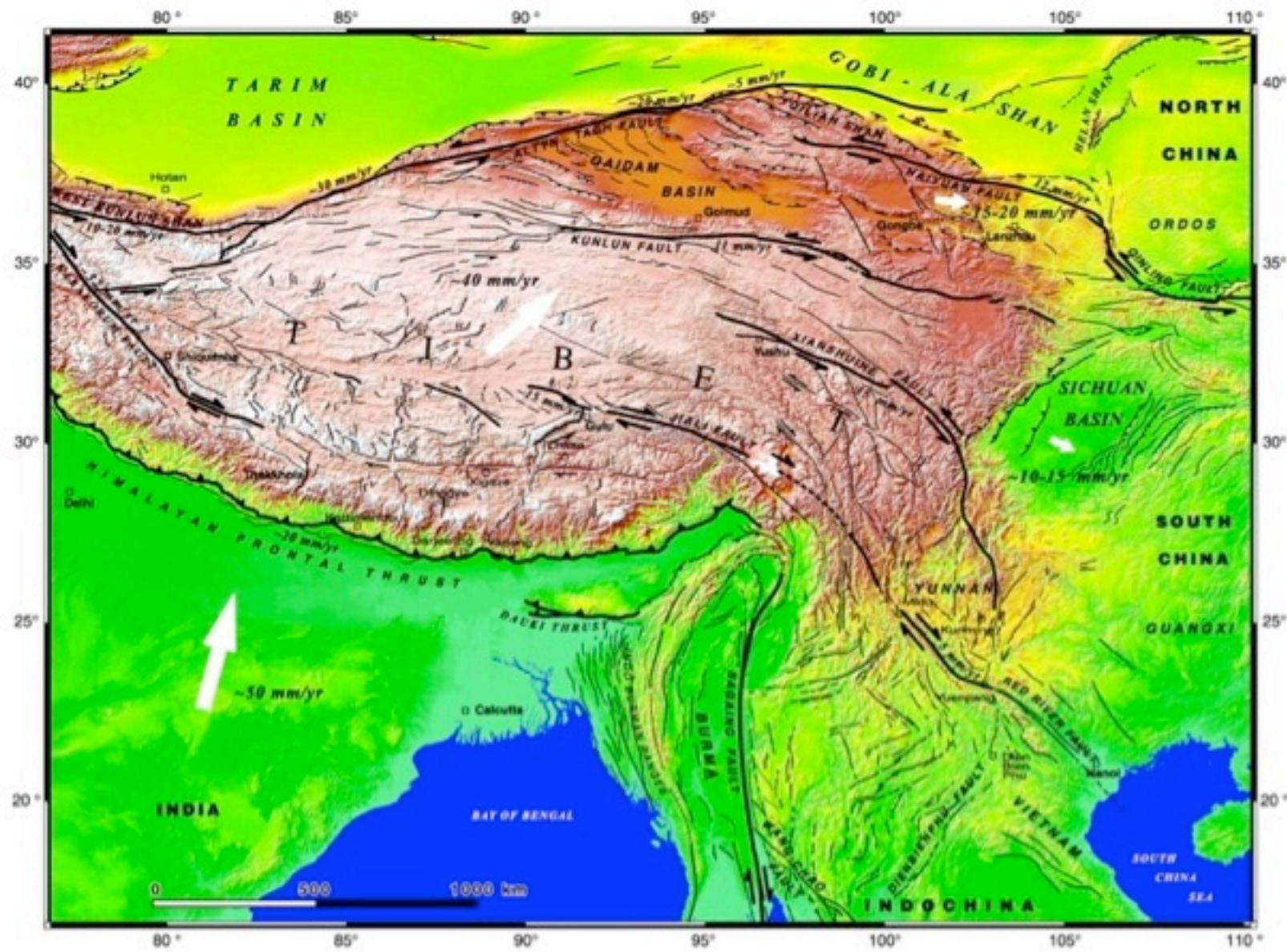
P_2 = Pressure at depth d below locality 2

After Moores and Twiss, 1997

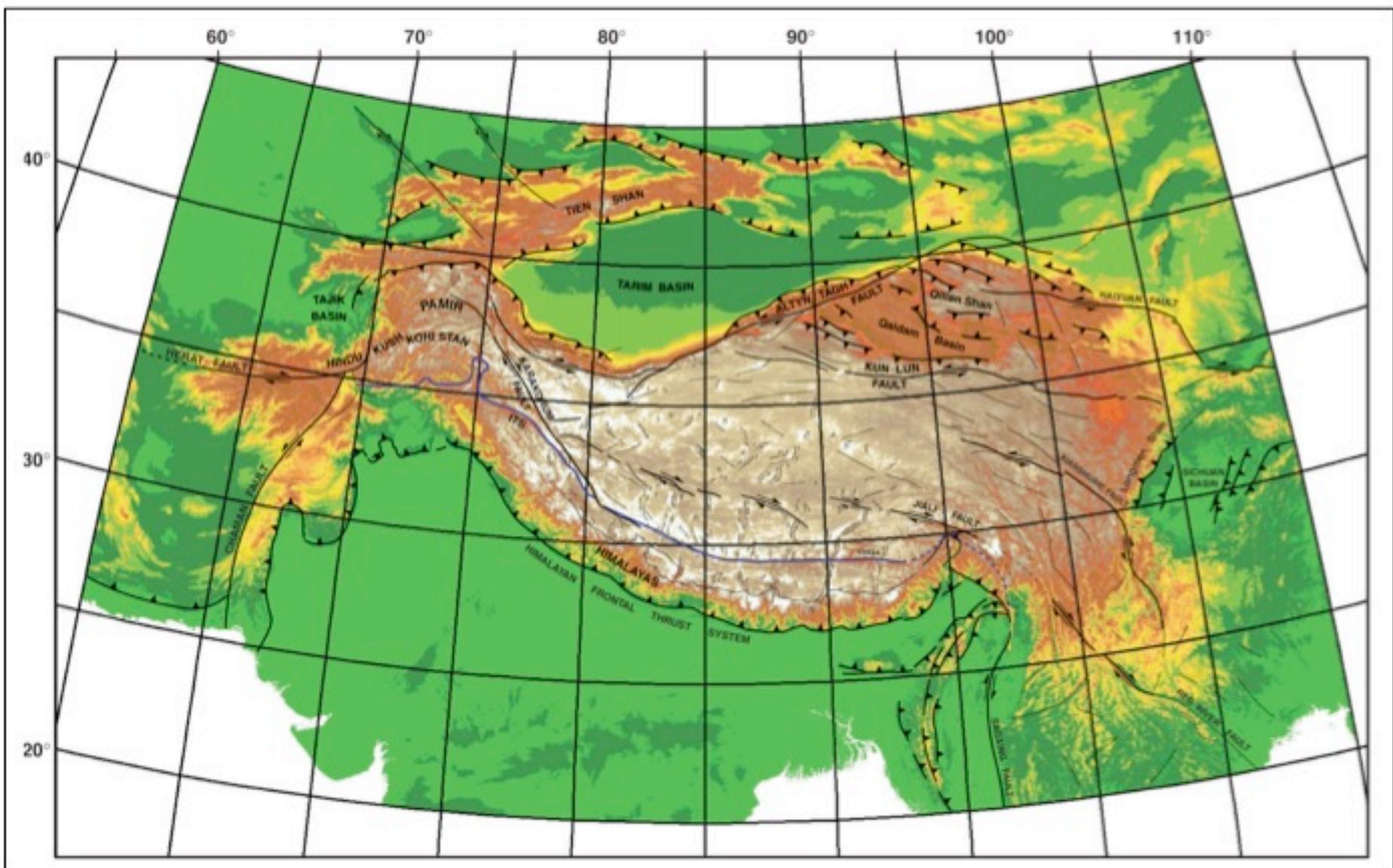


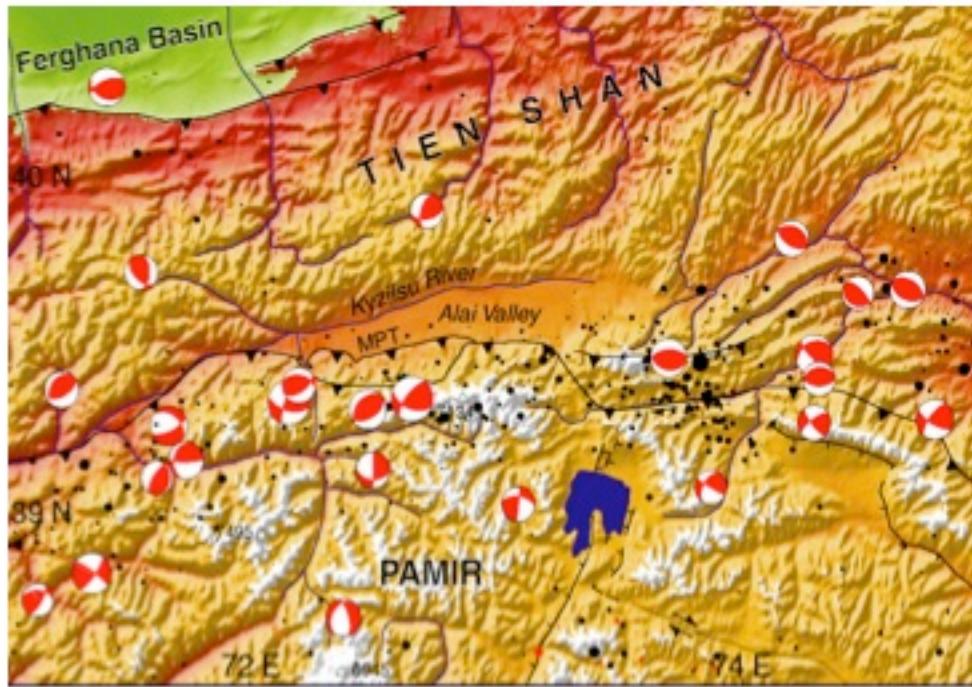
The World Stress Map





Central Asia





Arrowsmith and Strecker, 1999

80

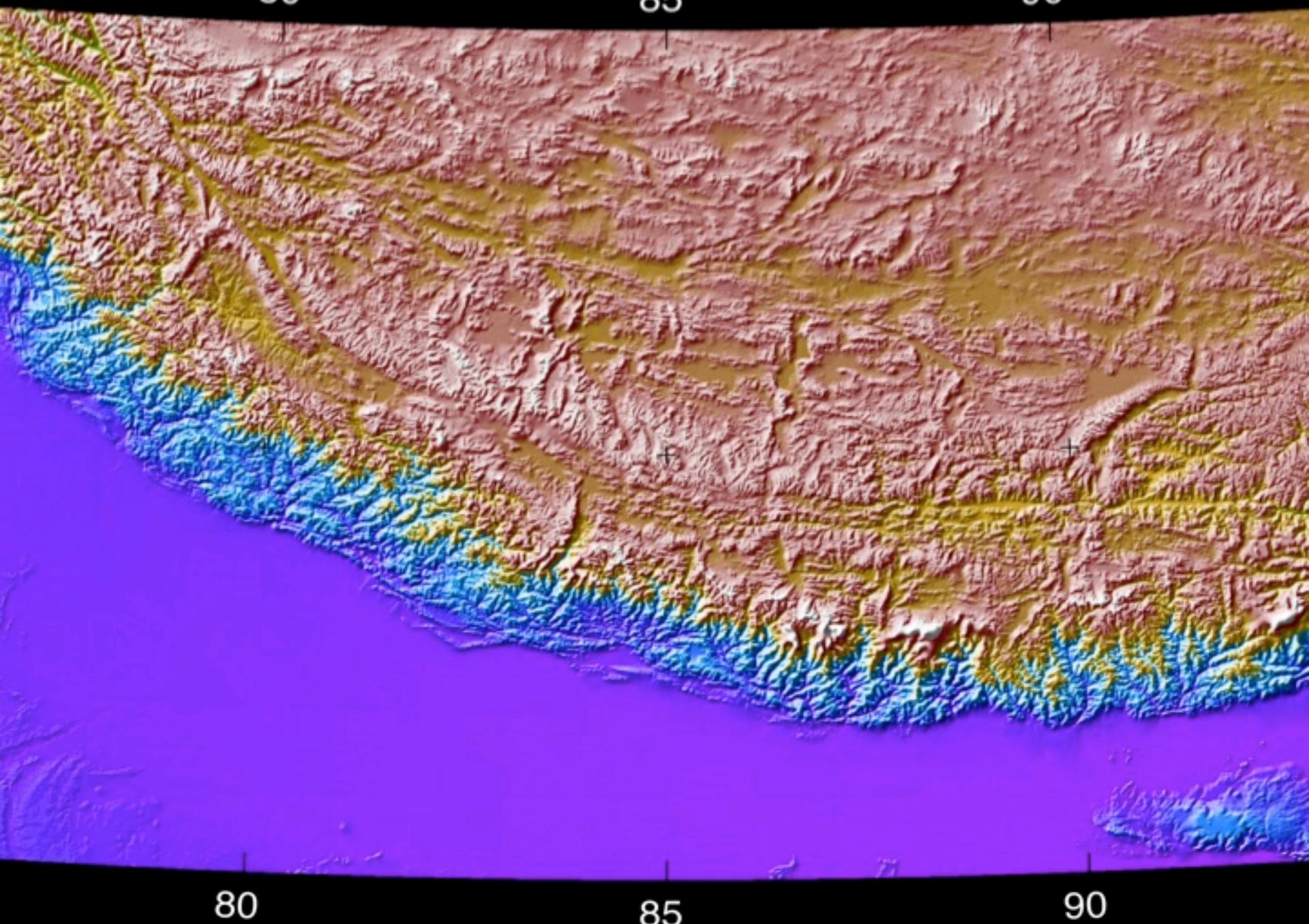
85

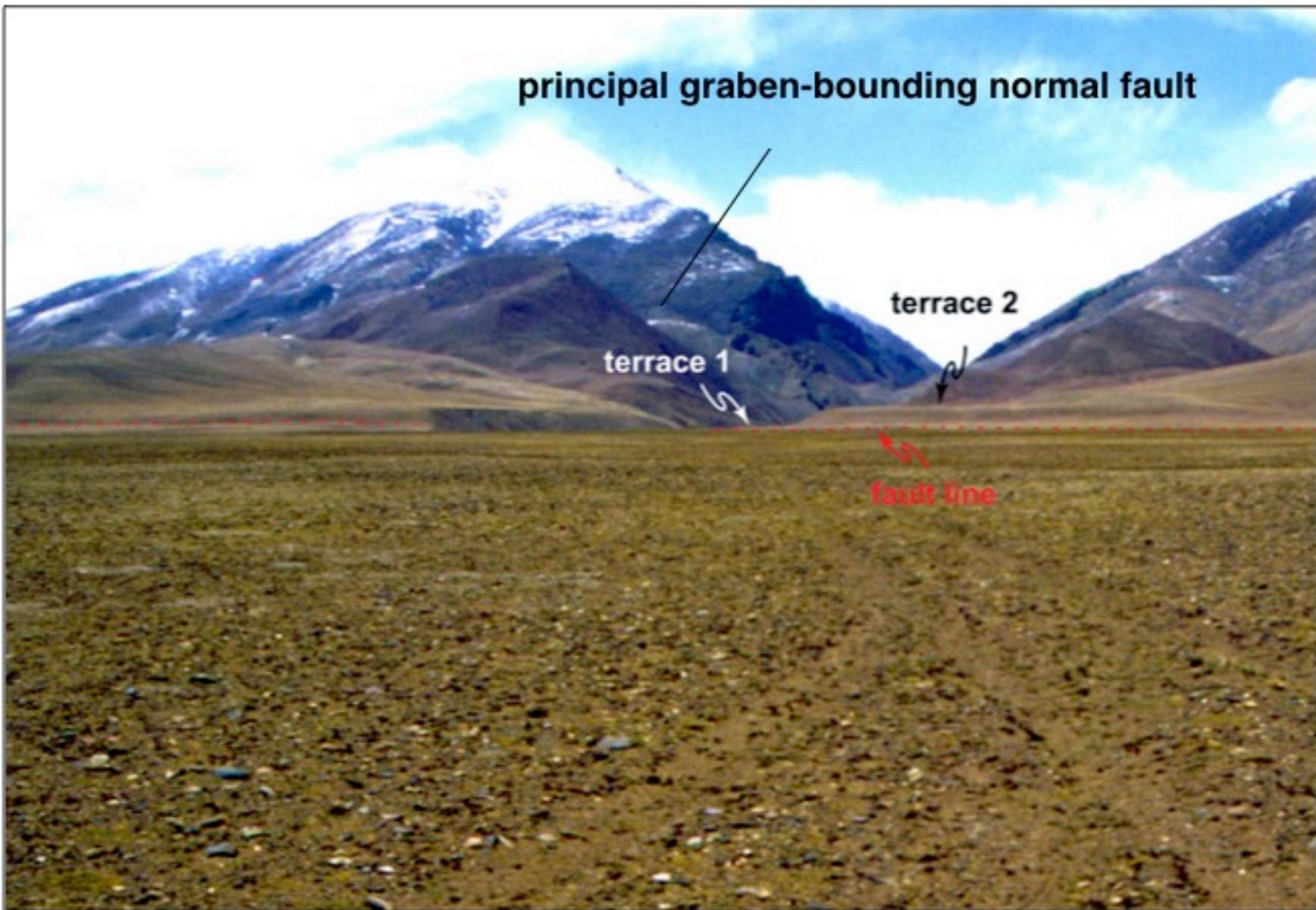
90

80

85

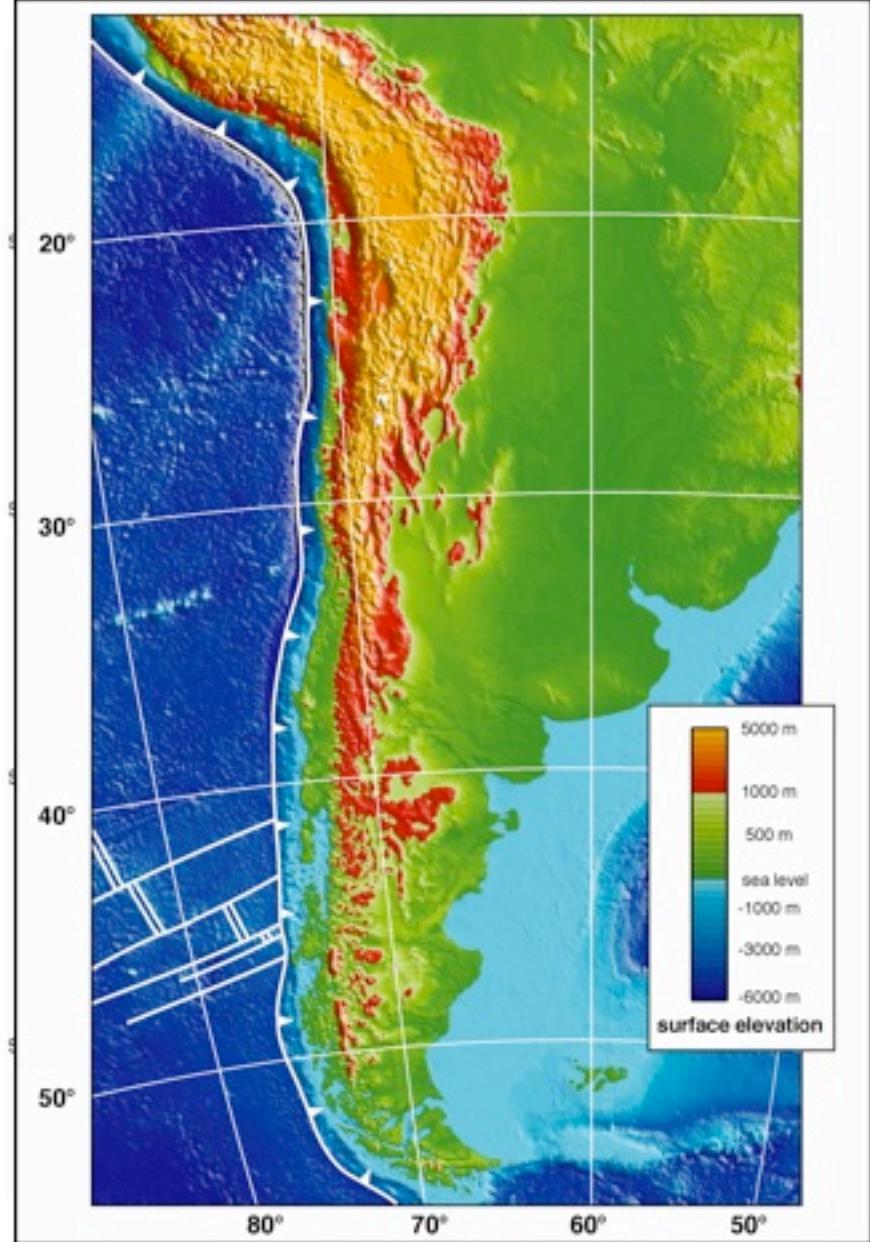
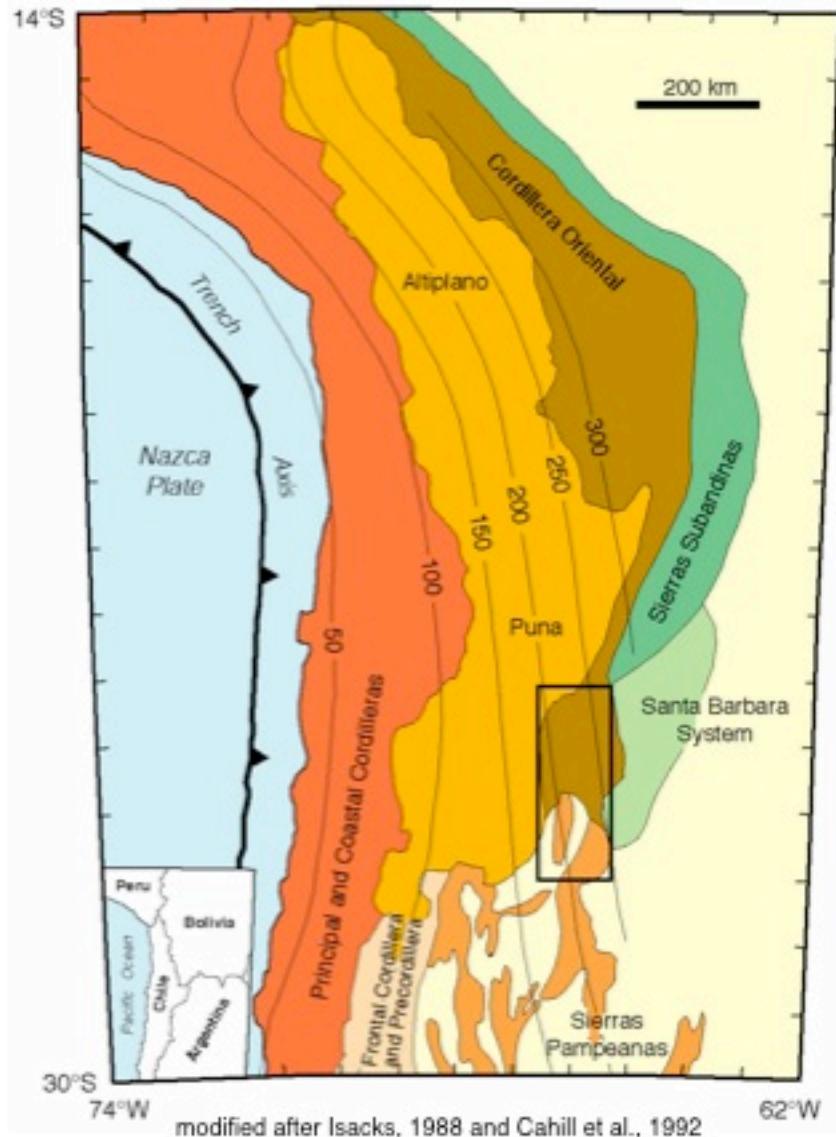
90

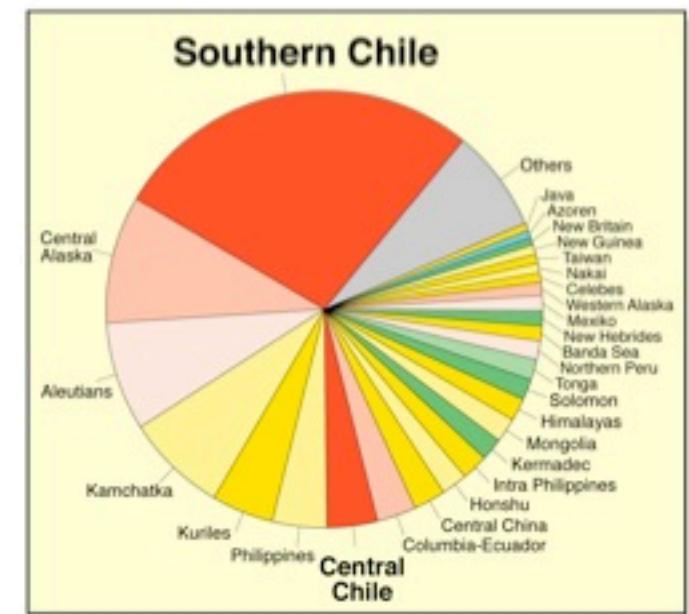
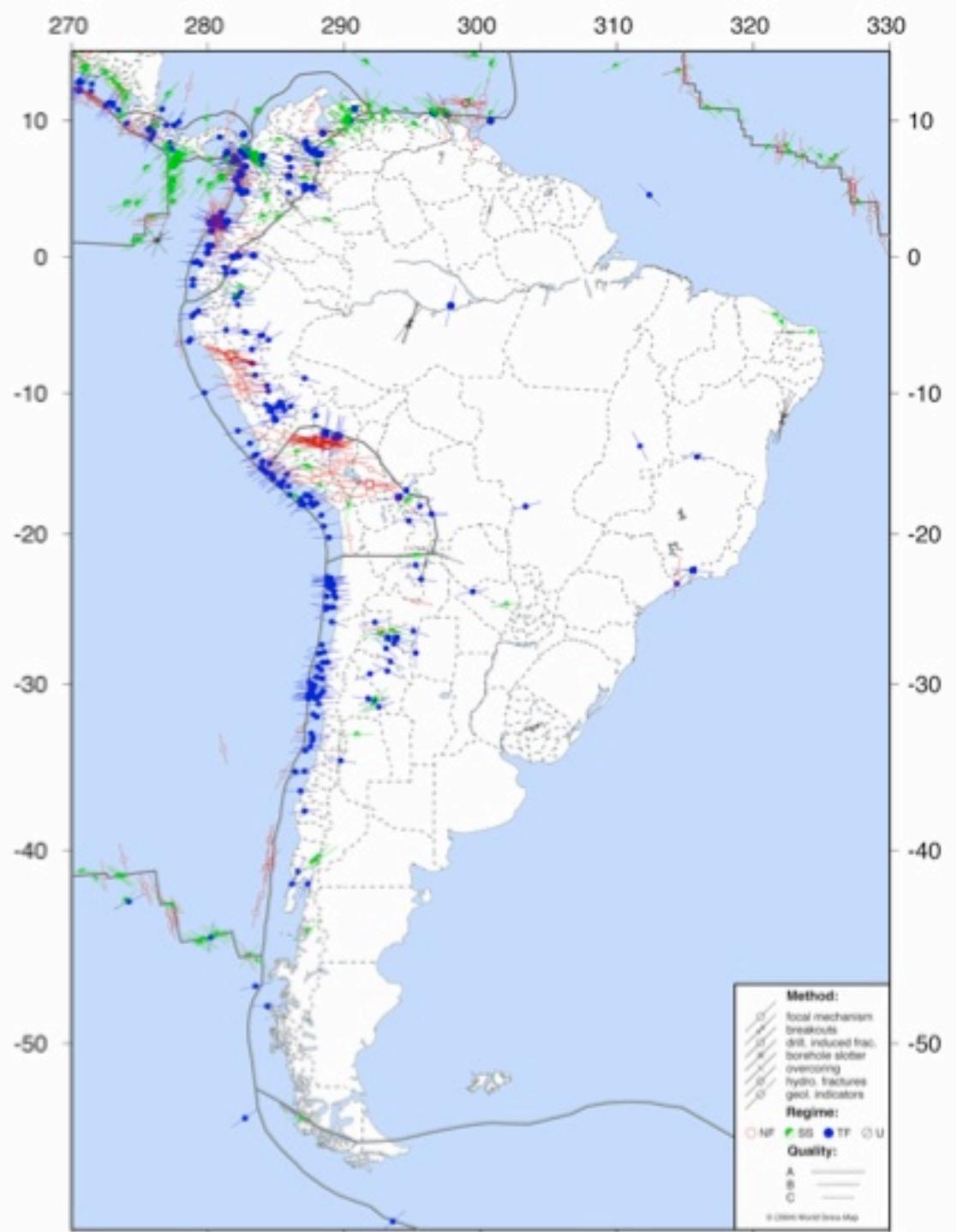




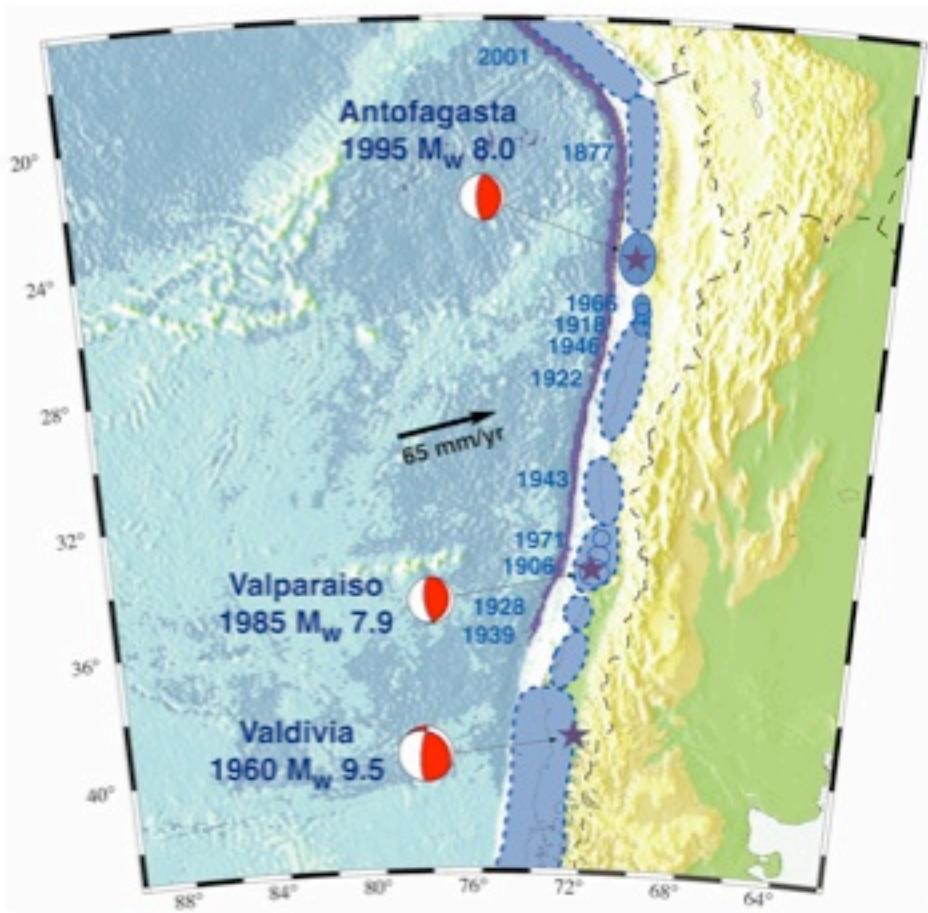
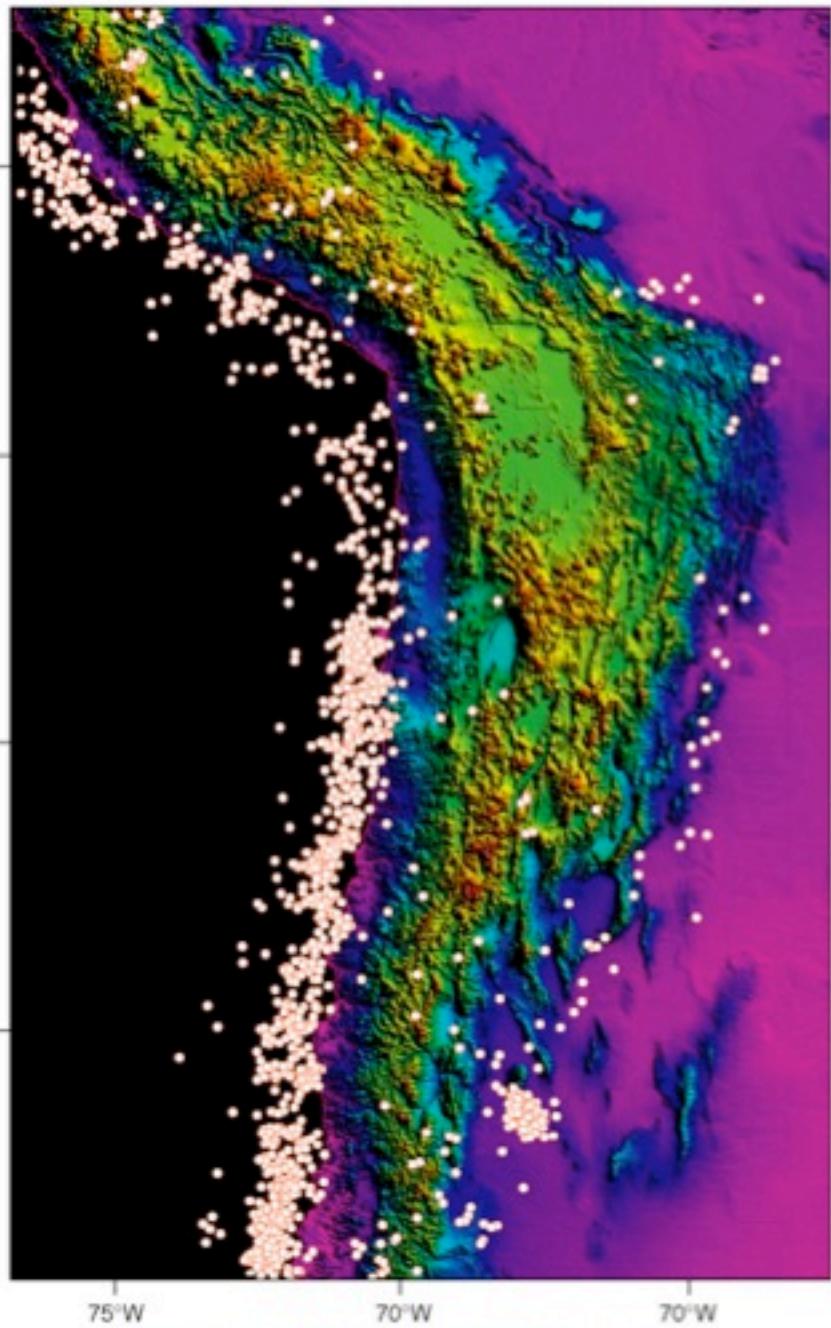
Shuang Hu graben: view of the SW graben margin, looking W

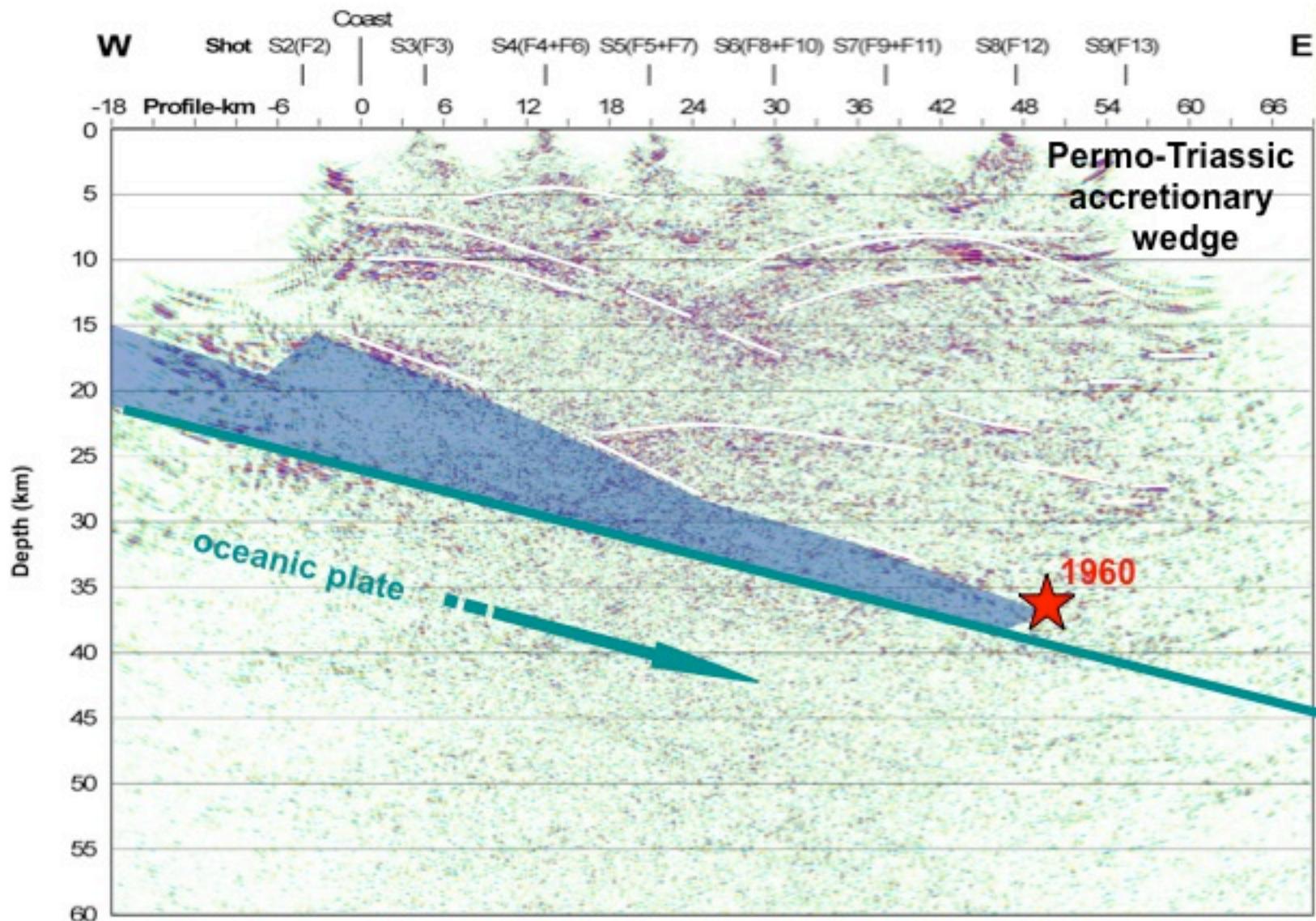
South America





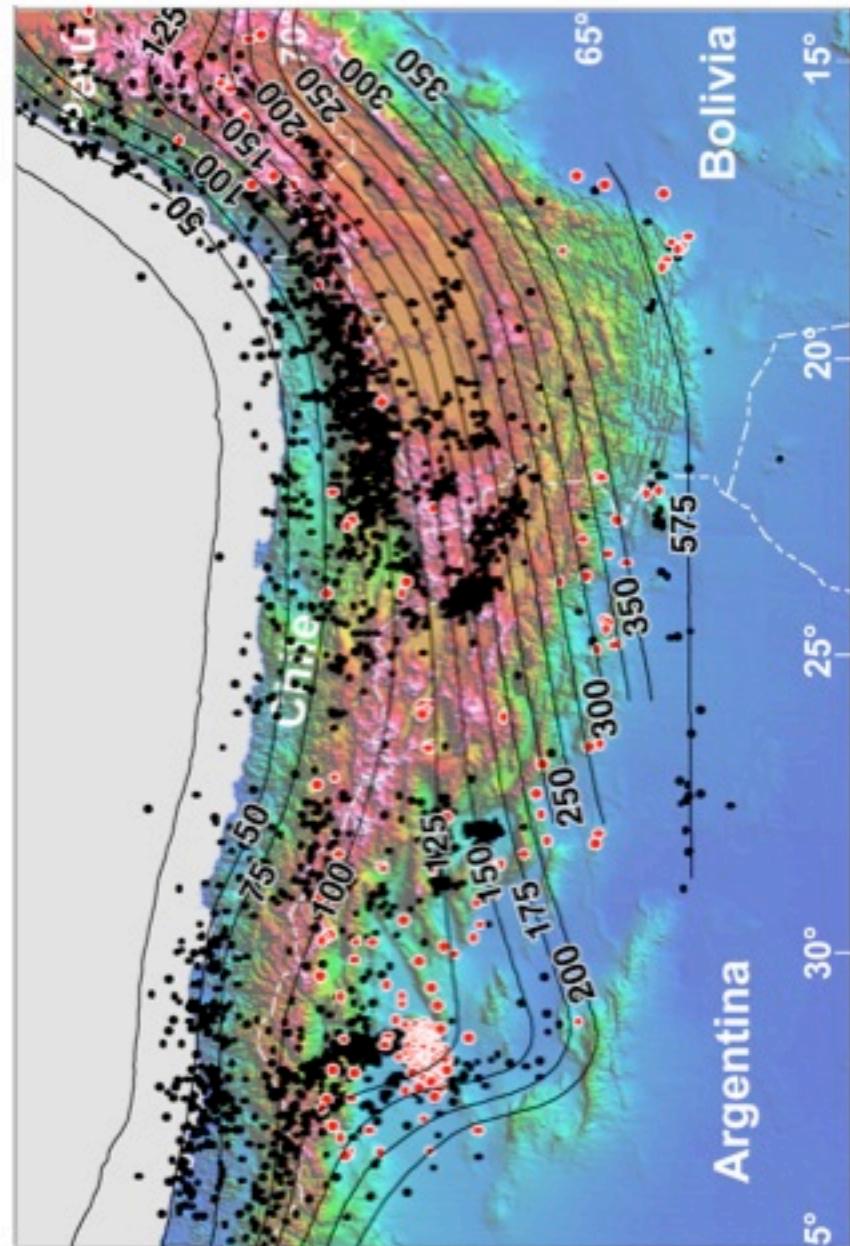
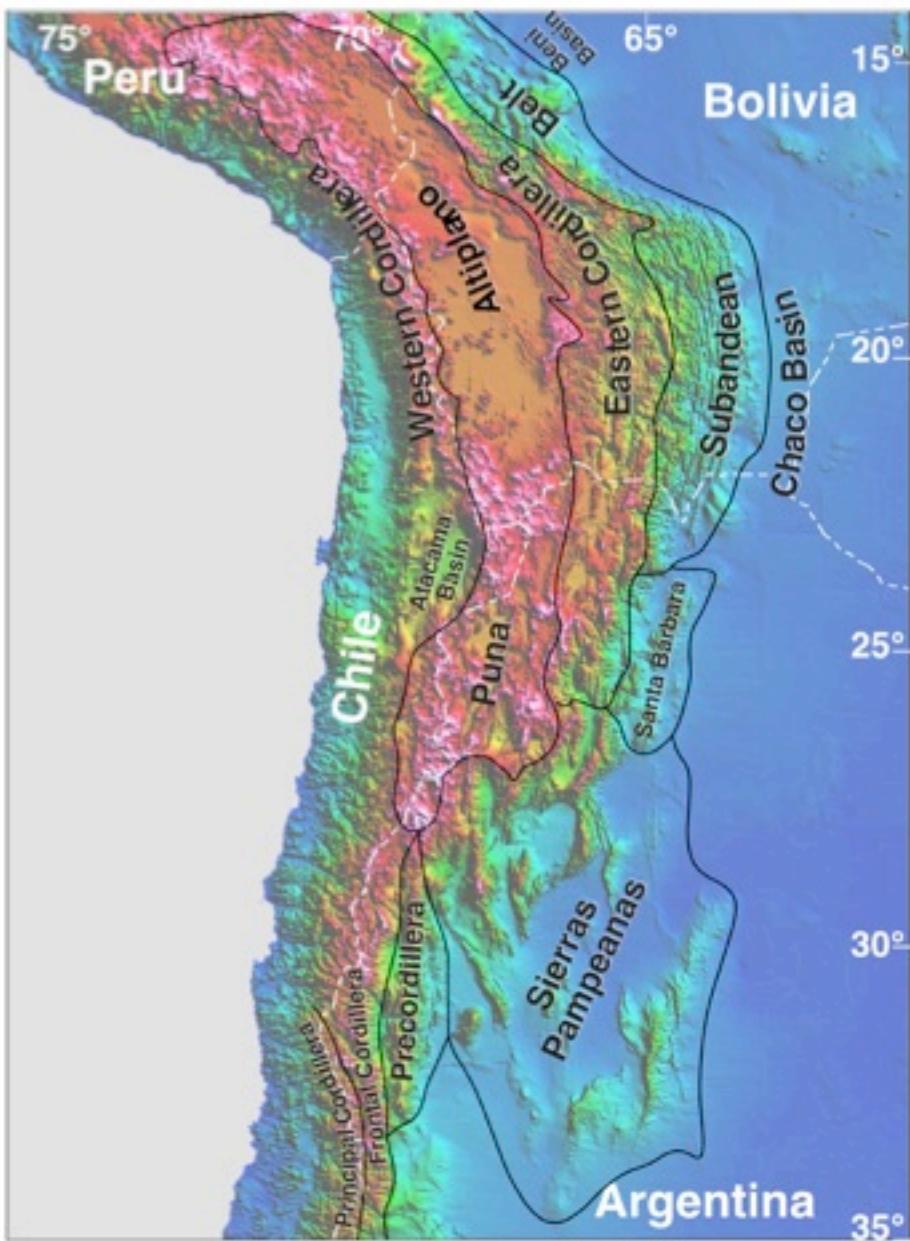
Seismic energy release in the 20th century





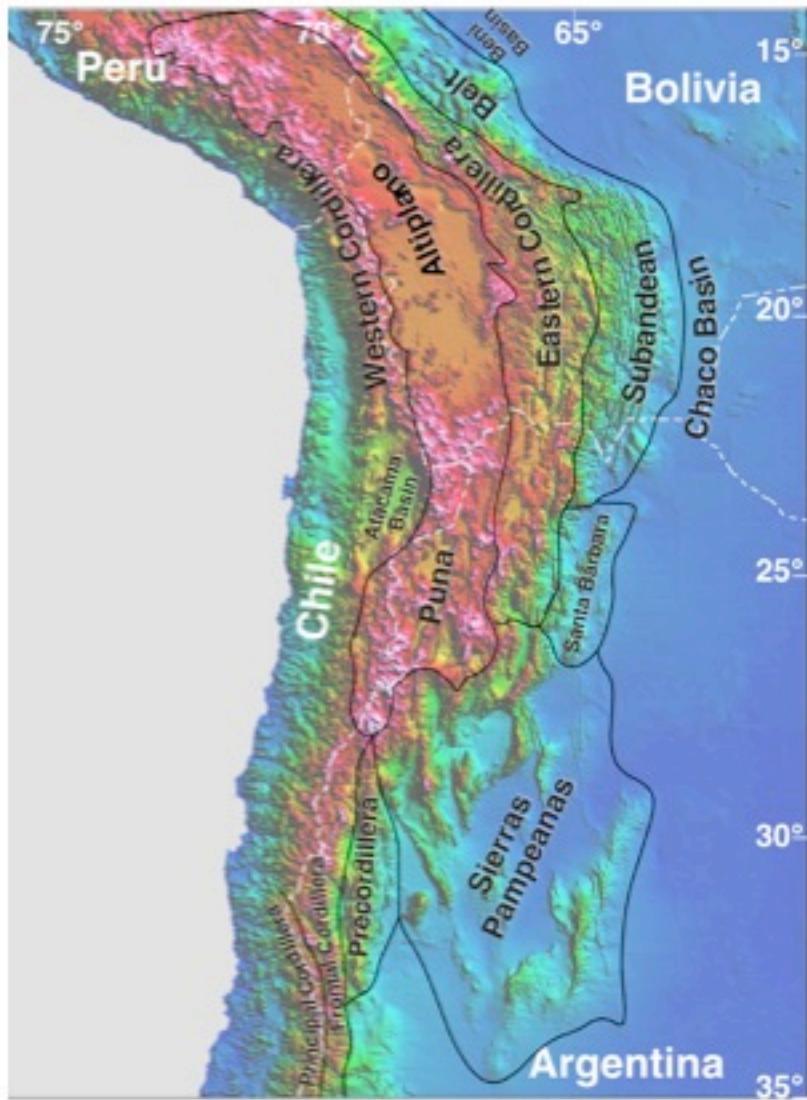
(modified after Krawczyk & The SPOC Team 2003; regional seismicity: Bohm et al. 2002)

The Central Andes

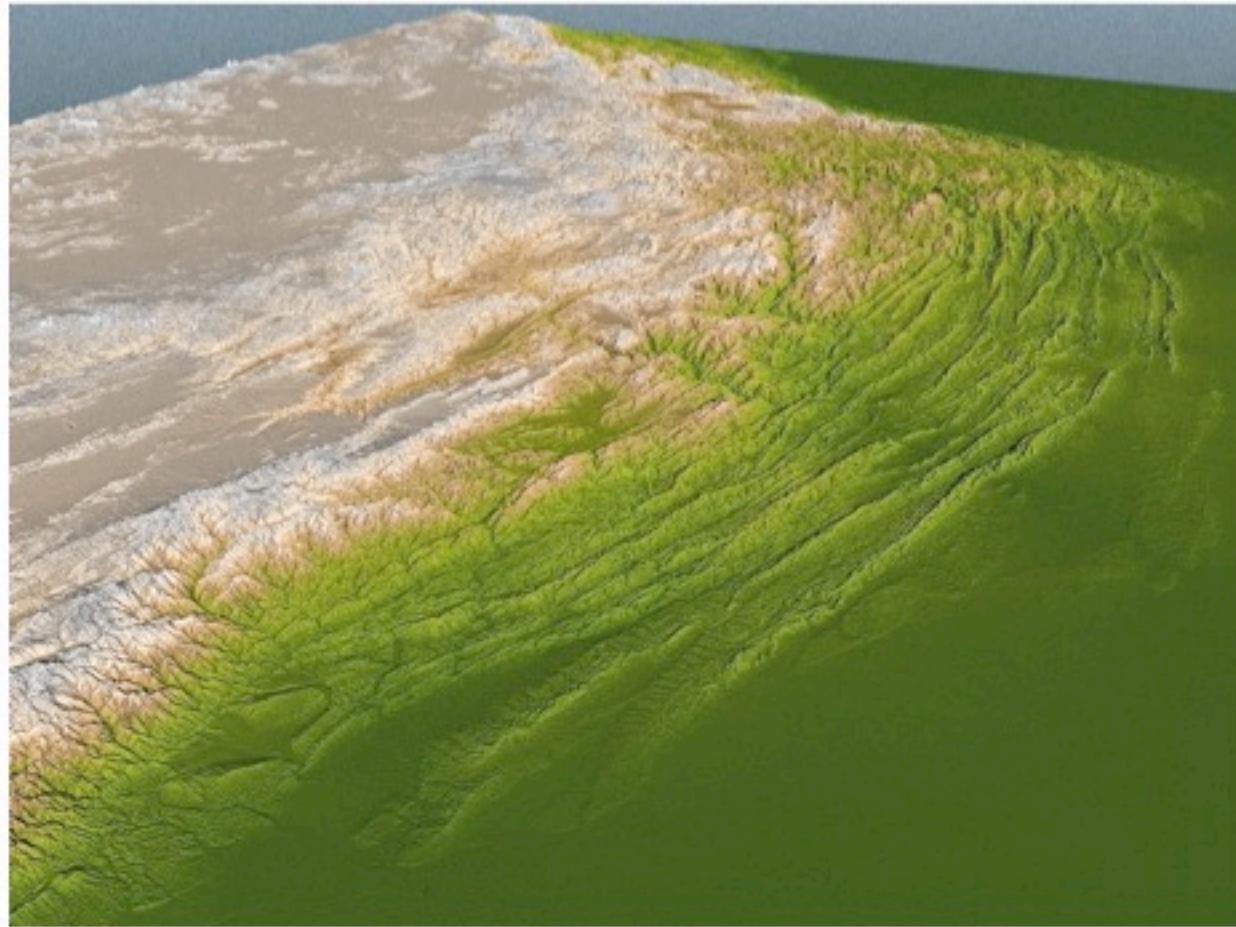
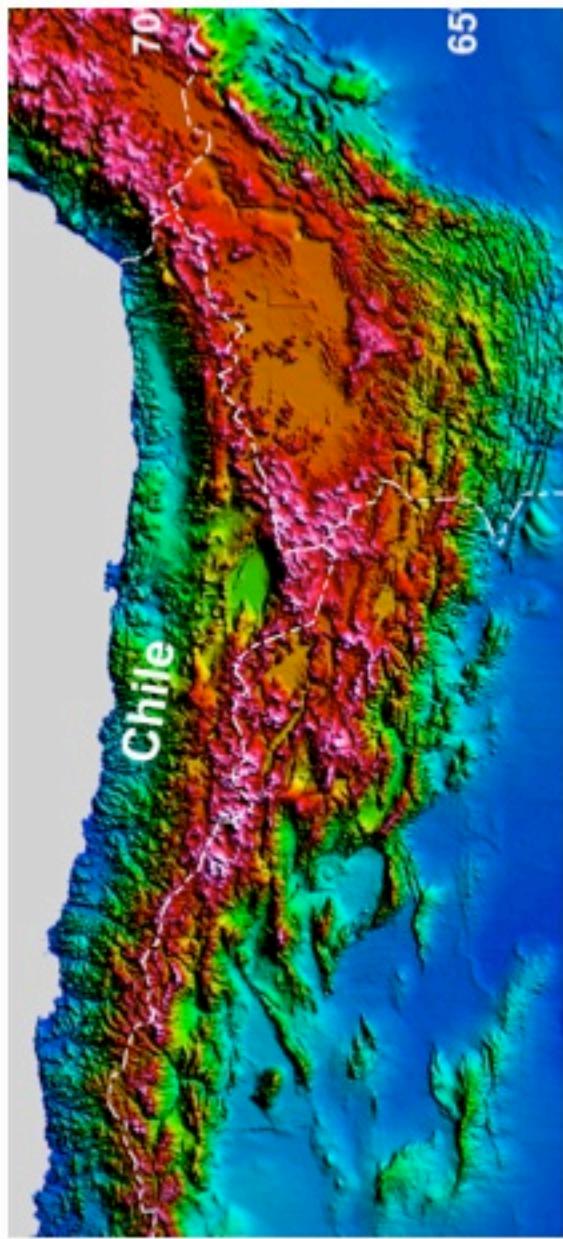


Modified after Jordan et al., 1983 and Cahill et al., 1988

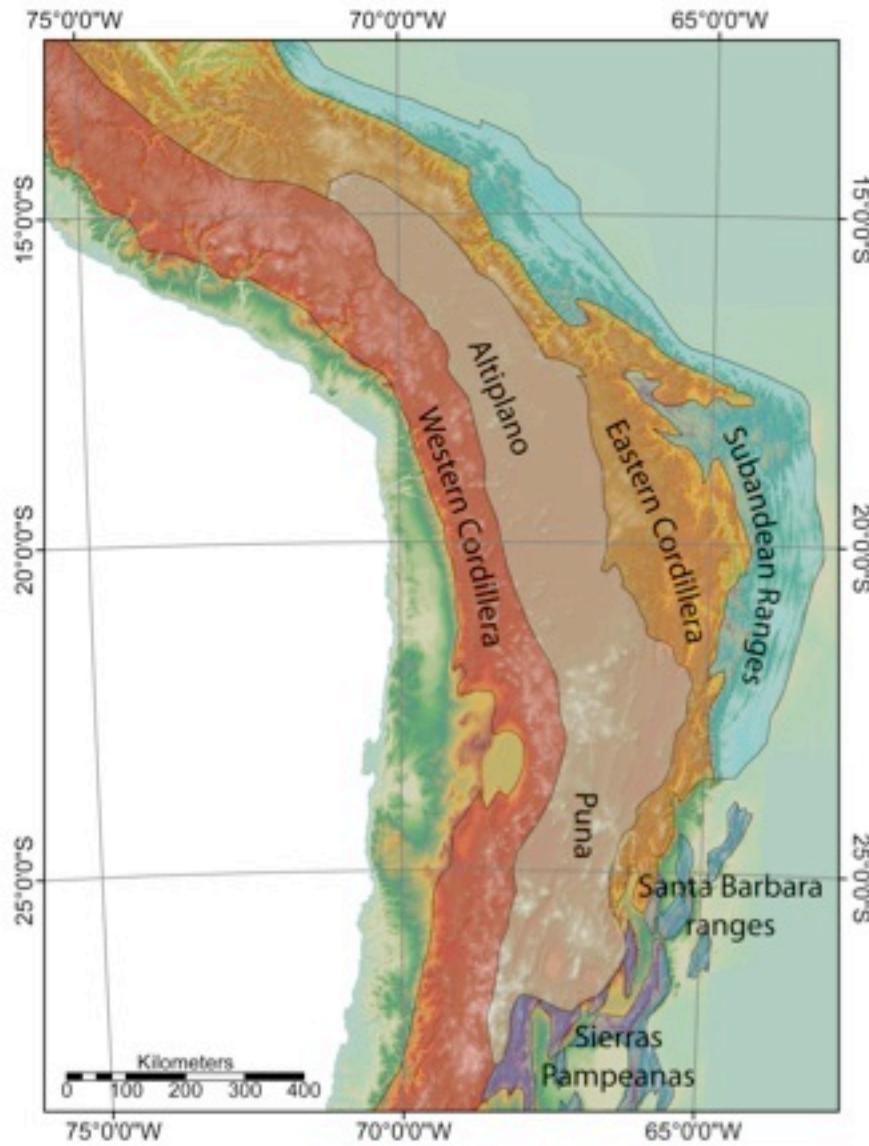
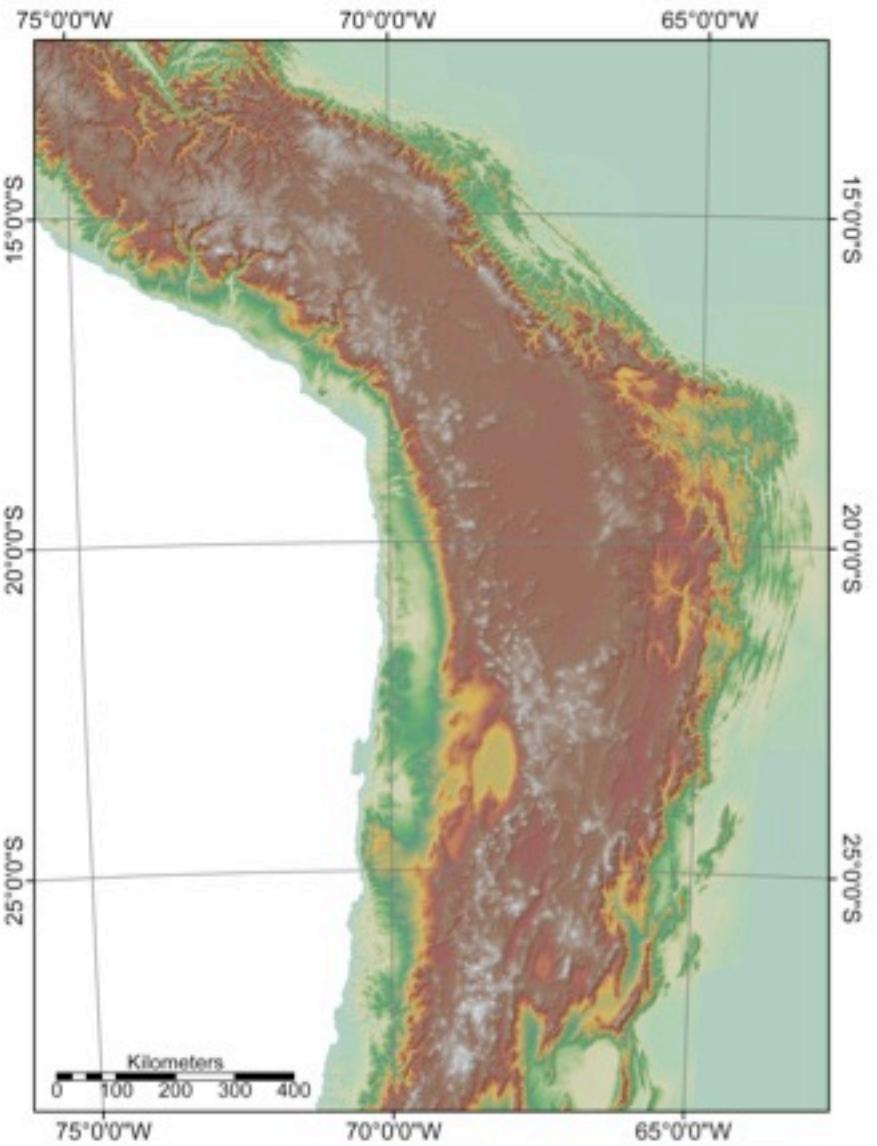
The Flat-Slab Region: Precordillera & Sierras Pampeanas Provinces

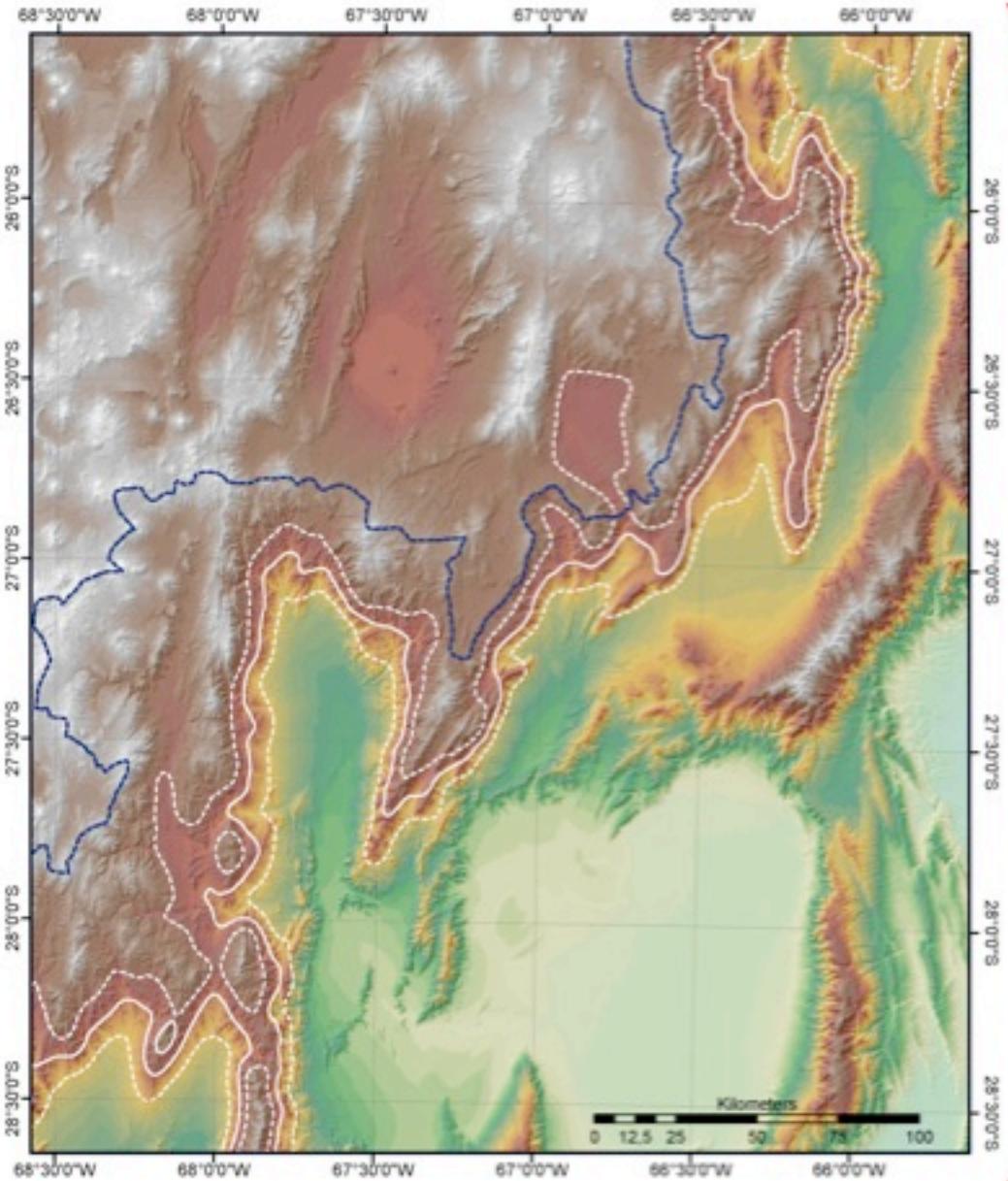


Subandes & Eastern Cordillera



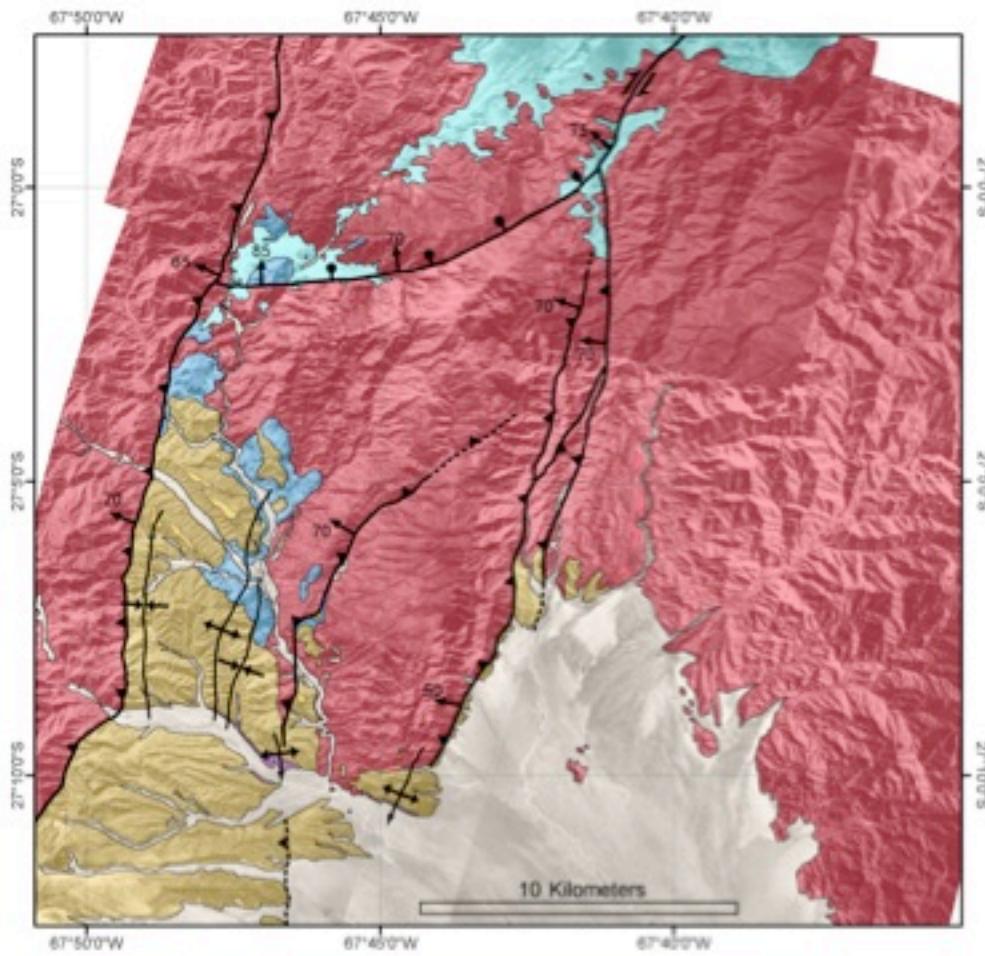
Altiplano-Puna Plateau



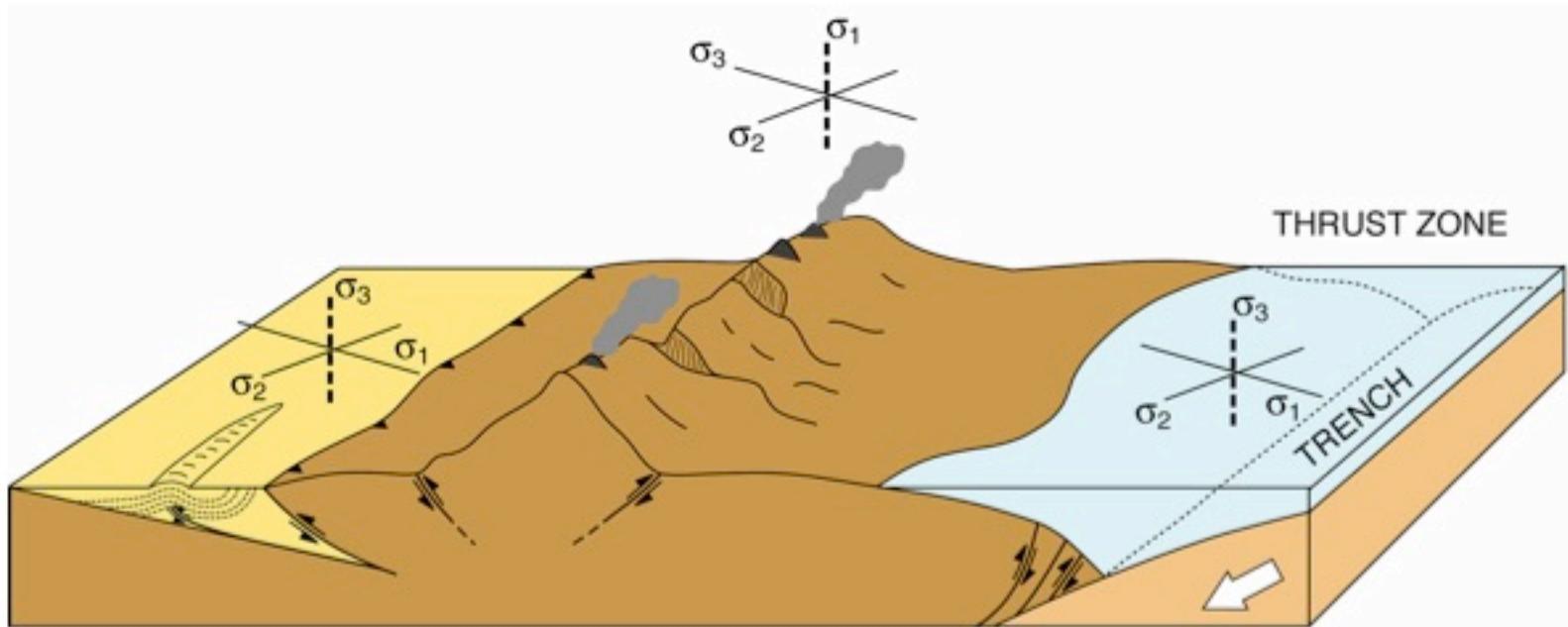


- Southern Puna Plateau
 - Region of internal drainage
 - 3000 +/- 500 m contour



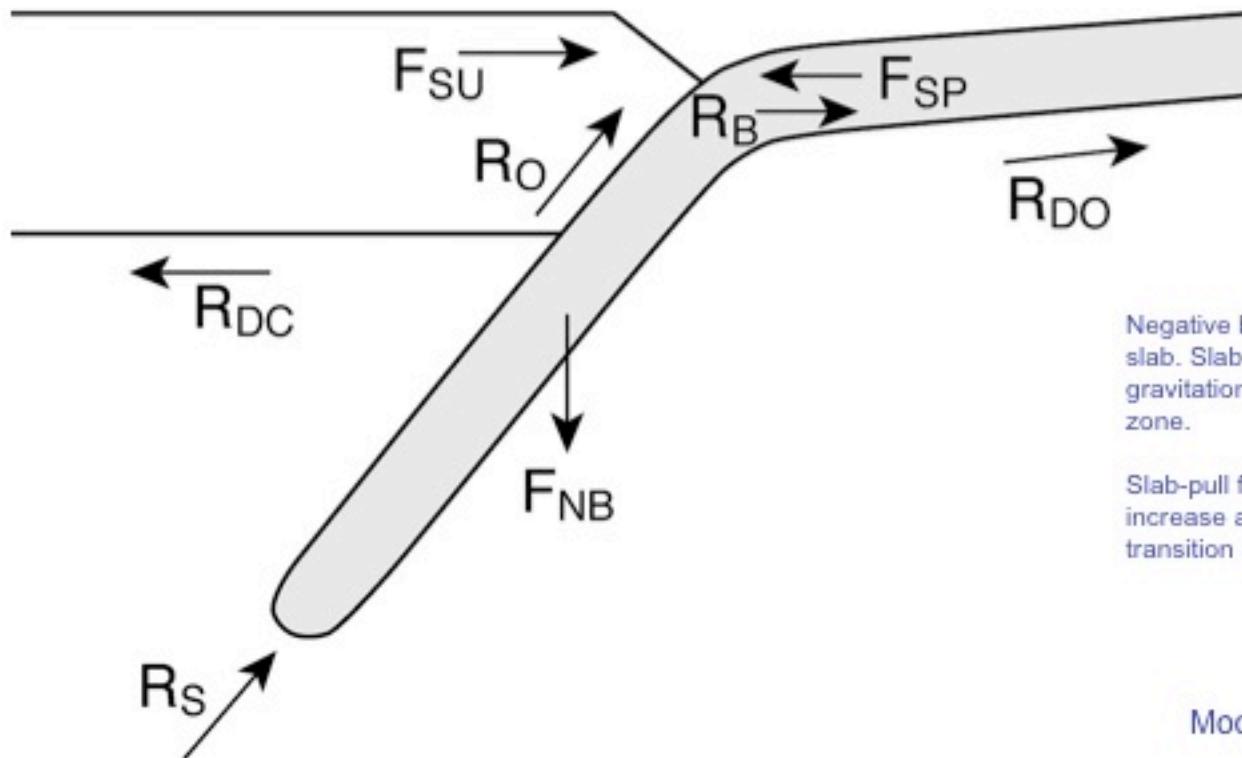


Schoenbohm and Strecker, in prep.



Nakamura and Uyeda, 1980

Slab-pull force

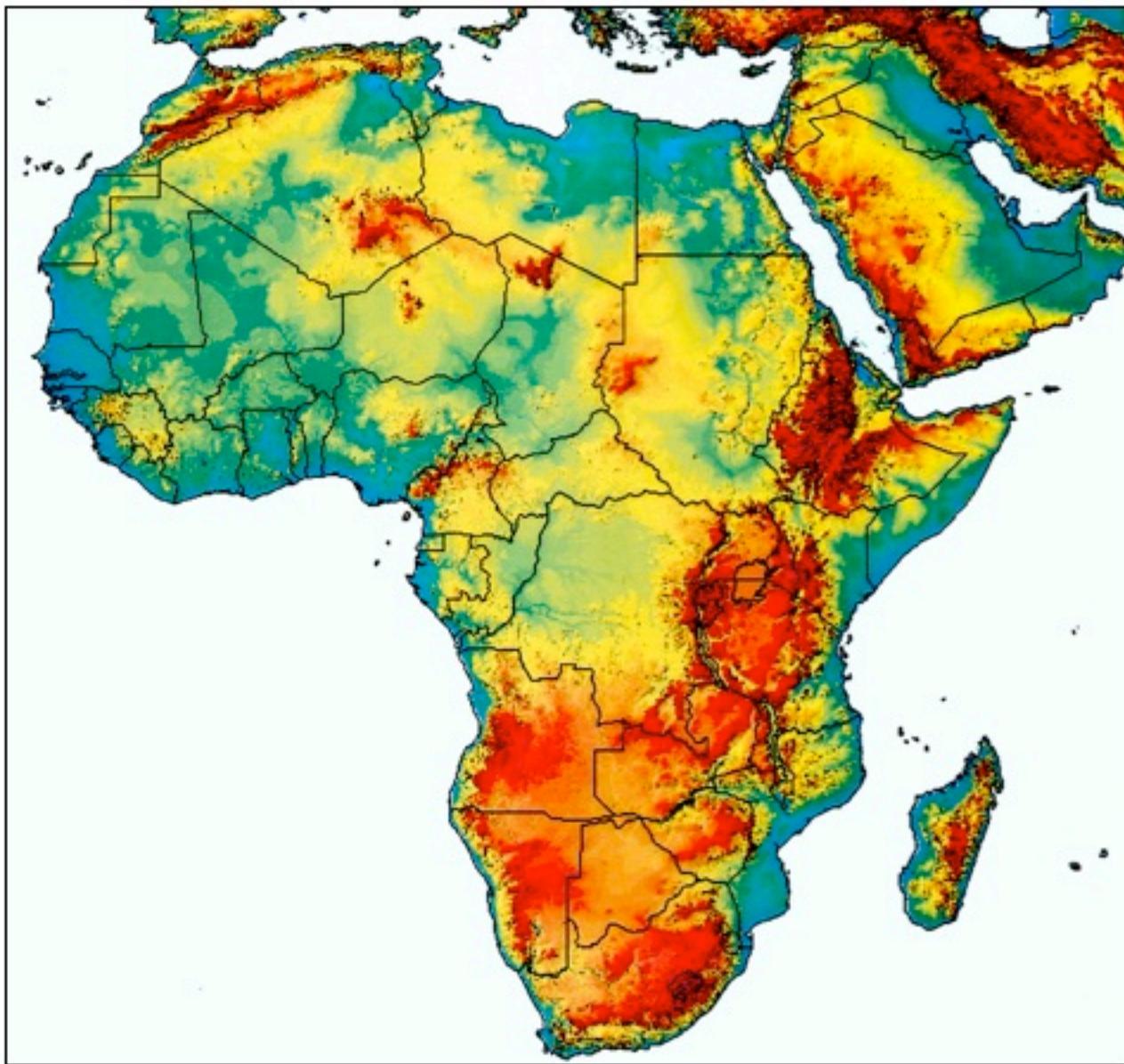


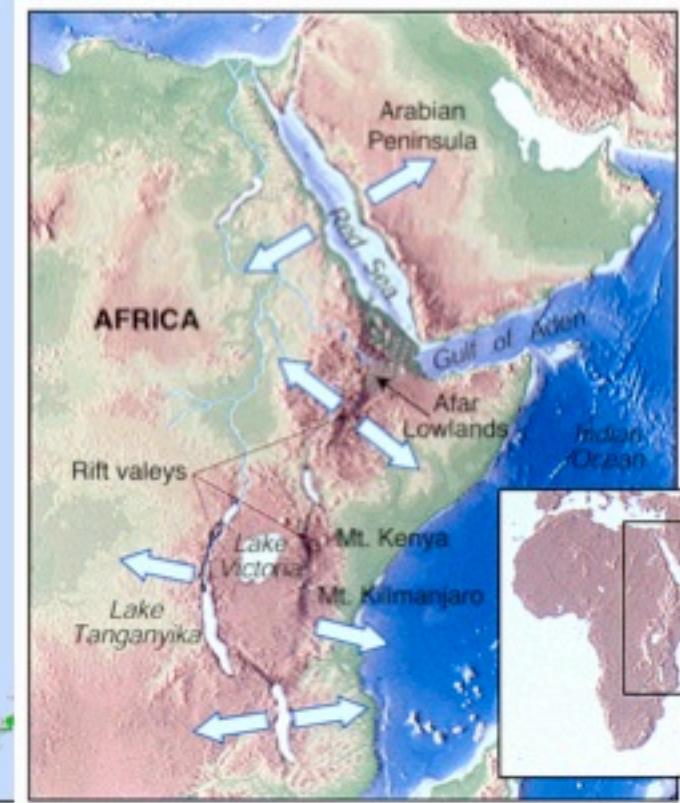
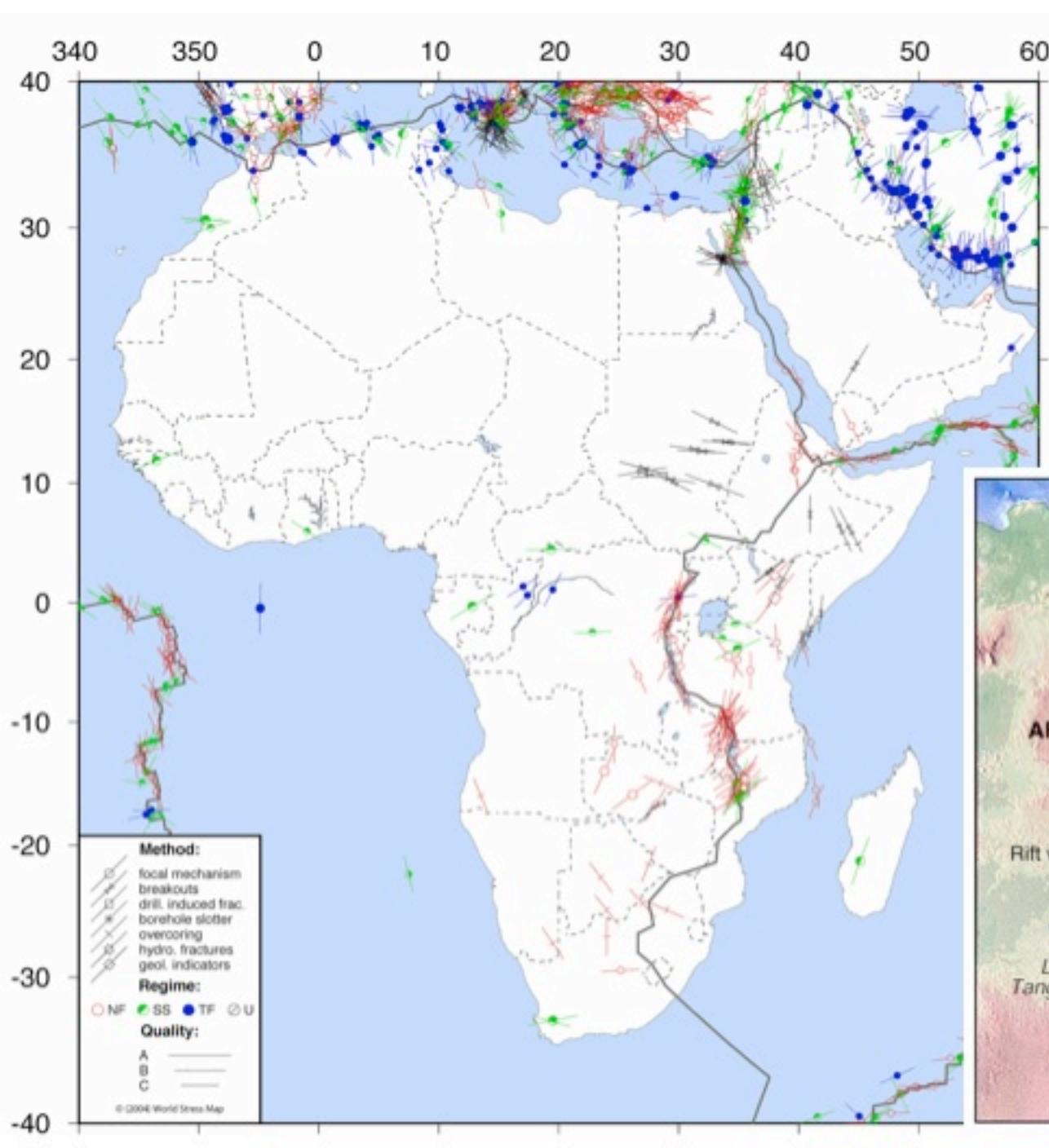
Negative buoyancy pulls dense oceanic slab. Slab pull thus results from a gravitational force in the subduction zone.

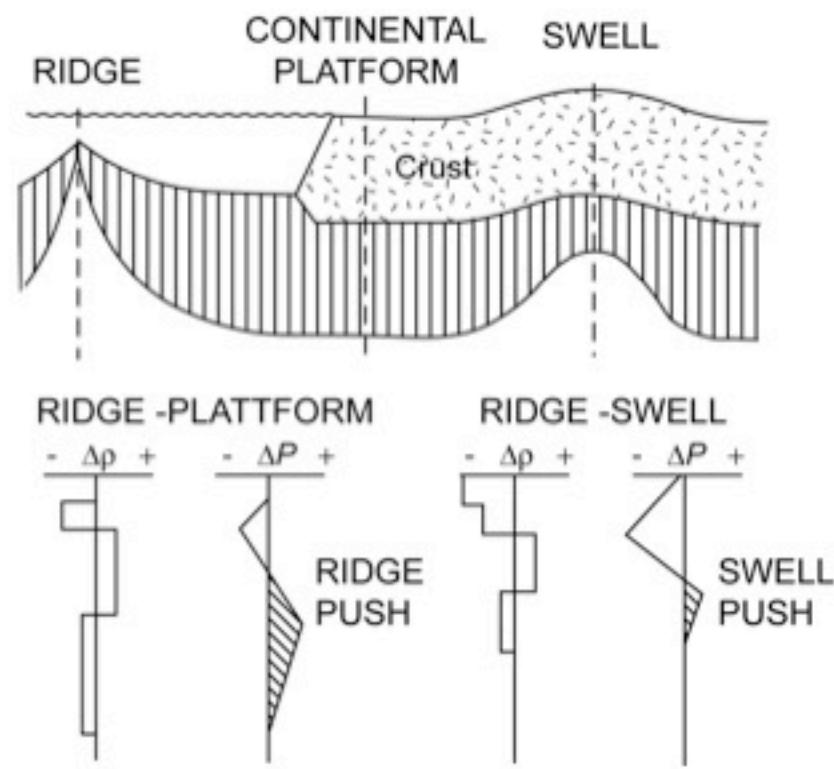
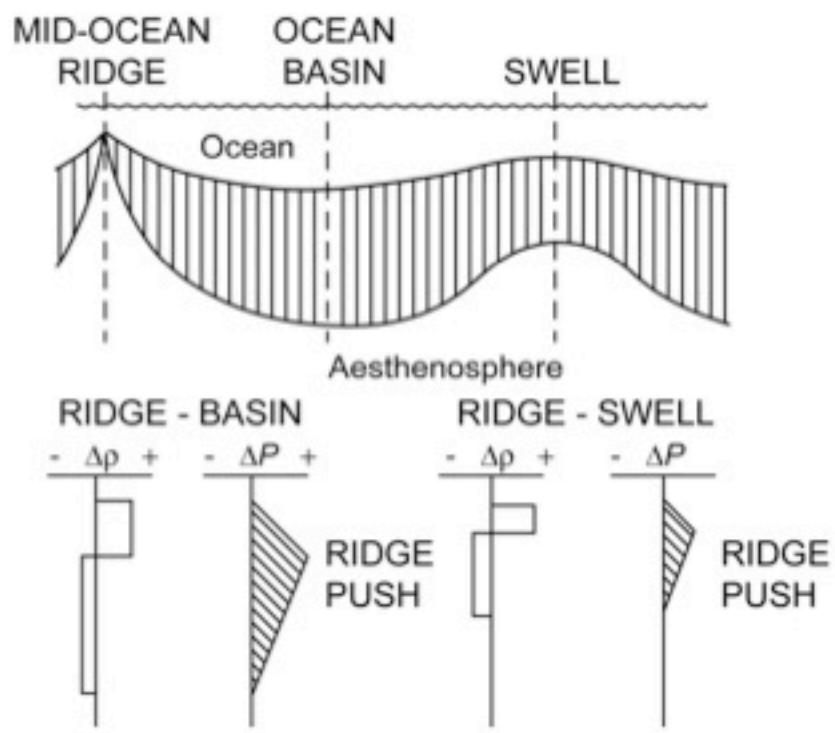
Slab-pull force is increased by density increase at depth due to olivine-spinel transition at 400 km

Moores and Twiss, 1997

Africa







(3) Neotectonics in intraplate regions

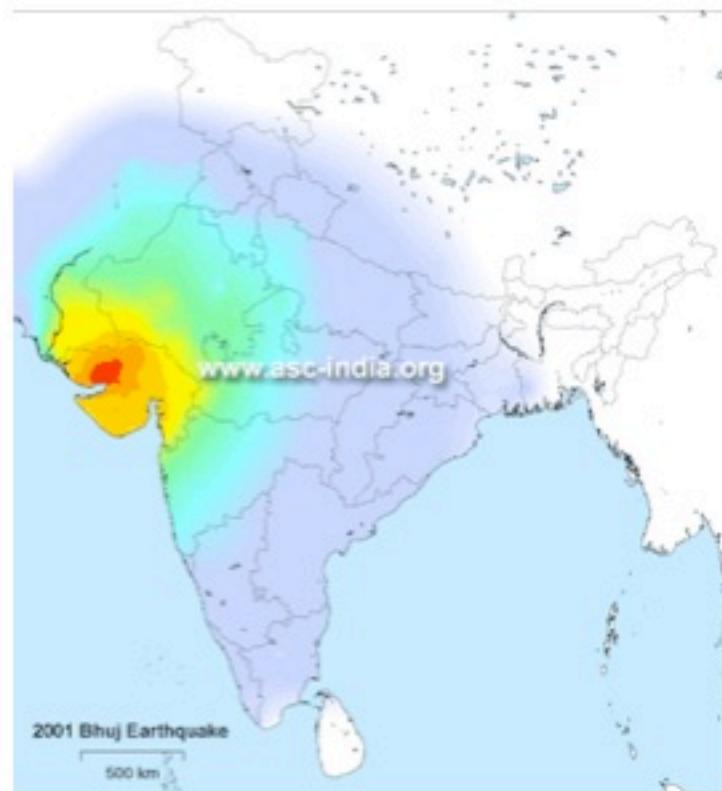
Worldwide seismic hazard



Highest seismic hazard

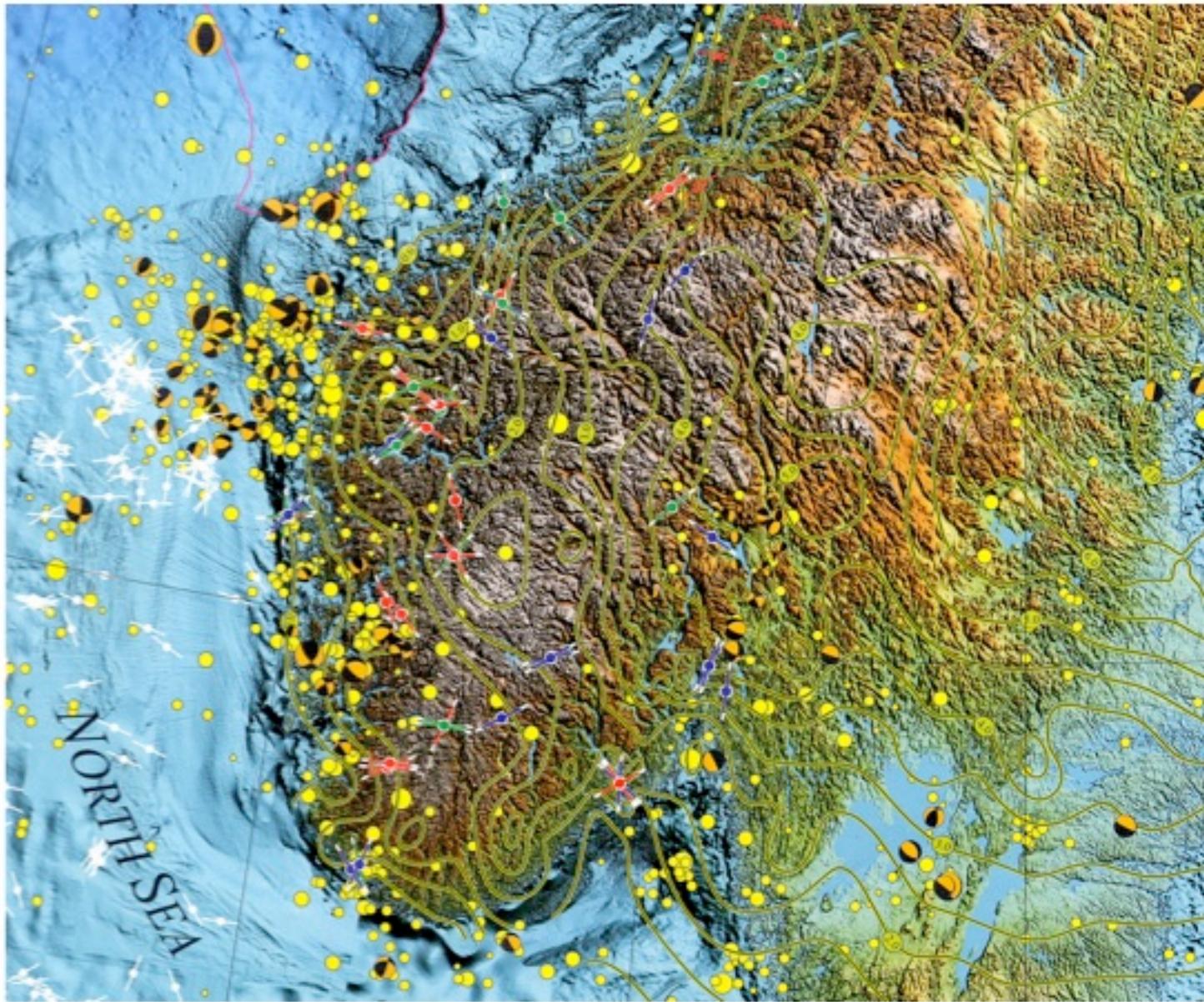


Lowest seismic hazard



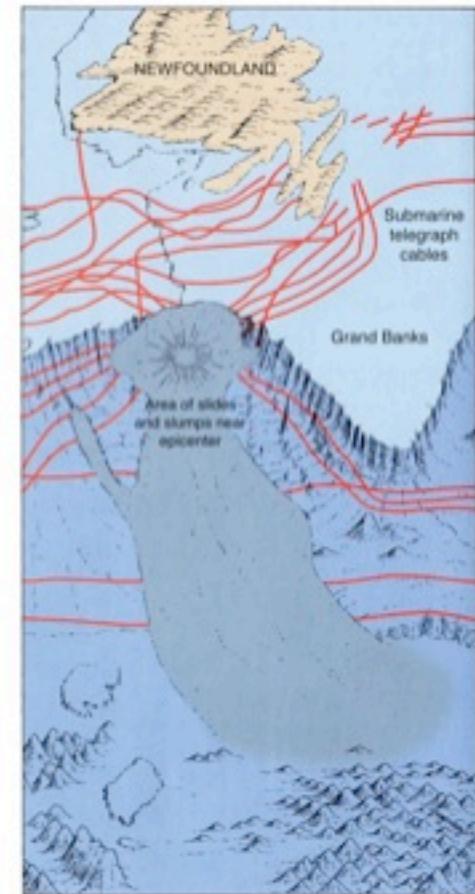
Gujarat Bhuj earthquake, 2001

Seismicity in Scandinavia

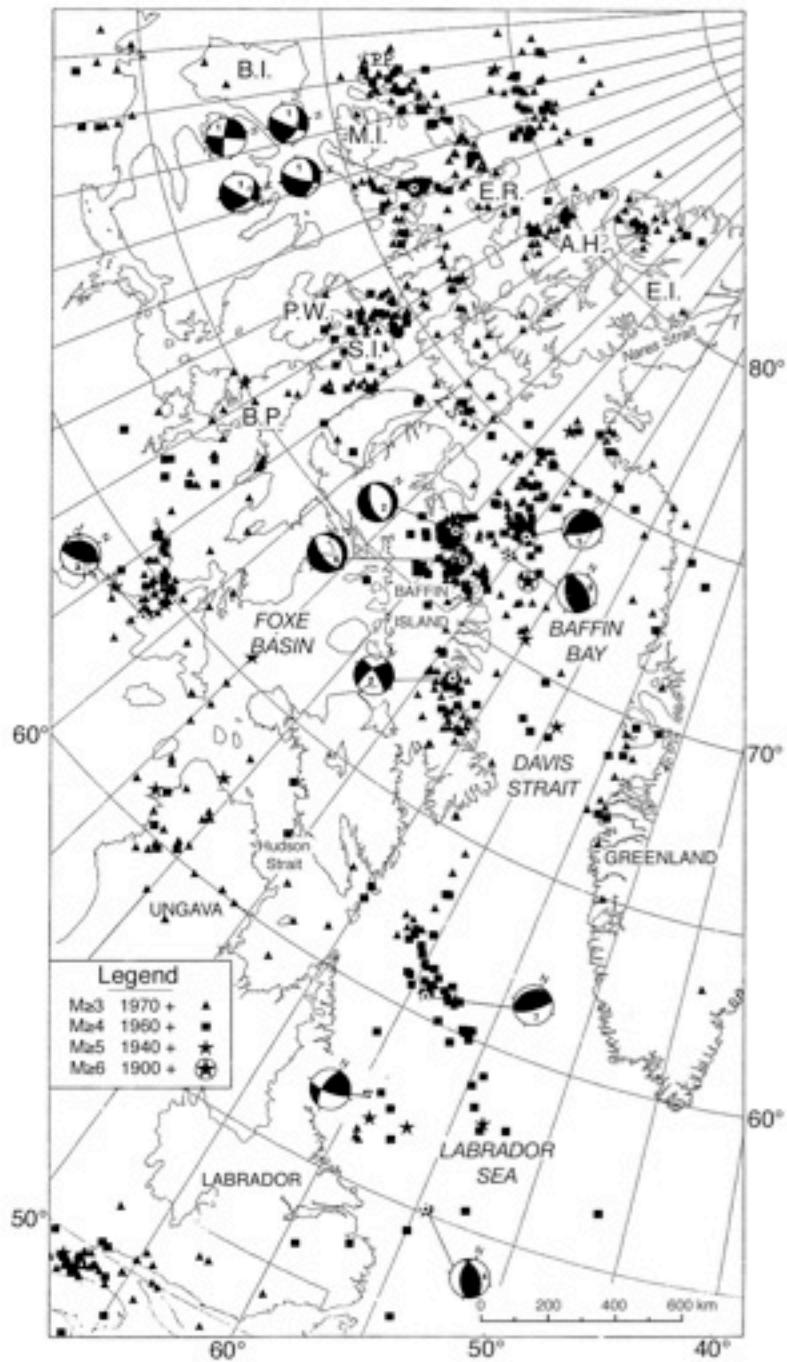


Neotectonic map, Norwegian Geological Survey

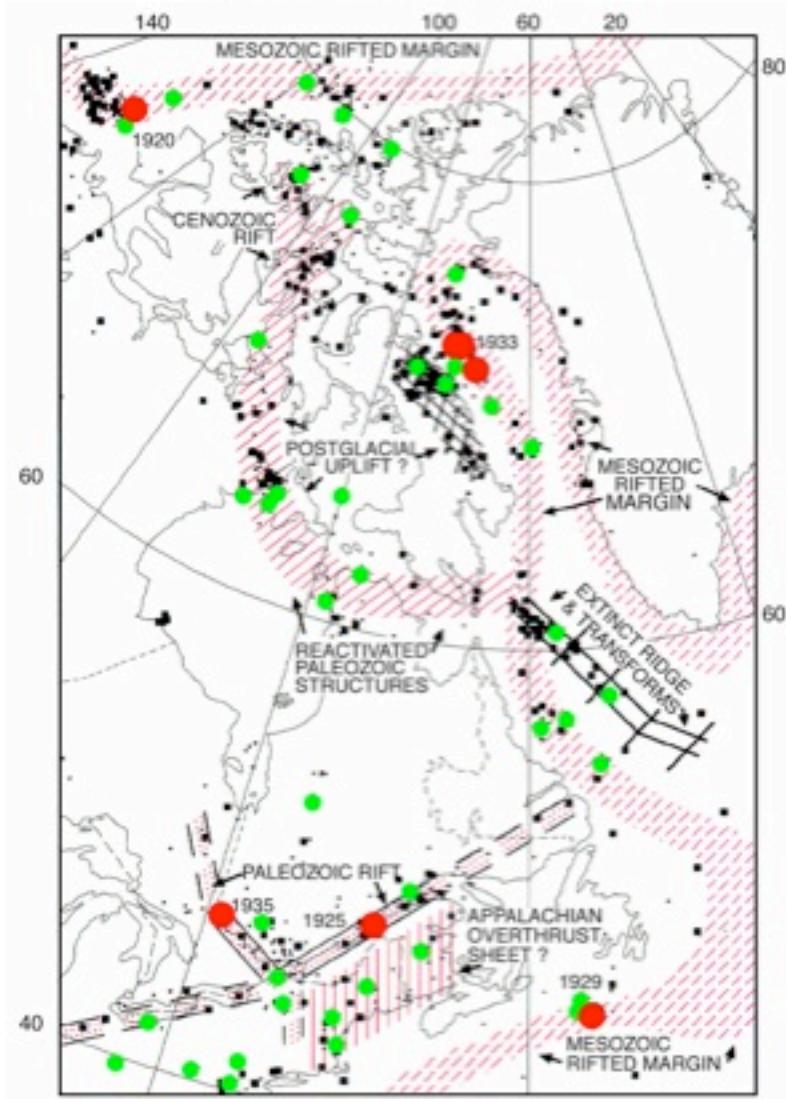
Seismicity in the St. Lawrence lowland and New Foundland, Canada



Adams and Basham, 1991



Seismicity in NE Canada

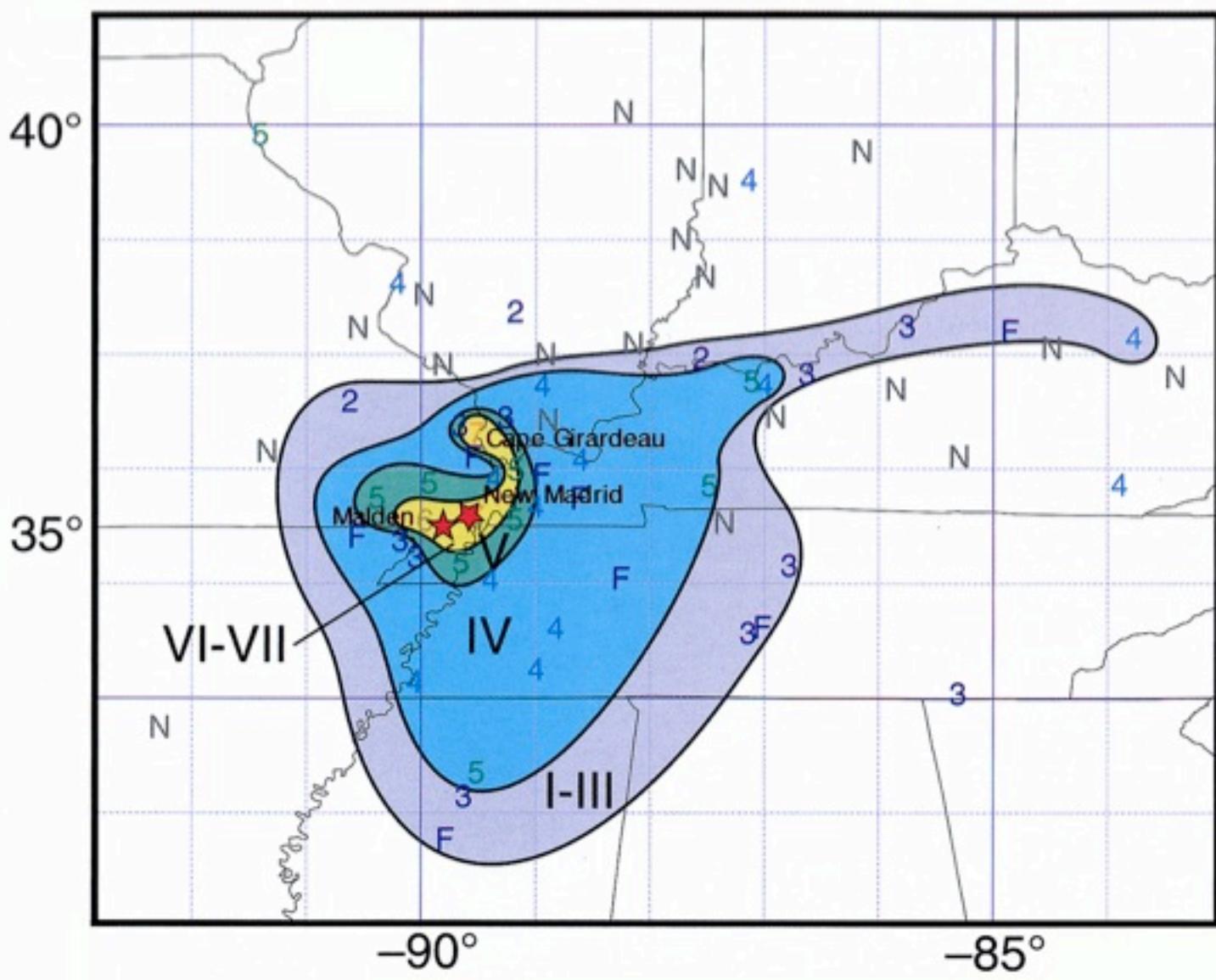


Adams and Basham, 1991

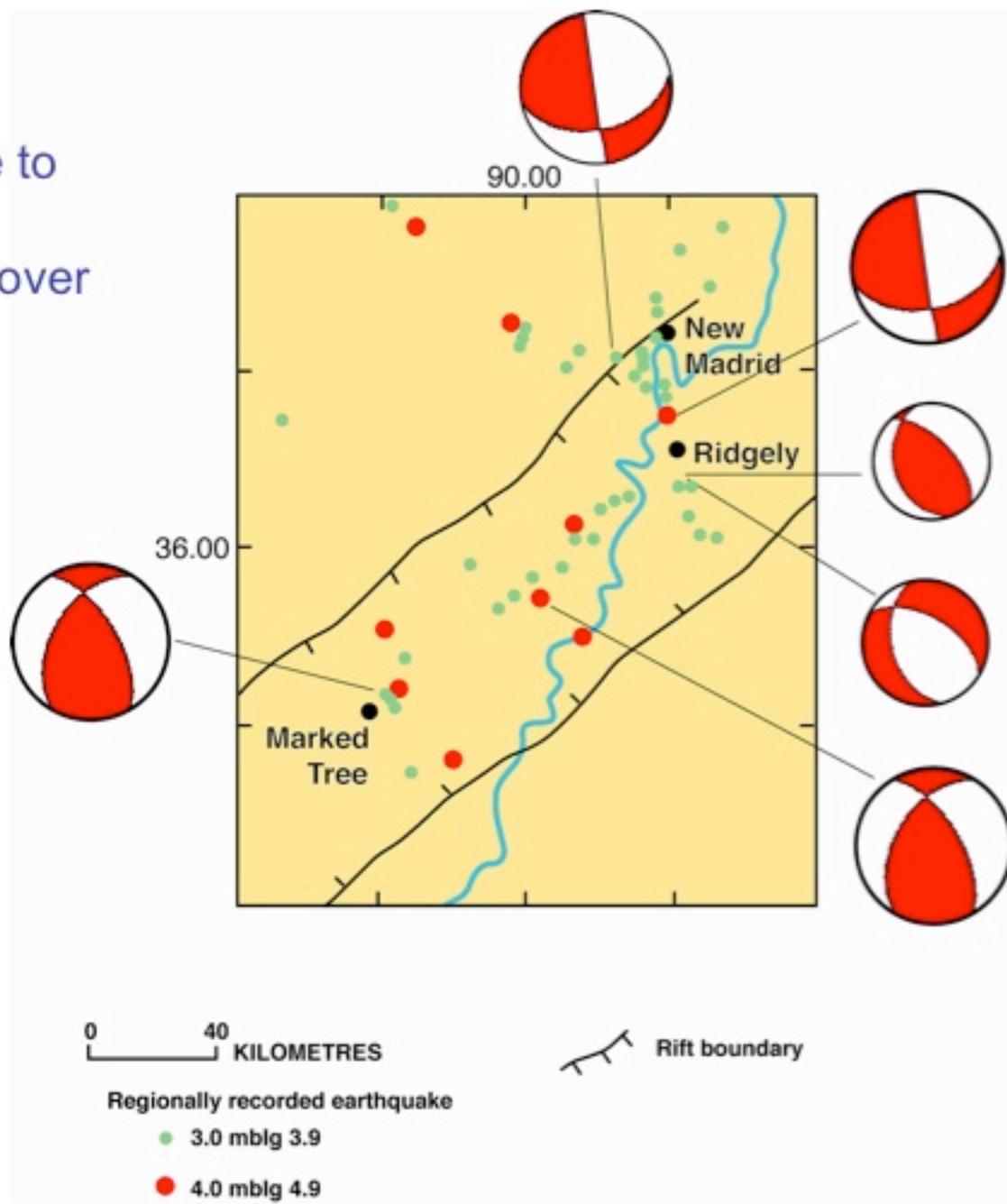
Neotectonic deformation and seismicity in the Reelfoot Rift, New Madrid seismic zone



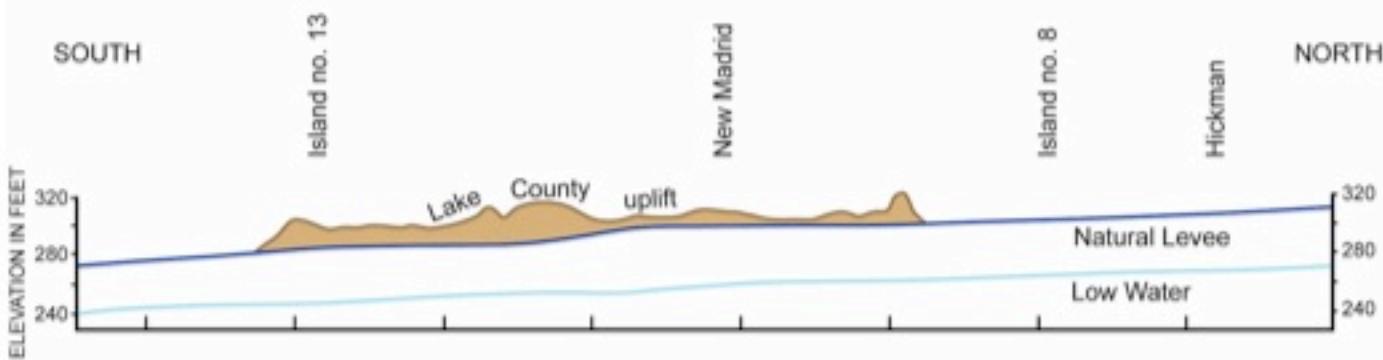
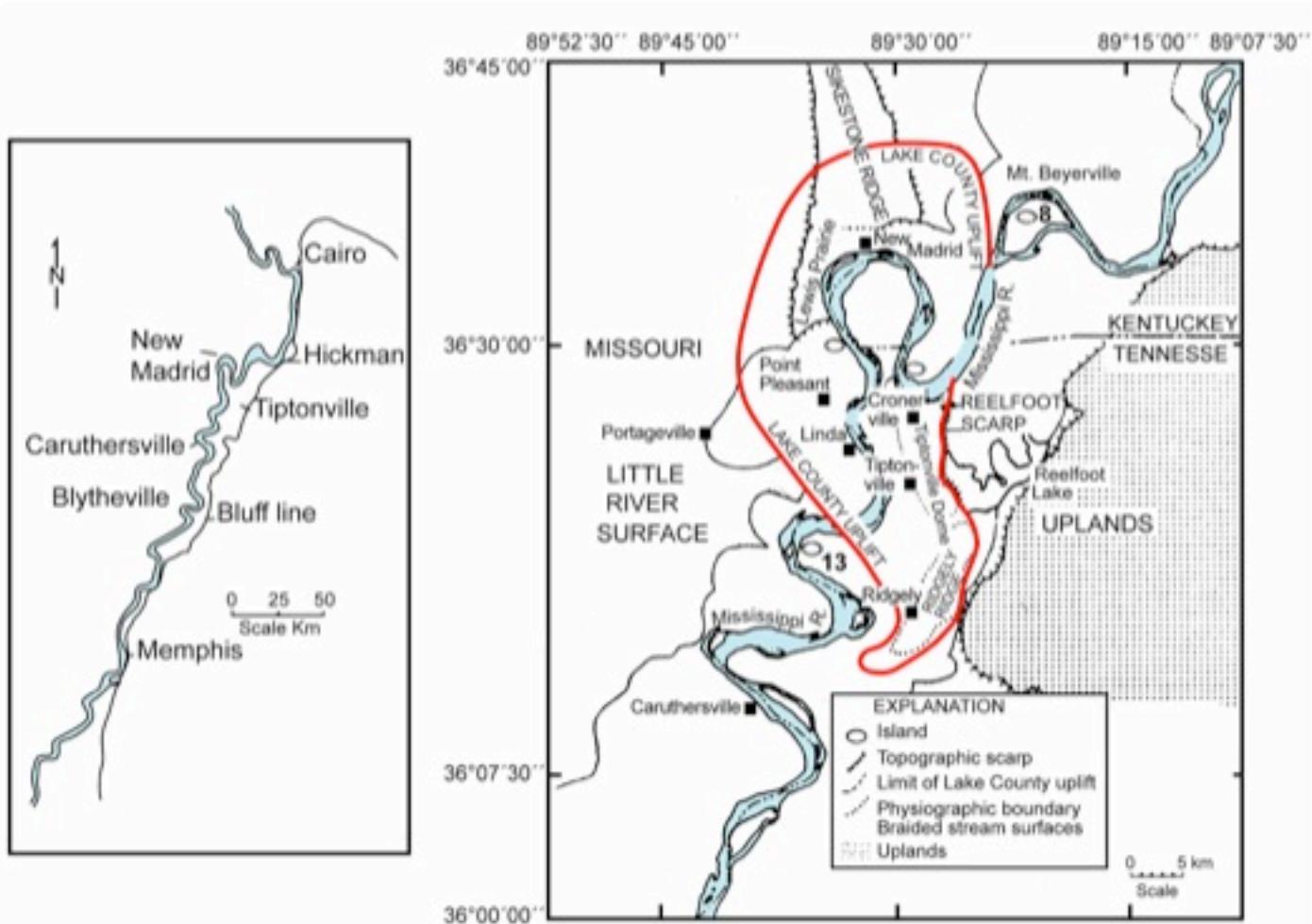
New Madrid earthquakes, 1811, 1812



Earthquake occurred due to dextral slip on two right-lateral faults in a left stepover







Slope deformation

River adjustment

Uplift



Profile

Pattern

bar-braided
or
meandering-talweg
braided

aggradation

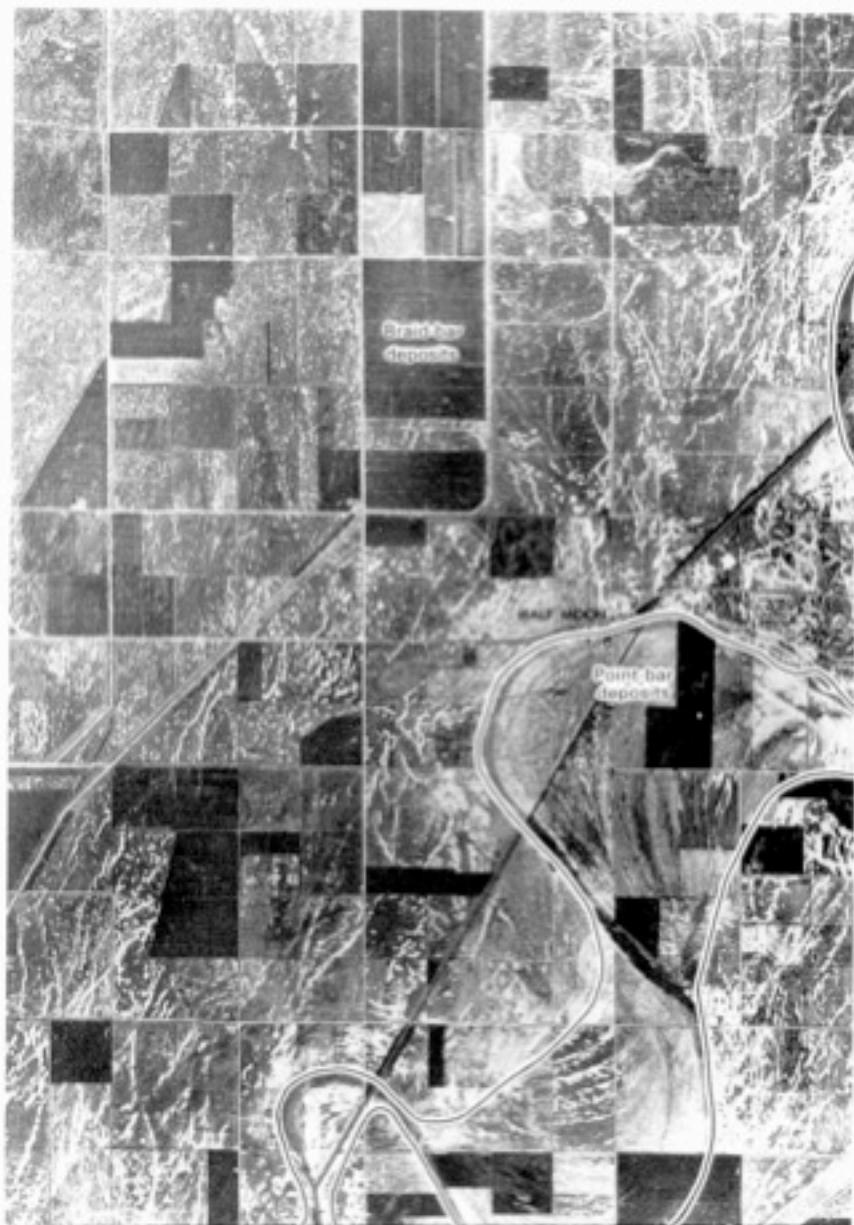
degradation

aggradation



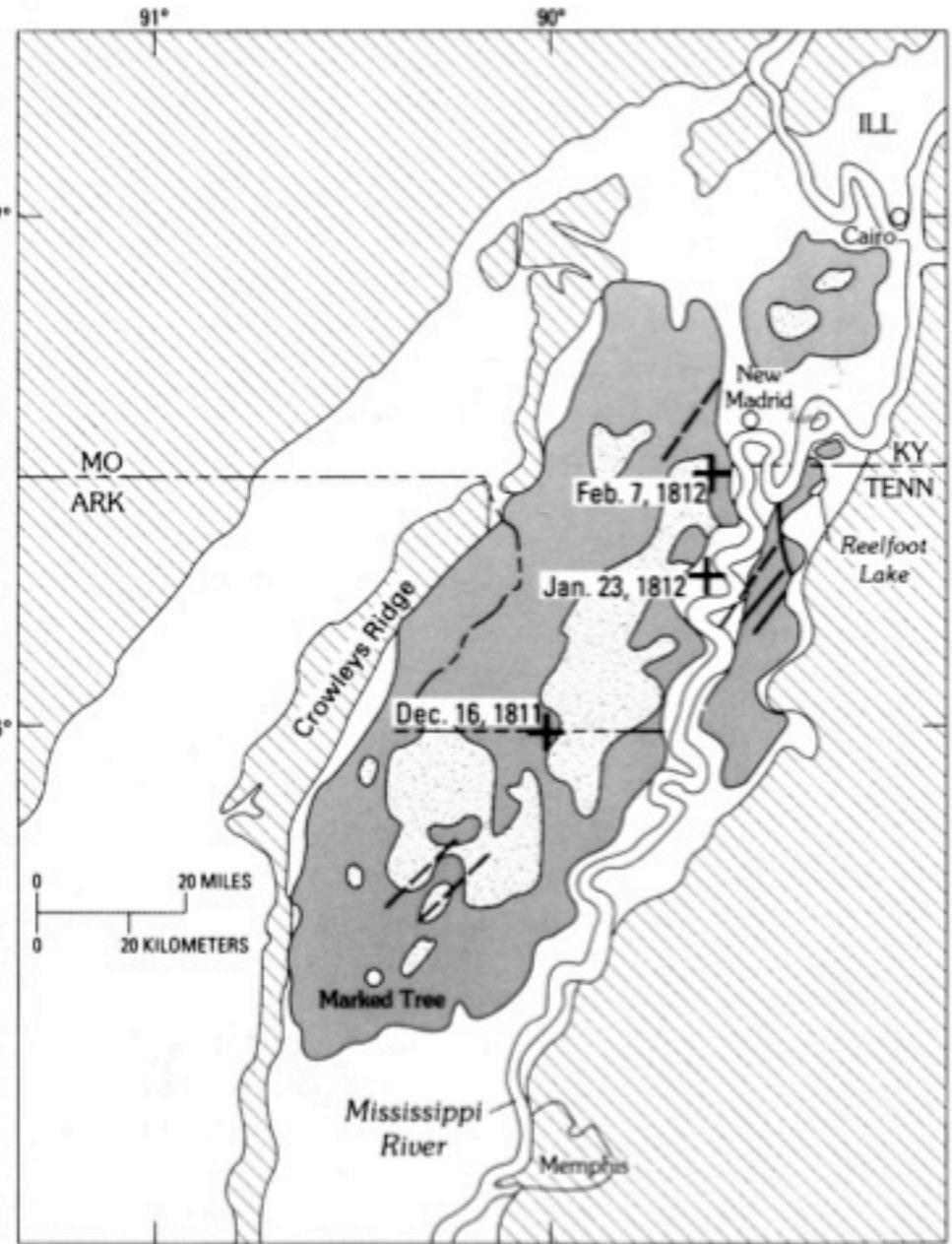
terraces

bar-braided



sand blows and
dykes





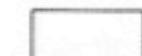
EXPLANATION



Upland areas



Alluvium with more than
25 percent of ground surface
covered by sand-blow deposits



Alluvium with few or
no observed sand blows



Alluvium with recognizable
sand-blow deposits (>1 per-
cent ground coverage), but
not major ground coverage



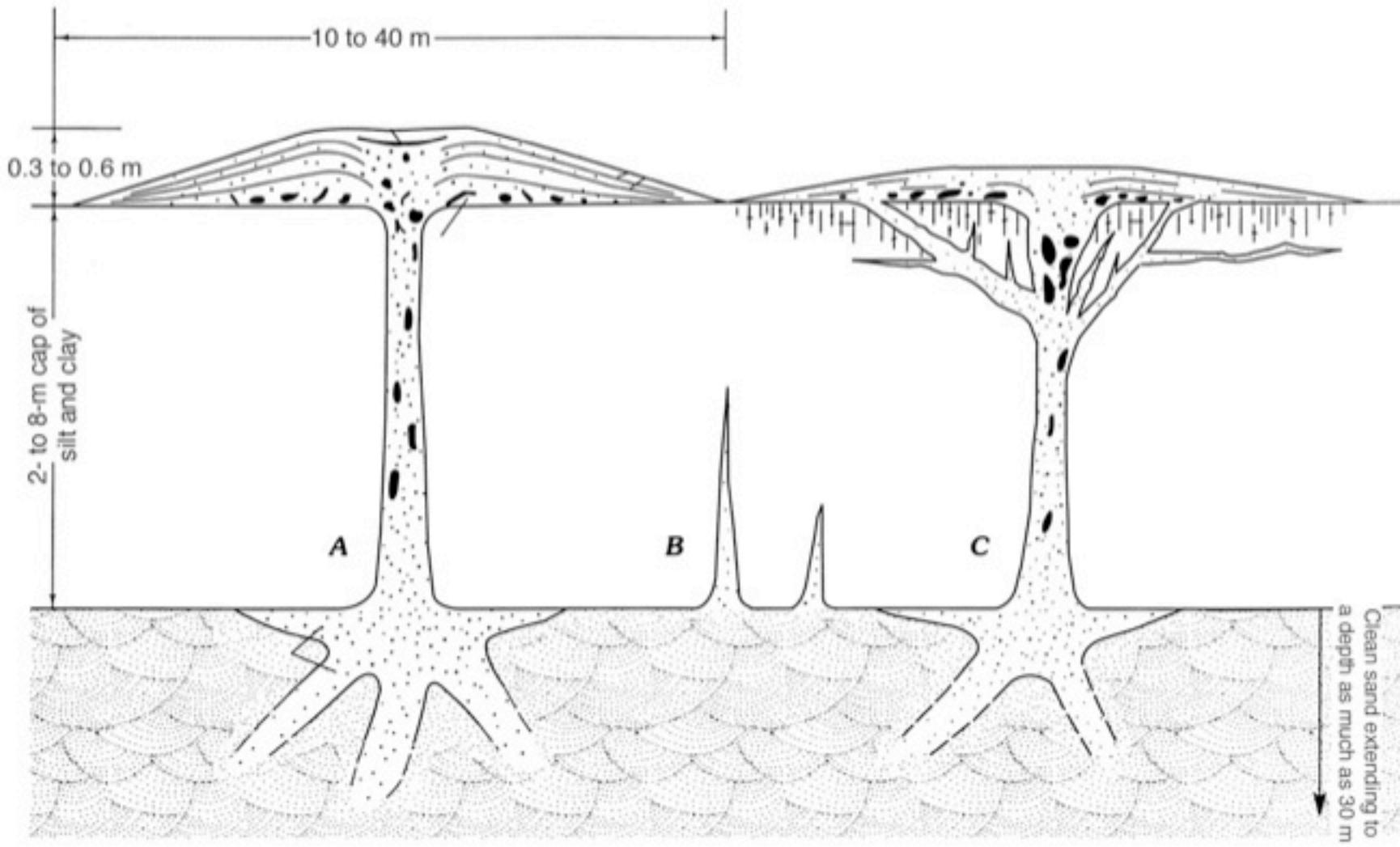
Epicenters for three largest
1811-1812 earthquakes
according to Nuttli (1979)



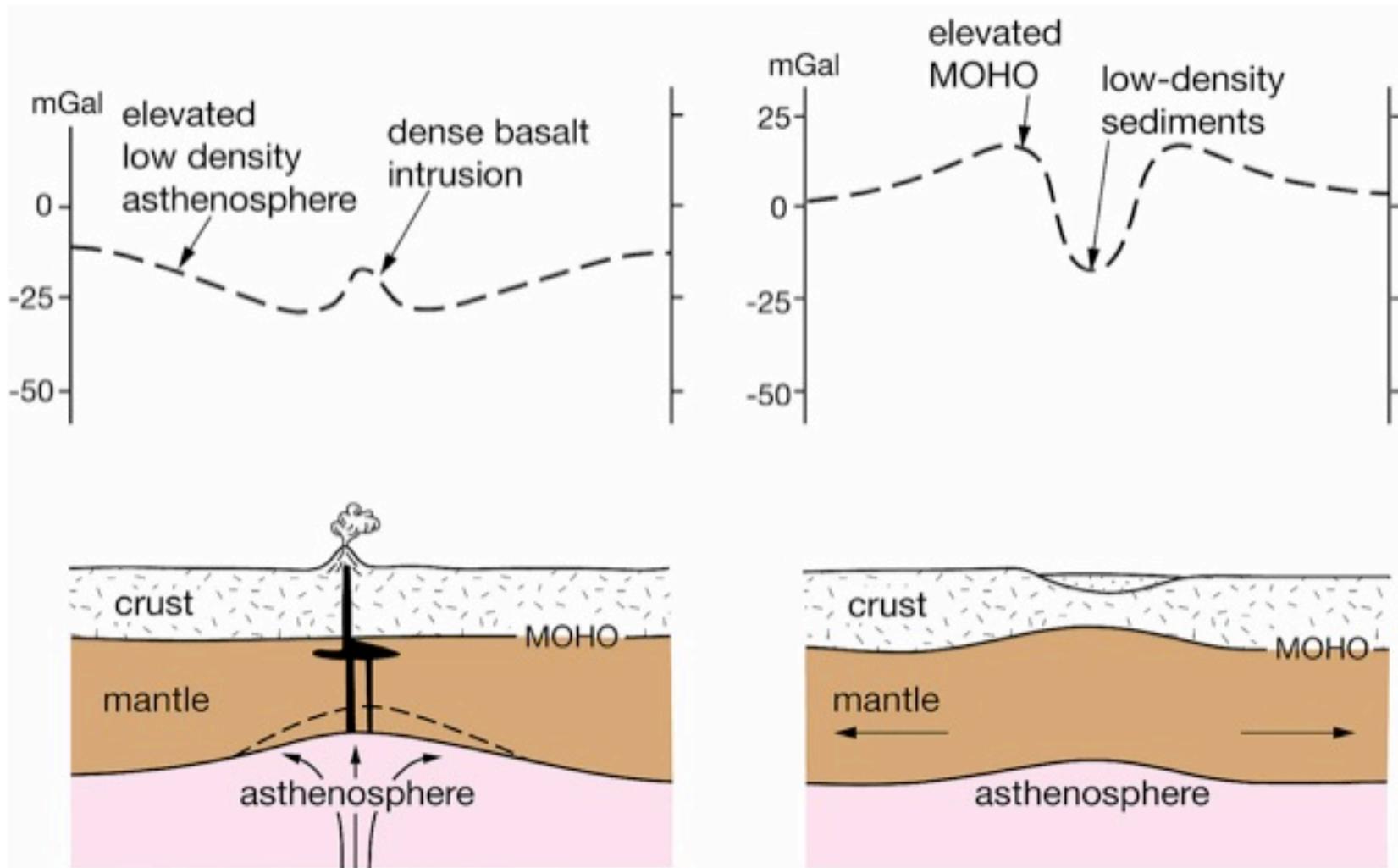
Fault



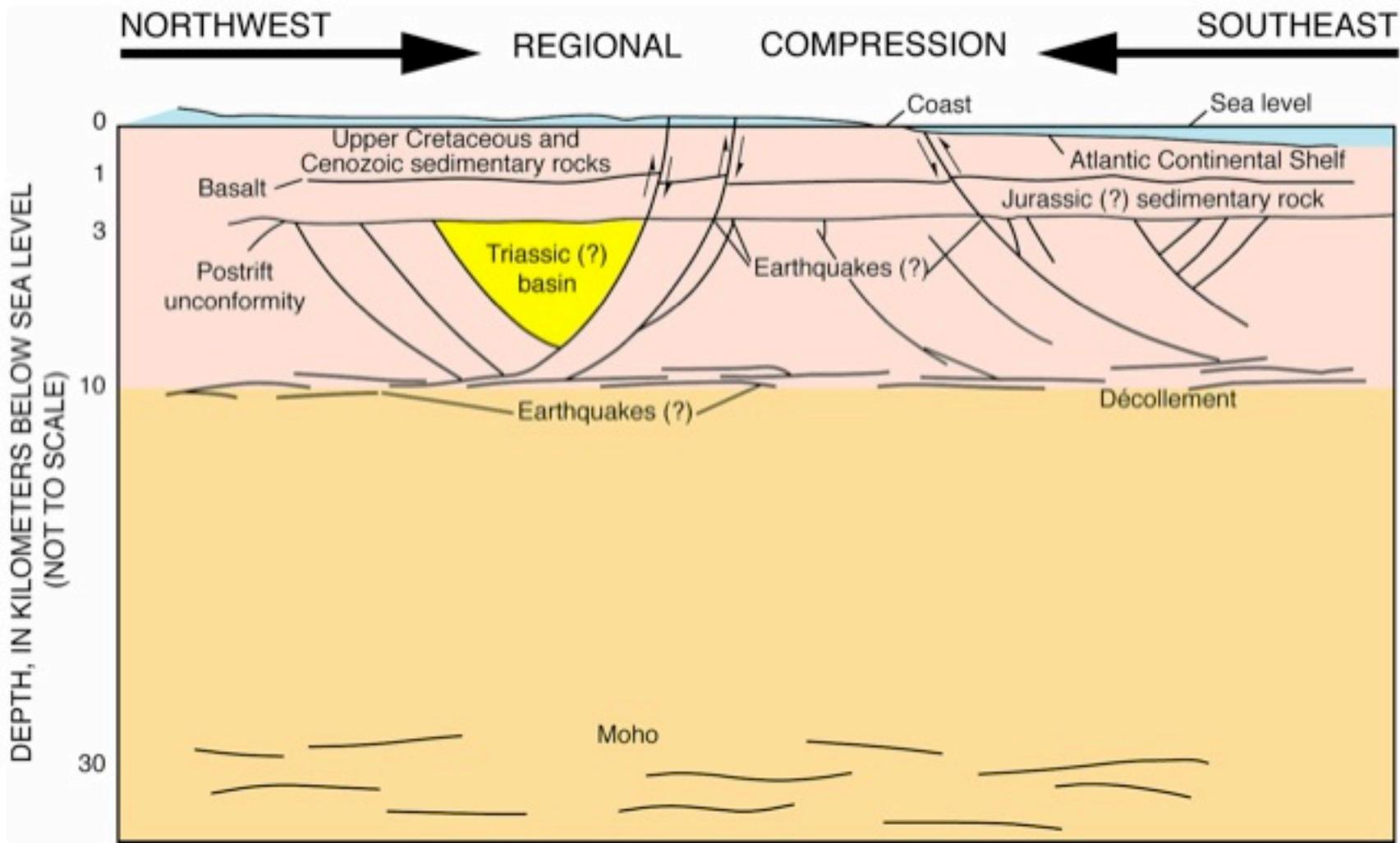
Fault zone



Seismicity in continental interiors and paleo-rifts



Seismicity along the U.S. East coast



Gravity signatures of paleo-rifts

